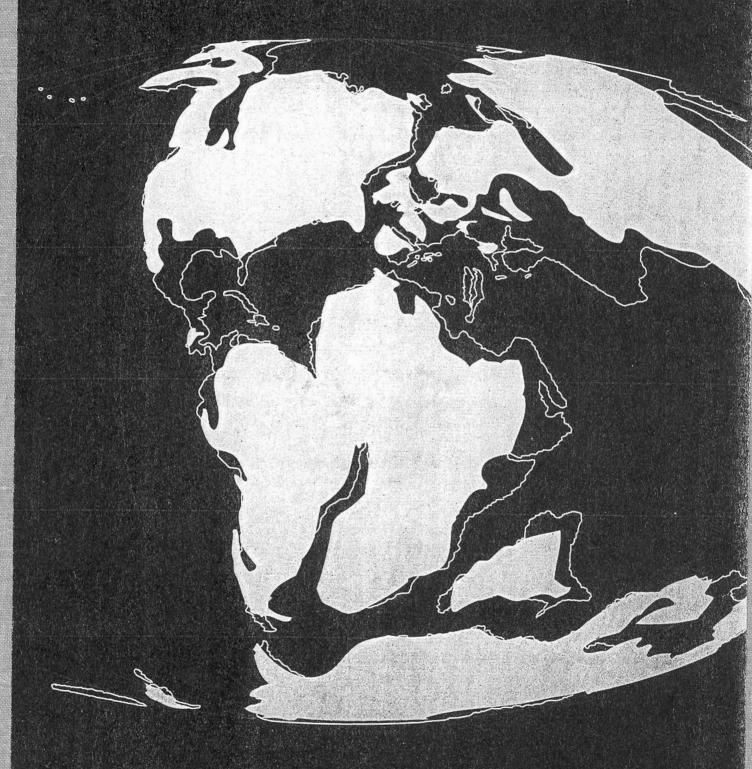
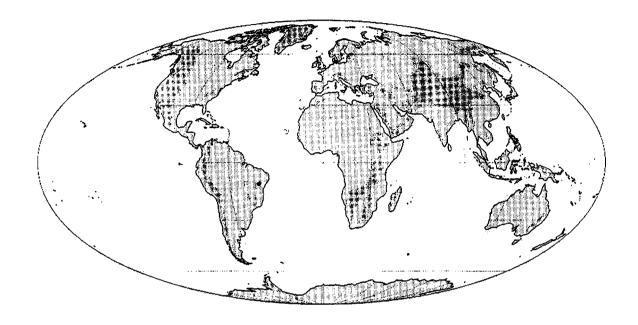
ATLAS OF MESOZOIC AND CENOZOIC COASTLINES



LAN G. SMITH, DAVID G. SMITH and BRIAN M. FUNNELL

The inferred positions of global paleoshorelines through the 240 million years of the Mesozoic and Cenozoic are presented within this atlas. Thirty-one maps, generally corresponding to stratigraphic stages, provide a snapshot of the continents and their shorelines at approximately 8 million year intervals. The maps provide a representation of the gross changes in the distribution of land and sea throughout the Mesozoic and Cenozoic plotted on Mollweide projections of paleocontinental reconstruction. They do not distinguish between well and poorly defined shorelines, but the information sources are set out in a bibliography numbering more than 2000 primary paleographic references. Work up to 1980 was used in the initial compilation, and the final maps use additional unreferenced and unpublished data up to about 1987.

This is a unique global compilation that presents the first attempt at delineating global shorelines at stage level.



Present-day coastline map.

ATLAS OF MESOZOIC AND CENOZOIC COASTLINES

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Preface

The maps in this atlas originated in concurrent Extramural Research Awards by British Petroleum from 1981 to 1985 to the Department of Earth Sciences at the University of Cambridge and the School of Environmental Sciences of the University of East Anglia. The group at the University of East Anglia (B. M. Funnell, R. V. Tyson and T. Austin) compiled paleocoastlines and paleofacies on present-day base maps of the individual continents, from published paleogeographic maps and other data. The Cambridge group (A. G. Smith, R. A. Livermore, L. A. Rush, J. L. Tarry) prepared the present-day coastlines, bathymetry, geographic grids and sutures, assessed the data for making global reconstructions and developed the ATLAS map-plotting and paleocontinental reconstruction software and associated databases. The Global Paleoreconstruction Group at BP Research (initially under I. R. Hoskin, then C. P. Summerhaves, then D. G. Smith), drawing on the work of both university groups, developed their own GLOBE computer reconstruction software with which regional plate tectonic studies were conducted and global reconstruction atlases were developed for BP's internal use. The paleocoastlines in the present atlas were initially compiled as an internal report. The BP group also supported the Paleogeographic Atlas Project at the University of Chicago, under Alfred Ziegler, and their global data compilations were also drawn upon at various stages of this work.

The maps presented here represent very many man-years of work, and we gratefully acknowledge the following for their input at various stages: Peter Llewellyn, Richard Field, Ian Hoskin and Andrew Fleet, who were responsible for setting the project up in the former Stratigraphy Branch of BP Research's Exploration and Production Division; David Burwell, Raymond Jackson and Brian Biart, who wrote and maintained the GLOBE continental reconstruction and database software; Neil Parkinson, Helen Turnell, Michael Daly and Martin Heffernan, who were all closely involved with the regional and global studies with BP's Global Paleoreconstruction research group that contributed to the reconstructions herein. Several generations of geology students were employed as temporary technical assistants at BP Research and made valuable contributions. BP Exploration, through Michael Daly, provided finance for the editing in Cambridge of the bibliography and maps by Clare Glover and Graham Cocksworth. Throughout the project there was considerable help from secretarial staff in all three organizations. Finally, we are grateful to BP International plc for permission to publish these maps.

Introduction

This atlas is one outcome of a substantial research programme that was conducted under the sponsorship of British Petroleum International. BP established concurrent research projects over a four-year period, 1981–1985, at the Universities of Cambridge (Department of Earth Sciences) and East Anglia (School of Environmental Sciences). The aim was to produce up-to-date computer-drawn maps on which were plotted global paleogeographic data at the stratigraphic resolution of one map for each Stage for most of Mesozoic and Cenozoic time

The Cambridge project, under the direction of Alan Smith, drew on that department's expertise, dating back to the work of Bullard, Everett and Smith (1965) and including earlier collections of maps by Smith, Briden and Drewry (1973), and Smith, Hurley and Briden (1981). They developed the ATLAS reconstruction software, initially on a mainframe computer but more recently on a PC, and prepared the large amounts of digitized data that make up the continental fragments, including present-day coastlines, bathymetric contours, fragment boundaries, and intra-continental sutures.

The University of East Anglia project was responsible for the paleogeographic and paleoenvironmental data to be used on the resulting global reconstructions. R. V. Tyson conducted a systematic search of the recent geological literature, back to a cut-off date of about 1960. Rather than attempting to collect primary data, the search was for data already compiled into paleogeographic maps, whether at local, regional or global scale. As far as possible, the maps were sought relating to a stratigraphic interval no longer than a Stage. The search resulted in maps from about 2000 sources, which are listed in the Bibliography. The Stage was chosen as the appropriate stratigraphic unit because it is the smallest chronostratigraphic division that can usually be recognized globally. The information from these maps was compiled, a Stage at a time, on to regional maps, generally covering a continent each, and using a present-day geographic base. The 31 resulting maps are the first known attempt to draw a continuous series of global maps with such a fine resolution. In some cases, especially for the Triassic, adjacent Stages have had to be combined because of lack of data, but the average duration of the stratigraphic units for the whole series is about 8 million years (Ma) (31 maps covering 245 Ma).

Meanwhile, the research group at BP developed their own paleoreconstruction software (GLOBE), with an associated database from which lithofacies data could be plotted directly on to paleocontinental reconstructions. This system was used to conduct a series of regional plate tectonic studies for BP Exploration, on regions that included the Arctic (Smith 1987), northern Australia/Papua New Guinea, the Caribbean, Southeast Asia (Daly et al. 1991), northwest Europe, East Africa/Arabia, the South Atlantic and New Zealand. Originally for internal company use, the results of all of those studies contributed to the development of BP's in-house global reconstructions and thus to the maps in this Atlas. Although initially based on the Cambridge reconstructions, BP's global reconstructions inevitably came to diverge in some respects from those of ATLAS, and the reconciliation of the two is described below.

The final stage in BP's involvement with this project, postdating the conclusion of the sponsored research in both universities, was the merging of the topographic data on East Anglia's continent-by-continent compilations, first on to present-day global base-maps and then on to reconstructions appropriate to the age of the paleogeographic information. The resulting map series was transferred to Cambridge for reconciliation with the ATLAS reconstructions, and computer-plotting of the maps as they appear in this Atlas.

Global shoreline maps for the Cretaceous, based on the original 1986 versions, were published in 1990 (Funnell 1990). More detailed Cretaceous shoreline maps for Europe only, based on the 1986 continent by-continent facies and shoreline maps, were first published in 1987 and republished in 1990 (Tyson and Funnell 1987; Funnell 1990).

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Global reconstructions

2.1 Making a continental reassembly

The method of making global continental reconstructions is now routine. It has been fully described elsewhere (e.g. Smith, Briden and Drewry, 1973; Smith, Hurley and Briden 1981) and only an outline is given here.

The first step is to subdivide the present-day world into the continental fragments that existed at the time concerned. We use the word 'fragment' because the various pieces to be reassembled do not all coincide neatly with individual continents. Each continental fragment consists essentially of a digitized outline, each point on which must be moved to a different position for each new reconstruction. As well as the outline, other digitized data may also be associated with the fragment, and moved with the outline. Thus, we normally include the present-day coastline, not for any paleogeographic significance, but simply in order that the fragment may be readily identified. Continental fragment boundaries rarely coincide with present-day coastlines, but are more usually located at continent—ocean boundaries or at intracontinental sutures.

The primary source of information for making Mesozoic and Cenozoic reconstructions is the ocean-floor magnetic anomaly record. All magnetic chrons recorded in the magnetic anomalies of the Atlantic and Indian Oceans are fitted together as well as possible. The information required for the computer to transform each pair of present-day coordinates into the corresponding pair of paleo latitude—longitude coordinates is expressed as an angle of rotation and the coordinates of the pole of rotation, using Euler's principle (see Cox and Hart 1986, for a comprehensive account of the method). We have used only published finite rotations in this work (Table 1), because the effort of reexamining the original ocean-floor anomaly data is a major research project in itself. The precision with which anomaly matches can be made has been greatly increased by using SEASAT data (e.g. Cande, LaBrecque and Haxby 1988). Modifications to the geologic time-scale have also increased the precision of the reconstructions (Harland et al. 1990). The increase in precision that could be obtained by using the most recent revision of the Late Cretaceous to Recent magnetic polarity time-scale (Cande and Kent 1992) is not significant on the scale of the maps.

Table 1. Sources of rotation data

Region	Reference
Atlantic Ocean data	
S. America to Africa	
C4N to C34N	Cande et al. 1988
'M0' (Late Aptian) to closure	Klitgord and Schouten 1990
N. America to Greenland	
C13N	Srivastava and Tapscott 1986
C21N-'M0'	Roest and Srivastava 1989
Pre-closure	Smith and Livermore 1991
N. America to Europe	
C5 to CRT	Srivastava and Tapscott 1986
Closure	Bullard et al. 1965
Indian Ocean data	
Africa to Somalia	
C5	McKenzie, Davies and Molnar 1970; Le Pichon and
	Francheteau 1978
Closure	Bullard et al. 1965
Madagascar to Somalia	Segoufin and Patriat 1981
Antarctica to Madagascar	
C3N-C5N,C7N-C8N,C21N	Cande and Leslie 1986
C6N	Sclater et al. 1981
C13, C22N, C34N	Fisher and Sclater 1983
C28N	Lawver, Sclater and Meinke 1985
Australia to Antarctica	
C5N-C13N, C22N, C31N	Stock and Molnar 1982
C18N, C24N, C34N	Mutter et al. 1985
Closure	Weissel et al. 1977
India to Australia	
C22N-C28N	Norton and Sclater 1979
C34	Fisher and Sclater 1983
M1R and closure	Norton and Sclater 1979

Where the anomaly data do not coincide in time with the age required for a particular reconstruction, the fit is interpolated between the finite rotations of the times closest to it. This procedure gives continental reassemblies for all the major continents back to the age of the oldest ocean-floor (mid-Jurassic). The continental reassembly of earliest Jurassic time is assumed to have been the classic Pangea obtained by fitting the continental edges together, but the Triassic maps have been modified to take into account Triassic extension and motion

along a postulated transform zone (Smith and Livermore, 1991) that makes the Early Triassic reconstruction very similar to that of Van der Voo and French (1974) as modified by Walper (1980).

Where triple junctions have existed between three plates, such as during part of the opening of the North Atlantic Ocean between North America, Greenland and Europe, the reassembly can be made in two ways. For example, Greenland can be reassembled relative to Europe, and the Europe–Greenland cluster can be moved to North America. Alternatively, Greenland and Europe can be independently brought to North America. The routes chosen were considered at the time to have the best-known rotations but in the future new data may suggest a different route.

2.2 Sources of finite rotation data

Table 1 summarizes the sources of the data used for reconstructing pairs of continents.

2.3 Areas affected by orogenesis

The major uncertainties in the maps are the positions of those continental fragments that have undergone compressional deformation at a later time than that for which the map is required. For example, all continental fragments in the Alpine–Himalayan chain have been affected by Cenozoic deformation. Their positions in pre-Cenozoic reassemblies are the least well known of all the continental fragments.

There have been no major new insights into how such fragments can be repositioned in a reassembly, but paleomagnetic data from these and so-called exotic or suspect terranes will require significant shifts in the positions of some of these fragments. But because the areas affected by Mesozoic and Cenozoic deformation are relatively small, errors in the reconstruction of these fragments do not significantly affect the global picture. Despite an increase in paleomagnetic and other relevant data from Southeast Asia it has been arbitrarily kept attached to Asia in its present-day position and configuration. Paleomagnetic data that became available after the compilation of the maps suggest significant modifications to the positions of China and neighbouring regions to those shown on the Jurassic and Triassic maps (Enkin et al. 1992).

The maps show some continental fragments such as northern and southern Tibet sweeping across the Tethys. This representation is largely schematic; these fragments probably did traverse the Tethys, but their shapes and the timing of their migration are illustrative rather than definitive. No attempt has been made to document the data on which the positions of these fragments and other fragments caught up in the Cenozoic–Mesozoic orogenic belts are based. Differences in the positions of these fragments will not alter the global area covered by sea but only the distribution of that area.

2.4 Orienting a reassembly

The reassembled continents must then be oriented so that the global mean paleomagnetic pole of the time concerned coincides with the map projection pole. The availability of paleomagnetic catalogues (Piper 1988) and machine-readable paleomagnetic databases (Lock and McElhinny 1991) has led to an increase in the number of poles that may used for this purpose.

Reliable paleomagnetic pole positions for the tectonically stable areas are rotated back to a reference frame - in this case the reference fragment used is Africa. For Mesozoic and Cenozoic time there is no ambiguity about whether the pole is the North or South pole. A window of 30 Ma has been applied to the paleomagnetic data and the global mean pole on Africa has been determined at 20 Ma intervals. Thus the mean paleomagnetic pole for a map, say, for 100 Ma has been found by averaging all poles lying in the 85-115 Ma age range. The mean age of the data will in general differ from the mean age of the window and the actual pole position at 100 Ma is interpolated between two bracketing ages, e.g. between 97 Ma and 118 Ma. The global mean pole for each stage is the interpolated position of the pole whose age is the mean age of the stratigraphic stage concerned. The estimate of the Earth's spin axis on the reassembly, i.e. its polar axis, is taken as the global mean pole. A map is then made by projecting the reassembly as a global map using the global mean paleomagnetic pole as the geographic pole for the time concerned. The new databases have decreased the uncertainty of the global paleomagnetic poles for Mesozoic and Cenozoic reassemblies but have not greatly changed the positions of the mean poles. Alpha-95 values - estimates of the uncertainty in pole position and therefore of paleolatitude - rarely exceed 5° and are generally much less.

Despite the availability of all these new data, comparisons of the present maps with previously published maps (e.g. Smith et al. 1981) show that the new maps are refinements of previous reconstructions rather than representing radically new continental distributions. It is clear that any new data are most unlikely to alter the positions of any of the major continents in any significant way, particularly on the scale adopted.

2.5 Choice of map projection

Projecting a sphere on to a plane always poses a problem. Since the intention is to show a global view then Mercator's projection cannot be used. Originally, it was intended to use two projections, one highlighting the equatorial regions and another for the two polar regions. However, this would have tripled the number of maps. The atlas uses Mollweide's projection, an equal-area projection commonly used to show the whole globe. Because the maps are computer generated, it is in practice a trivial matter to generate maps on a variety of other projections. All maps depict the present-day coastline as well as the paleocoastline. The present-day coastline is shown so that the continental fragments can be readily recognized. The geographic grid that has been superposed on the maps is centred on the mean paleomagnetic poles of the time concerned. Latitude and longitude lines are spaced at 30° intervals.

_ 3 _

Biostratigraphic time and magnetic polarity scales

3.1 General statement

The biostratigraphic subdivisions and numerical ages in Ma assigned to them are those of Harland et al. (1990). Compared with other time-scales, this scale gives one of the more extreme (older) values for the Jurassic/Cretaceous boundary (at 146 Ma). Modifications to parts of the Cenozoic scale have been made (Berggren et al., 1991) and others suggested to the Mesozoic scale (Gradstein et al., 1993). The magnetic polarity scale and the ages assigned to polarity changes are also taken from Harland et al. (1990).

The Cenozoic magnetic polarity scale is also under revision (Cande and Kent 1992). The revisions are of two types: those that change the relative spacing of the polarity changes and those that modify their numerical ages. Changes in relative spacing, if confirmed, will affect all subsequent magnetic polarity scales; modifications to the numerical ages are likely to be matters of discussion and significant refinement for some time to come.

3.2 Stratigraphic subdivisions used

Thirty one maps have been drawn for Mesozoic and Cenozoic time. Their stratigraphic age ranges, alphabetic abbreviations and numerical ages (Table 2) are taken from Harland et al. (1990). Where two or more Stages are covered by a single map, e.g. the Messinian—Tortonian, making up the Late Miocene, the younger Stage is listed first. For the older Jurassic and part of the Triassic, the numerical ages have been rounded to the nearest 5 Ma. The Bathonian and Aalenian maps have been excluded for lack of data. However, the Callovian and Bajocian maps, which bracket the Bathonian, are similar to each other as are the Bajocian and Toarcian maps, which bracket the Aalenian. The unpublished University of East Anglia Bathonian and Aalenian shoreline maps show them to be intermediate between the bracketing Stages.

Atlas of Mesozoic and Cenozoic coastlines

Table 2. Stratigraphic and numerical time-scale

Stratigraphic age	Abbreviation	Numerical age (Ma)
Tertiary		
Neogene		
Pliocene	Pli	5
Late Miocene (Messinian-Tortonian)	Mio3	10
Middle Miocene (Serravallian-Langhian)	Mio2	12
Early Miocene (Burdigalian-Aquitanian)	Mio1	20
Paleogene		
Oligocene	Oli	30
Late Eocene (Priabonian)	Eoc3	37
Middle Eocene (Bartonian-Lutetian)	Eoc2	45
Early Eocene (Ypresian)	Eoci	53
Paleocene (Thanetian-Danian)	Pal	60
Cretaceous		
Late Cretaceous		
Maastrichtian	M22	70
Campanian		80
Santonian	Cmp San	85
Conjacian		
Conjacian Turonian	Con Tur	88 90
		90 95
Cenomanian	Cen	75
Early Cretaceous	A 11	105
Albian	Alb	105
Aptian	Apr	115
Barremian-Hauterivian	Brm-Hau	125
Valanginian-Berriasian	Ber-Vlg	138
Jurassic		
Late Jurassic (= Malm)		
Tithonian	Tửn	148
Kimmeridgian	Kim	153
Oxfordian	Oxf	155
Middle Jurassic (= Dogger)		
Callovian	Clv	160
Bathonian	Bth	Map not made
Bajocian	Baj	170
Aalenian	Aal	Map not made
Early Jurassic (= Lias)		
Toarcian	Toa	180
Pliensbachian	Plb	190
Sinemurian	Sin	200
Hettangian	Het	205
Triassic		
Late Triassic (Rhaetian-Norian-Carnian)	Tr3	220
Middle Triassic (Ladinian-Anisian)	Tr2	238
Early Triassic (= Scythian)	Tr1	245

Abbreviations from Harland et al. (1990). Stratigraphic ranges from Harland et al. (1990) but ages do not necessar correspond to the mid-points of each range.

_ 4 _

Plotting paleogeographic data

4.1 Initial map preparation

The paleogeographic data were compiled initially on base maps on a scale of about 1:10 m. At such a scale it was possible to distinguish a number of marine and non-marine facies. However, information at this level of detail was routinely available (in the public domain) only for western Europe and North America. These unpublished detailed facies maps are the basis for the paleocoastline maps in this Atlas.

The total Bibliography of sources for the information summarized in the shoreline maps given on pp. 55–99 provides an essential pathway into more detailed local or regional information. It is important to note, however, that this information covers only the period up to 1985. More recent literature is not included or referred to.

In addition to the literature compiled by the University of East Anglia, BP had access to the global reconstructions of the Paleogeographic Atlas Project of the Department of Geophysics, University of Chicago (A. M. Ziegler, personal communication), a version of which has been published (Ziegler, Scotese and Barrett 1982). This series comprised some 16 reconstructions, each representing a single Stage, for which data had been compiled from the literature at a density of up to one datapoint for each one degree latitude/longitude square.

Areas of higher relief for the time concerned are shown schematically on each map. These are generalized from BP's database, which is a modified version of the data collected by F. Ziegler's group at the University of Chicago.

The global shoreline maps that form the basis of this atlas were compiled by staff and vacation students over a two-year period at the BP Research Centre. The compilation was initially on a present-day geographic base, with a map for each Stage from the base of the Jurassic to the Pliocene. The University of East Anglia compilation did not include the Triassic, and maps for each of the Early, Mid and Late Triassic were added to the series at this stage.

4.2 Transfer to global reconstructions

The shorelines were then transferred to global continental reconstructions for each Stage, using reconstructions generated by BP's GLOBE system. Because of the termination of the Global Paleoreconstruction project in BP, it became necessary to transfer the paleocoastlines to ATLAS for final plotting for this publication. However, by this time the detailed rotations that specify the paleocontinental positions were significantly different between ATLAS and GLOBE. Also, because the reconstructed positions of the shorelines were digitized, but their present-day positions were not, it was necessary to go through a rather cumbersome procedure at this point. The shorelines had to be divided into segments corresponding to the continental fragments used in GLOBE, then back-rotated using GLOBE's rotation data to recover the original position of the shorelines on present-day maps. They were then rotated again using the ATLAS rotations, and tidied up before generating the plots presented in this volume. Difficulties were experienced with back-rotating the data for only two maps – the Valanginian and Berriasian map, and the Barremian and Hauterivian map. The shorelines on both these maps have been redigitized from their present-day positions rather than back-rotating the old reconstructed data.

4.3 Estimation of paleocoastline positions

The shoreline is the paleogeographic element that is perhaps most commonly singled out on maps of this kind. The scientific justification for selecting this particular feature is not obvious, because the coastline is a notably ephemeral feature – it may migrate backwards and forwards across large areas of a continent with considerable rapidity. However, it provides geologists with a very sensitive indicator of changes in base level, particularly in areas of low gradient where a small relative rise or fall in sea level will cause a correspondingly large movement of the coastline. The level of interest in the degree of control exercised by changing sea-levels on patterns of sedimentation and stratigraphic packaging is currently high, and maps of the present time are a potentially important contribution to this discussion (see below).

Ideally, placing a paleocoastline on a map is no more than drawing the boundary between the extent of marine sedimentation on one side of the line and continental on the other. Only in the most intensively researched parts of the globe, principally in western Europe and North America, is adequate information available. At its best this may involve well-dated sequences of marginal marine (paralic) facies where the saltwater/freshwater or marine/terrestrial interfaces can be defined in actual sedimentary deposits. More commonly only terrestrial-freshwater or marine-saltwater deposits are present, and the shoreline is defined from only one side. Particularly for periods of major marine transgression the marginal marine deposits may be subsequently eroded away, and their preserved extent is then only an indication of the minimum extent of the transgression. In these circumstances variations in the apparent depth of deposition of the marine deposits have been interpreted as indicators of the relative extent of transgressions and commensurate shoreline displacement. Further consideration of the

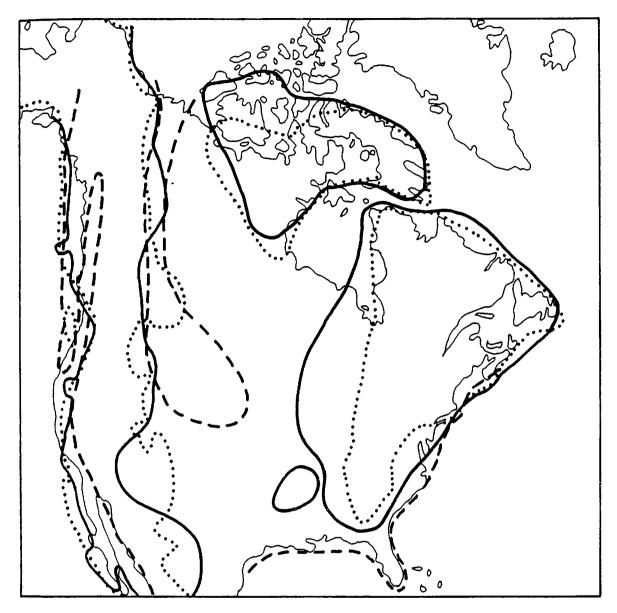


Figure 1. Three Maastrichtian coastlines for North America. The full line is from Map 10, this atlas; the dotted line is from E. C. Kauffman and D. A. Beeson (unpublished results) and the dashed line from Funnell (1990). Lambert equal area map with a central meridian at 90° W. See text for discussion.

limitations of global and regional maps in respect of paleocoastline and facies distributions have been given by Tyson and Funnell (1987) and Funnell (1990).

In practice the low density of the available data dictates that it is very rarely so simple, and paleocoastlines at the global scale must be regarded as very speculative. For many geological Stages, there are simply no data at all from many parts of the globe. Under these circumstances, it may be possible to make reasonable estimates of the likely shoreline position

from its position in preceding or succeeding Stages. In the less explored or researched parts of the globe such interpolation may extend over many Stages and real control over the shoreline position is very tenuous. An example is Antarctica, for which most of the maps simply assume that the continent was elevated for most of Mesozoic and Cenozoic time (except where marine deposits are known to be present), and we therefore show a paleocoastline that approximates to that of the present day.

As examples of the variation in interpretation that is possible between different research groups, the Cretaceous shoreline maps of Tyson and Funnell (1987) and Funnell (1990) can be compared directly with the maps in this atlas which were prepared by different researchers but from very much the same data (Figure 1). They may also be compared with a third set of maps for North America (E. C. Kauffman and D. A. Beeson, unpublished data). The philosophy adopted in compiling our maps was to illustrate, in cases of diverging interpretations for the same Stage, the maximum amount of marine transgression. The Maastrichtian maps are of particular interest because they show the greatest divergence. Funnell's (1990) map shows no through-going north-south seaway east of the Rocky Mountains. By contrast, the map in this atlas and that of Kauffman and Beeson show a wide seaway forking into two, with one branch extending to Hudson's Bay and beyond. Though broadly similar, the paleocoastlines of these two maps differ in detail. By the Paleocene, the western seaway had disappeared. The differences in all three maps reflect mostly the duration and age of the time-slice used. Kauffman and Beeson's map is for a very short time interval the Baculites clinolobatus chron in division R9 of the Bearpaw cycle - whereas the other two are representative of different parts of the Maastrichtian. Thus Funnell's (1990) map is probably representative of a later part of the Maastrichtian than the map in this Atlas. The latter probably represents the average Maastrichtian shoreline during the earlier Maastrichtian. All three maps are therefore acceptable on their own terms, but together they illustrate the difficulties of drawing and evaluating individual paleocoastline maps, particularly when large variations take place within a Stage.

4.4 An illustration of the use of the Bibliography

To give some idea of the number and distribution of the references used to compile a single global map, we list here the references used for the Albian Stage map.

In preparing the original shoreline and facies maps, data were generally consolidated on to regional base maps. These regions comprised: Arctic (including the USSR), North America (including the Caribbean), Europe, Far East (including Australasia), Africa, Middle East (including India and the Indian Ocean), and South America. References are cited as the initials of the first or sole author plus a two-digit date (20th century assumed), the first initial being that of the family or surname. For example, Bay, T.A. 1977 is abbreviated as bta77. In the case of more than one reference in the Bibliography being represented by the same initials and date, it will usually be evident from the context which reference is intended. A slash indicates more than one author; a hyphen represents a hyphenated name.

Arctic Albian references None were used.

African Albian references A total of 35 references was used: as/83, ap/83, bft72, bk74, bgp/66, bpf/83, bh66, bg70, cg/62, fr/79, fr75, fr81, fs/73, gpcde/82, gr75, k-km82, kca76, lz/76, lr76, me73, mpa61, mrc72, psw83, pc75, ru82, rra/76, rd66, sj78, sj80, sj/81, tb/79, wa82, wj82, ww83, wp/82

European Albian references A total of 80 references was used: ac78, a-va/79, ah64, aj/74, aj/75, ba/78, bgm/69, bpy/79, bpy/78, ba/79, bj76, bph82, cj79b, cp/80, cs76, cm/82, clr74, ci/81, cp/71, cp71, cl/72, cg/79, dra/71, dm/83, dr78, di67a, di67b, di68, dpvo70, dp/78, fp67, fp71, fp/73, frm/75, fw67, g-hm79, gaa/78, gmj63, ga80, ge82, gmp71, hm78, hbb69, jm66, j-sm79, j-sm81, km83, km63, lgk/80, ms/79, mc/80, me74, me79, md/81, mr68, mj/83, ndp81, ogh73, ogh75, ohg79, pb78, pc75, qr69, ra79, rp68, rpf80, rj79, rj/77, rf81, rc73, rn75, sw/69, sb/76, sp/77, tmja71, tka81b, wr67, we73, ww64, zpa82a

Far East Albian references A total of 22 references was used: aga78, amc79, bdd80, bda/68, bcm/80, dji/76, ejl72, epr81, enf/76, enf/80, hg69, kt74, kt80, lnh78, mt78, os82, ppe75/, sj/80, sgr/78, vjj/82, wgd66, wh/71

Middle East Albian references A total of 17 references was used: ba76, ba77, br71, bt/80, imw81, imw83, jga/65, kwj/83, lz/76, mrj80, phs79, plf/69, rr/68, sa78, wd68, wjl75, wr/80

North American Albian references A total of 70 references was used: bta77, bew63, bwl/83, bba70, bwf80, bwr/69, bdj68, cma71, ctp73, cac63, ctd/75, csp83, djc70, d-cz70, drl73, dj79, dj82b, drjw/70, ep74, ep77, ep83, fwl/69, fjm63, gmm77, gar/82, hpt70, jlf/75b, jja71a, jja75, jja76, jja77a, krs77, kkm/71, lm73, l-fj81, lmw/78, l-re81, mcgdp/72, mcgaa/68, mjr/81, maa/74, mad73, mad79, mch/66, ohg73, pbf74, rfg81, rjb57, rdd/80, rra64, sjs81, srw70, srw79, sdf73, sci70, sci81, snf/76, scr75b, sdf68, tm77, wrj83, wr80, wgd/75, whh74, wbj/83, wgr62, ycj/81, yrg60, yfg/76, zpa69.

South American Albian references A total of 37 references was used: ae/73, bfa74, bp/81, bc83, bve56, bve68, bk70, bh73, cg70, cj79, db78, drh/82, e-sf/76, fa/63, hhj62, i.e.m75, jm73, jh74, jh77, lg78, mdim/83, mn/83, mhg71, mf79, mjc/64, mrc/75, nep/80, oha74, oha82, pe/69, rl67, rra/76, spc/76, sf79, wj82, wjj63, ze/72.

The dominance of references from Europe and North America is clear.

Readers can search additional information on individual Stages and groups of Stages by examining the Bibliography.

_ 5 _

Changes in land area through time

Regression, or the retreat of the sea, ideally produces a characteristic lateral sequence in the stratigraphic record, also seen in a vertical section. Similarly, advances of the sea, or transgressions, produce sequences of a different but equally distinct character.

The most important influence on the position of shorelines worldwide is global sea level. It has been generally recognised in recent years that geologically rapid changes in global sea level are not only a characteristic of Ice Ages, such as the Quaternary, but also occur at a rather less rapid frequency, throughout the geological record (Haq, Hardenbol and Vail 1988). Certainly significant changes in global sea level occurred within many of the Stages represented in this atlas. On the whole the paleocoastlines we have indicated correspond to episodes of maximum marine transgression (maximum global sea level) within the time period of the individual map.

Although we leave interpretation and further application of our maps to the interested reader, we have taken the opportunity of using them to estimate changes in land area (and hence in marine inundation) through the Mesozoic and Cenozoic, as this provides a unique source of information on the sea level. The area of land on all maps has therefore been calculated, directly from the digitized data. The distribution of land area through time is plotted in Figure 2 and listed in Table 3. It is difficult to estimate the uncertainties in these estimates, but the results of the calculations are shown to the nearest million square kilometers. This is about the size of the non-terrestrial areas between Australia and Antarctica on the Santonian to Cenomanian maps (Maps 12–15). It is unclear whether these areas had connections to the world ocean, as assumed in Figure 2 and Table 3, or were inland seas.

Except for the Maastrichtian and Campanian, the gross fluctuations in land area are remarkably similar to the changes in sea level inferred by Haq et al. (1988), at least on a time-scale of about 25 Ma or more. Such a relationship is to be expected. On the other hand, our sea level curve and that of Haq et al. have been derived from quite different datasets. It is not our intention here to discuss the significance of our data in this and other respects, but a recent interpretation of global sea level changes is given by Gurnis (1993).

Table 3. Non-marine areas and ages

Map and stratigraphic age	Age (Ma)	Non-marine area in 10 ⁶ km ²
Map 1. Pliocene	5	153
Map 2. Late Miocene (Messinian-Tortonian)	10	152
Map 3. Middle Miocene (Serravallian-Langhian)	12	150
Map 4. Early Miocene (Burdigalian-Aquitanian)	20	150
Map 5. Oligocene	30	147
Map 6. Late Eocene (Priabonian)	37	140
Map 7. Middle Eocene (Bartonian-Lutetian)	45	138
Map 8. Early Eocene (Ypresian)	53	137
Map 9. Paleocene (Thanetian-Danian)	60	138
Map 10. Maastrichtian	70	109
Map 11. Campanian	80	110
Map 12. Santonian	85	118
Map 13. Coniacian	88	118
Map 14. Turonian	90	121
Map 15. Cenomanian	95	118
Map 16. Albian	105	129
Map 17. Aptian	120	131
Map 18. Barremian-Hauterivian	130	125
Map 19. Valanginian-Berriasian	138	127
Map 20. Tithonian	148	130
Map 21. Kimmeridgian	153	130
Map 22. Oxfordian	155	132
Map 23. Callovian	160	128
Map 24. Bajocian (There is no Bathonian map)	170	138
Map 25. Toarcian (There is no Aalenian map)	180	139
Map 26. Pliensbachian	190	142
Map 27. Sinemurian	200	146
Map 28. Hettangian	205	150
Map 29. Late Triassic (Rhaetian-Norian-Carnian)	220	134
Map 30. Middle Triassic (Ladinian-Anisian)	238	139
Map 31. Early Triassic (= Scythian or		
Spathian-Nammalian-Griesbachian)	245	135

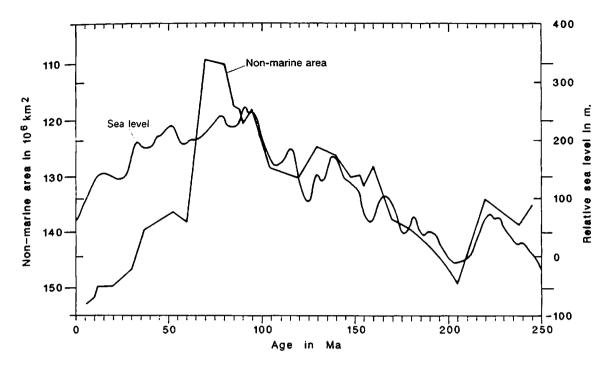


Figure 2. Area of land in millions of square kilometers plotted against time in Ma, superposed on sea level changes in meters versus time in Ma from Haq et al. (1988). The graph assumes that the elongate area between Australia and Antarctica for Santonian–Cenomanian time has a connection to the sea. If it was an enclosed sea, the 'non-marine area' for Santonian–Cenomanian time needs to be increased by about 1 million km² for all four stages.

6

Relevant literature

The relevant literature listed below comprises references cited in the text and tables together with subsidiary references that have not been specifically cited. The latter include details of Euler rotations and the geographical boundaries of minor continental fragments used in the reconstructions.

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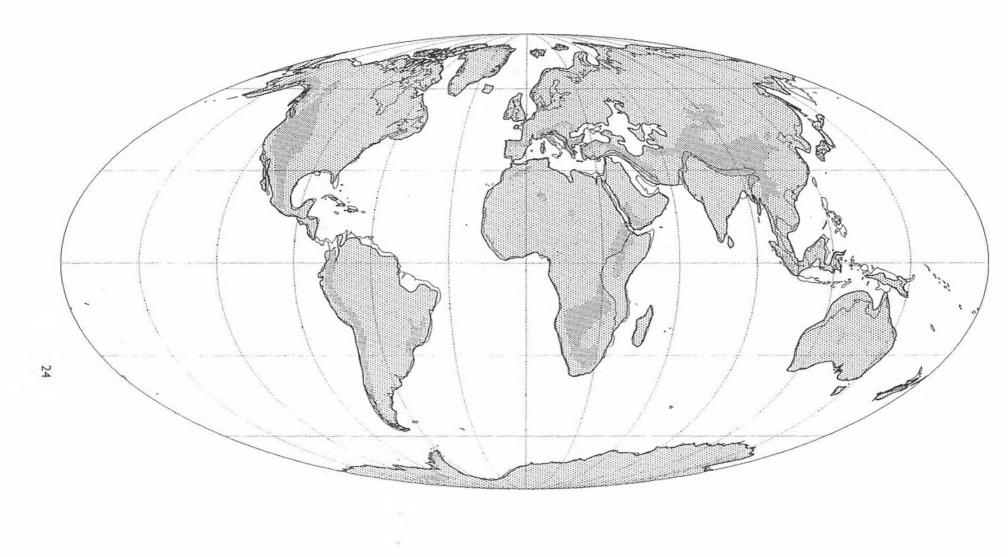
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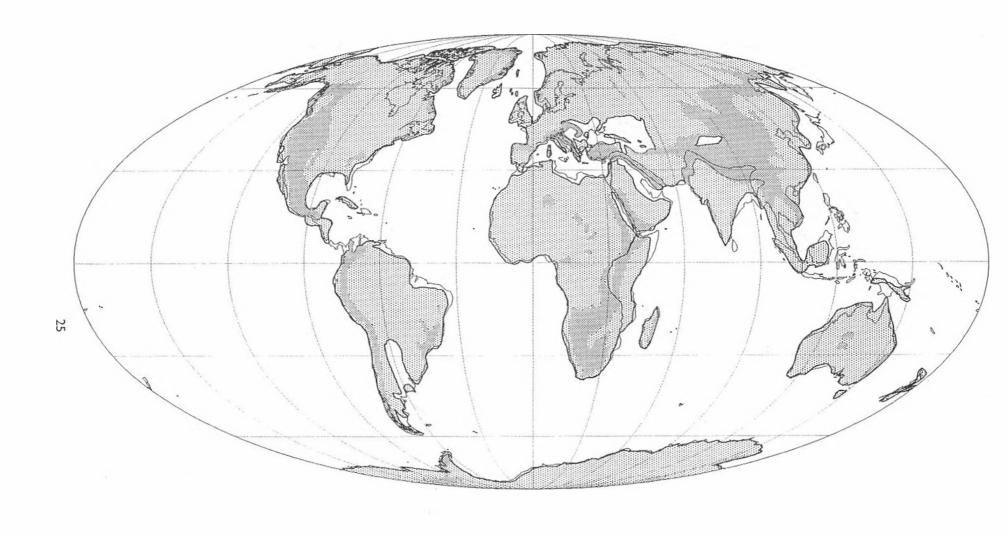
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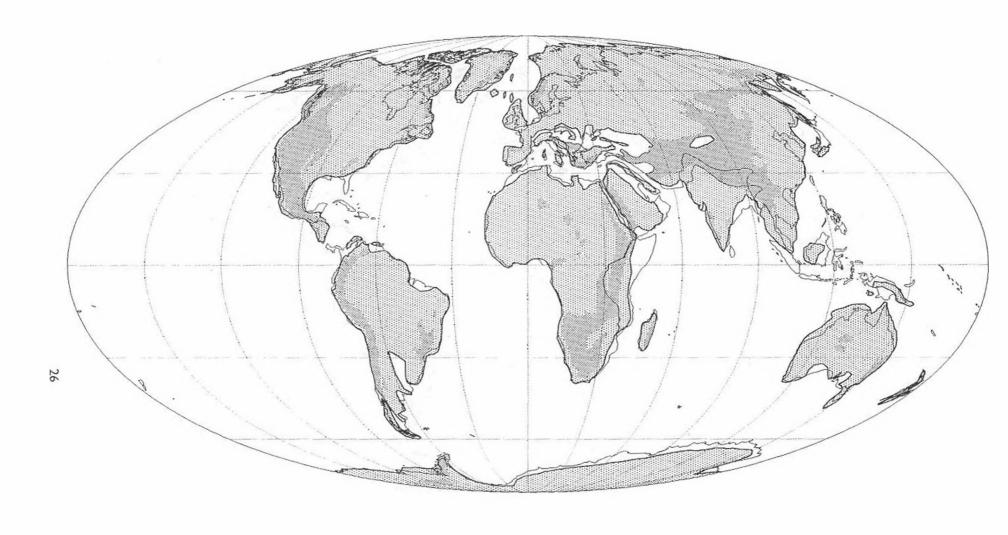
_ ⁷ _ Paleocoastline maps



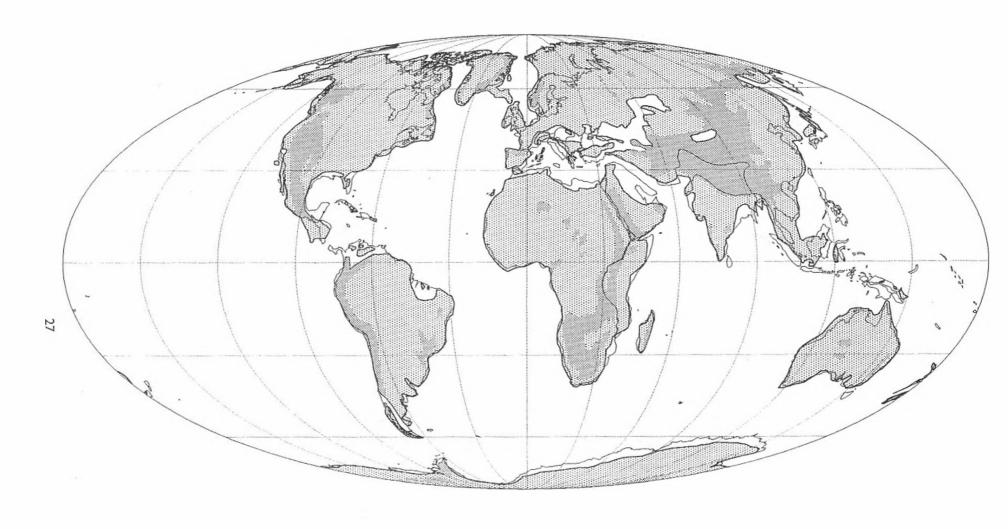
Map 1. Pliocene 5 Ma paleocoastline map.



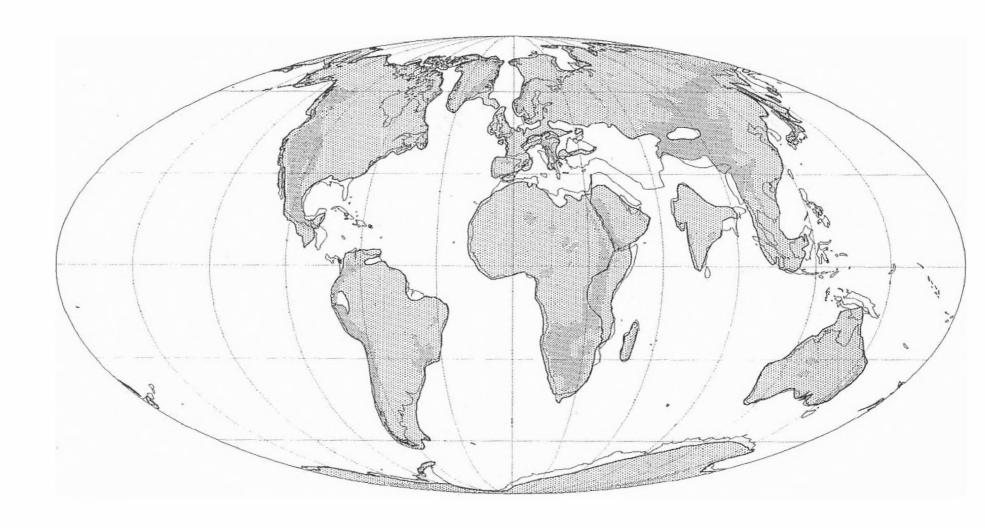
Map 2. Late Miocene (Messinian-Tortonian) 10 Ma paleocoastline map.



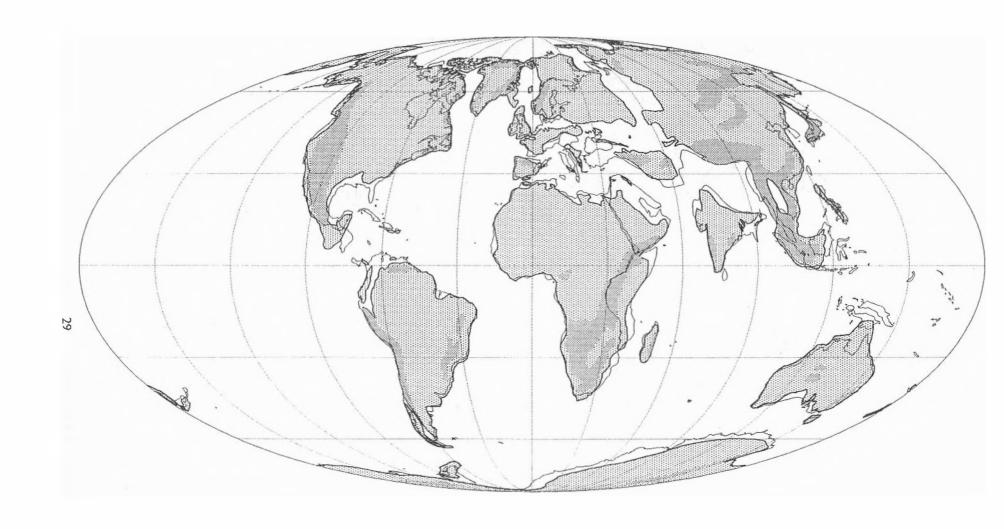
Map 3. Middle Miocene (Serravallian-Langhian) 12 Ma paleocoastline map.



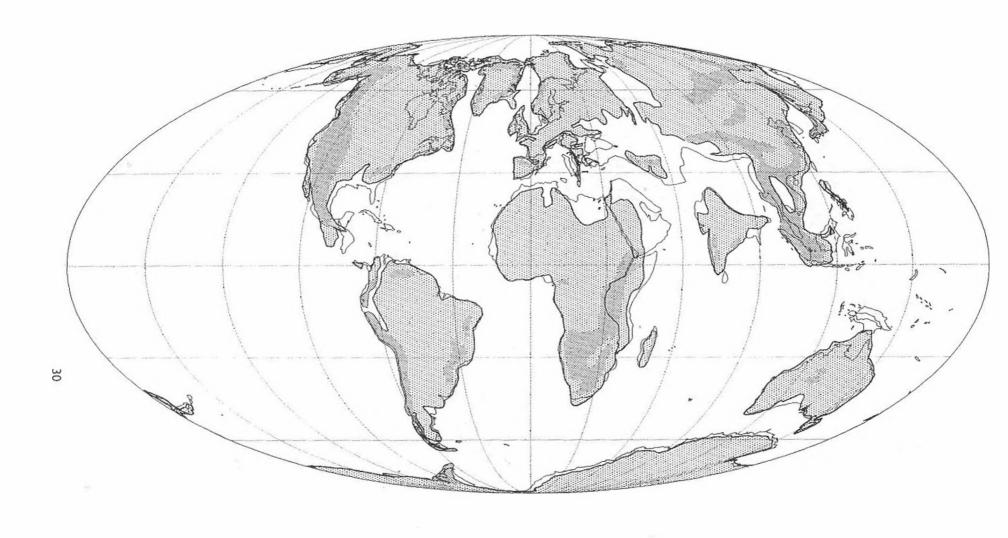
Map 4. Early Miocene (Burdigalian-Aquitanian) 20 Ma paleocoastline map.



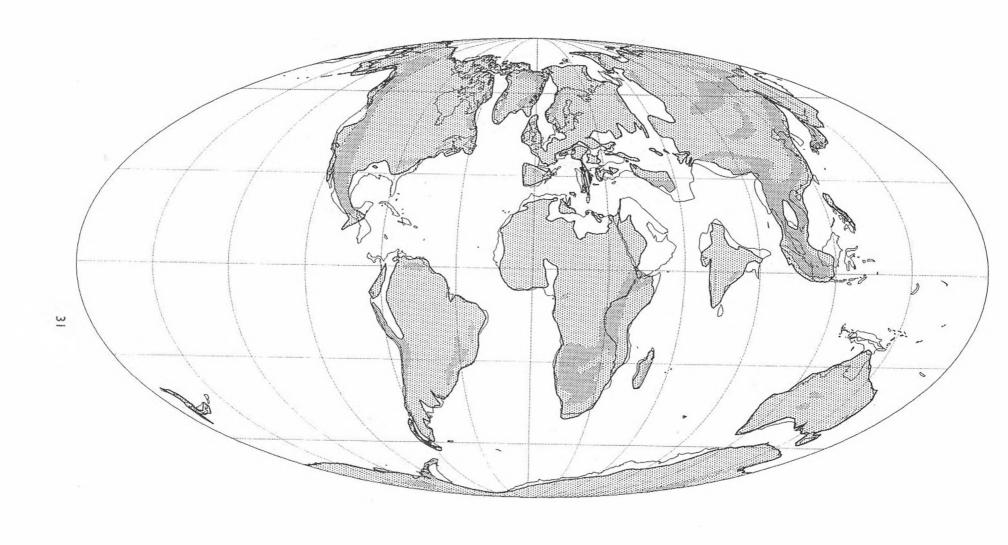
Iap 5. Oligocene 30 Ma paleocoastline map.



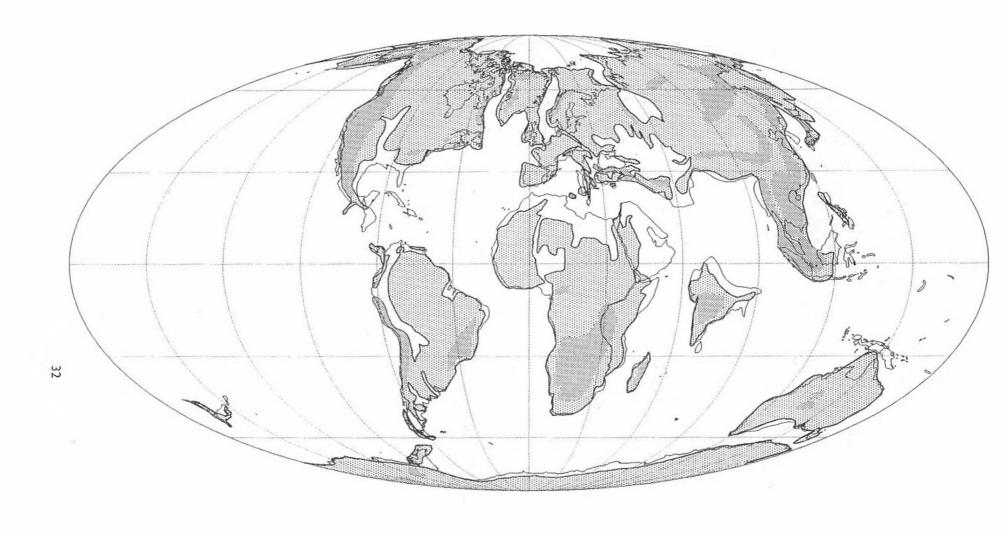
Map 6. Late Eocene (Priabonian) 37 Ma paleocoastline



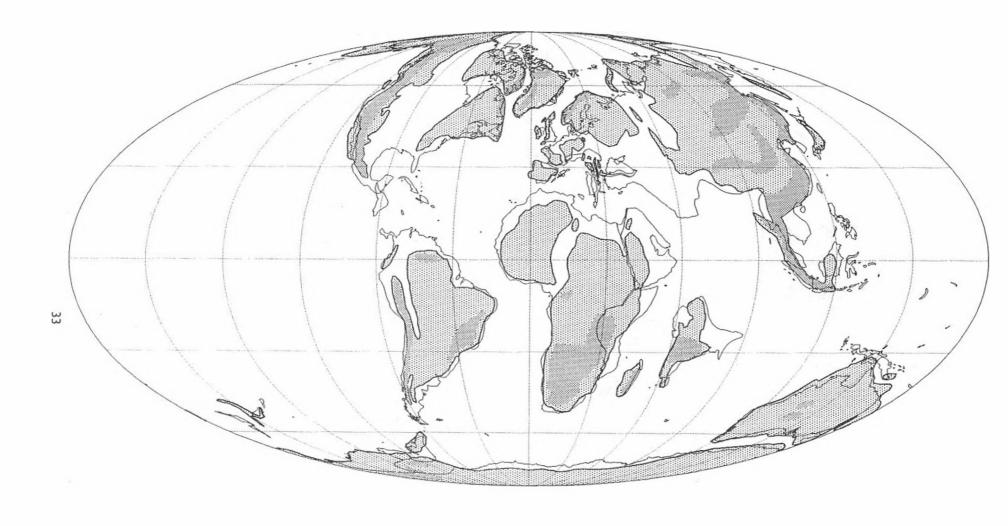
Map 7. Middle Eocene (Bartonian-Lutetian) 45 Ma paleocoastline map.



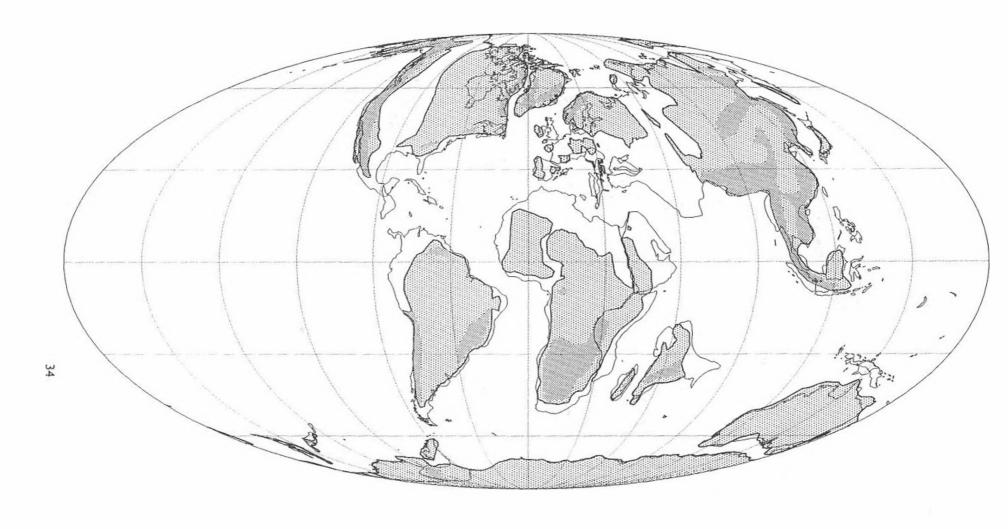
Map 8. Early Eocene (Ypresian) 53 Ma paleocoastline map.



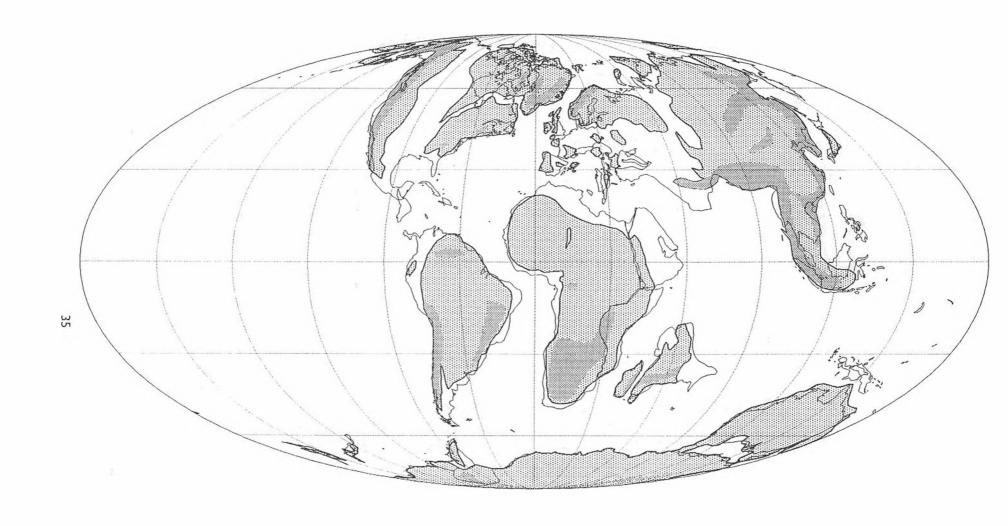
Мар 9. Paleocene (Thanetian-Danian) 60 Ма paleocoastline map.



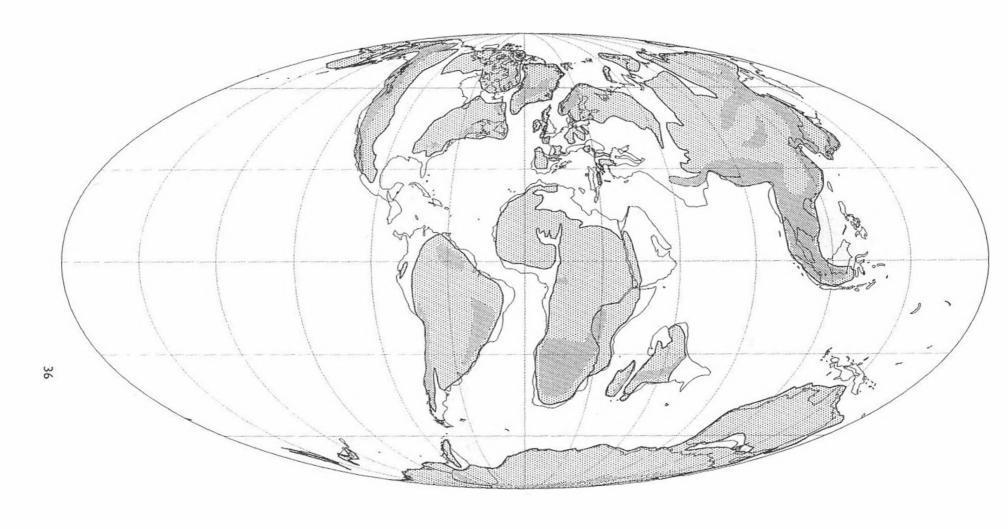
Map 10. Maastrichtian 70 Ma paleocoastline map.



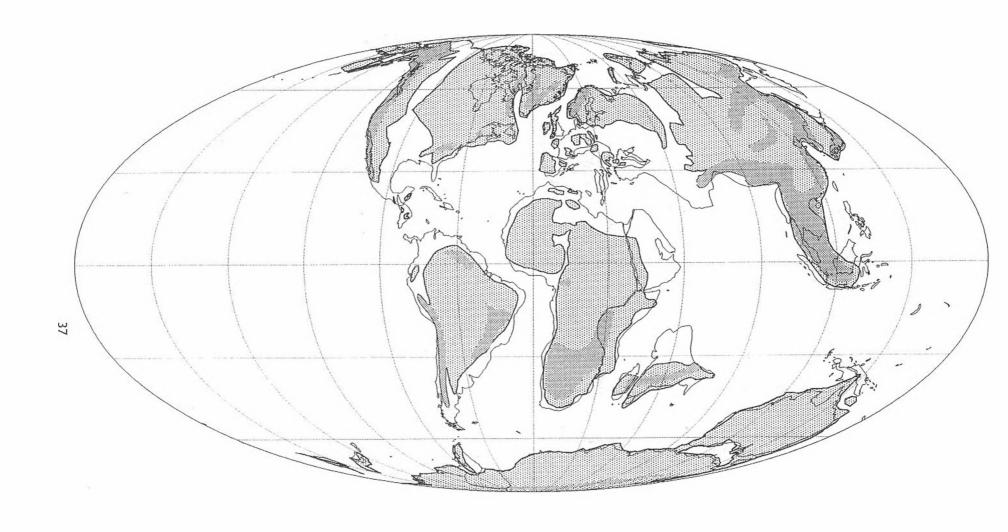
Map 11. Campanian 80 Ma paleocoastline map.



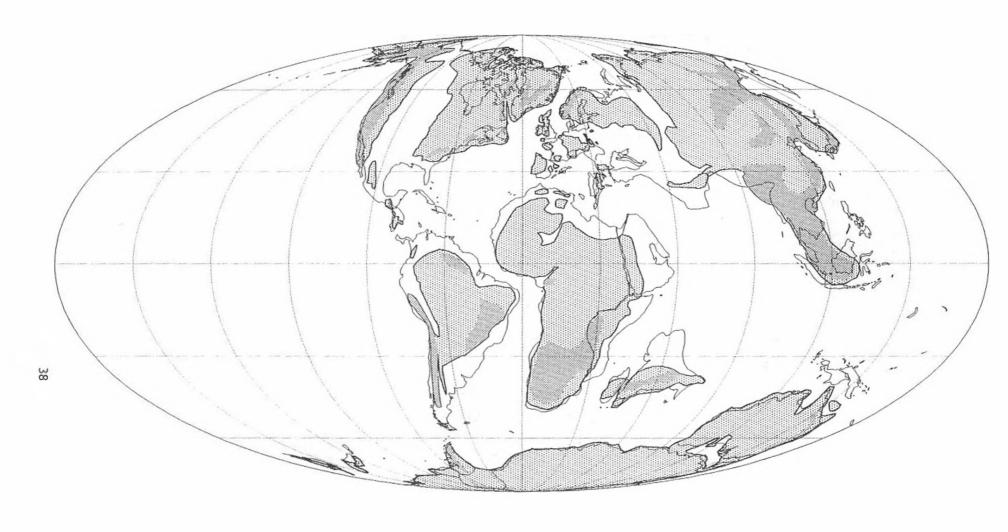
Map 12. Santonian 85 Ma paleocoastline map.



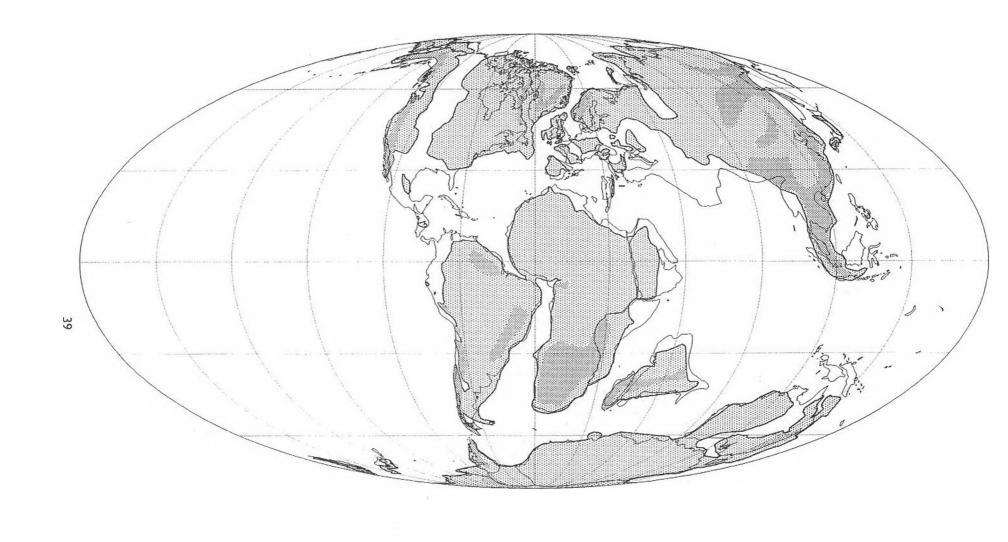
Map 13. Coniacian 88 Ma paleocoastline map.



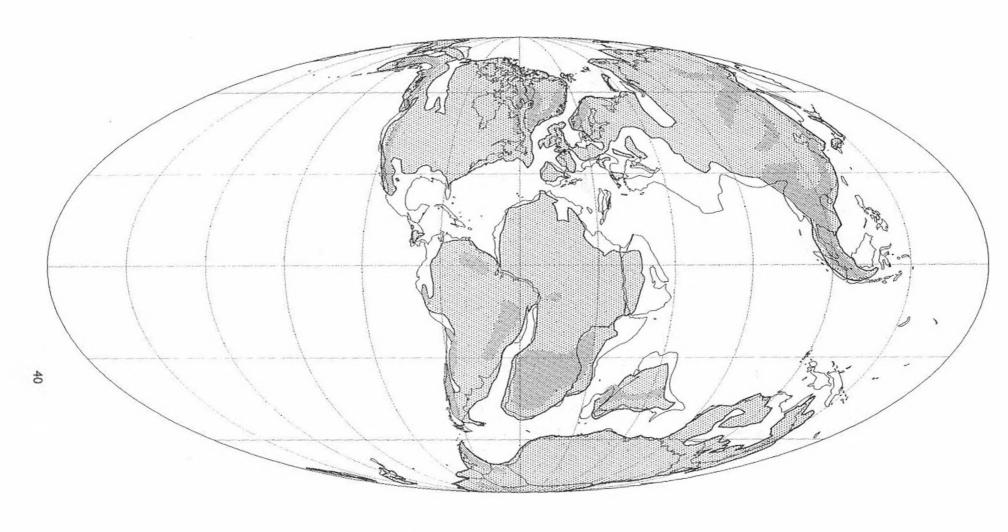
Map 14. Turonian 90 Ma paleocoastline map.



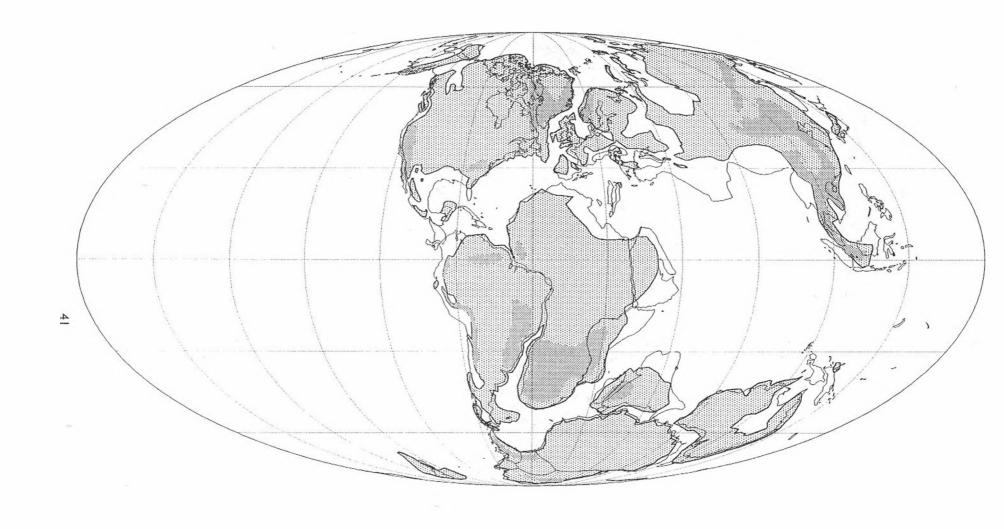
Map 15. Cenomanian 95 Ma paleocoastline map.



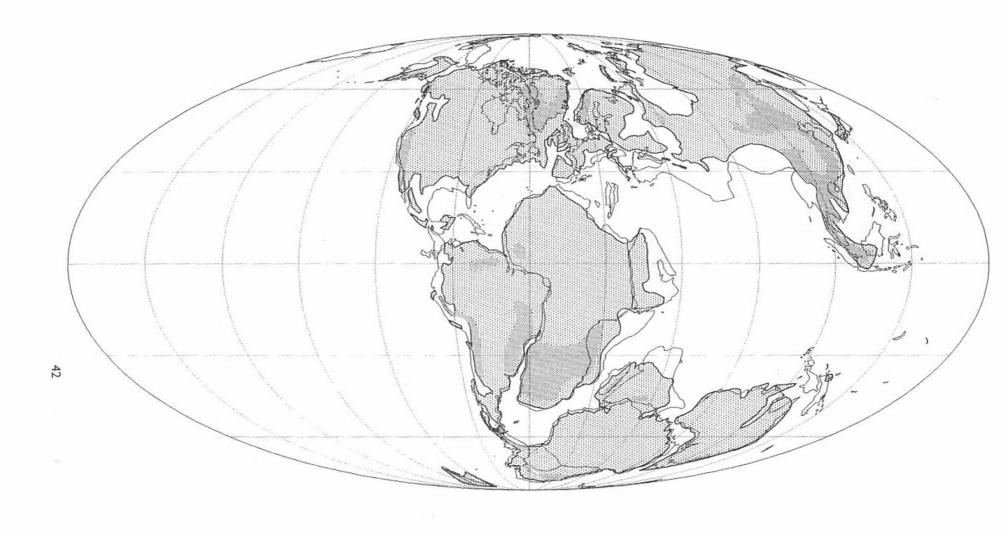
Map 16. Albian 105 Ma paleocoastline map.



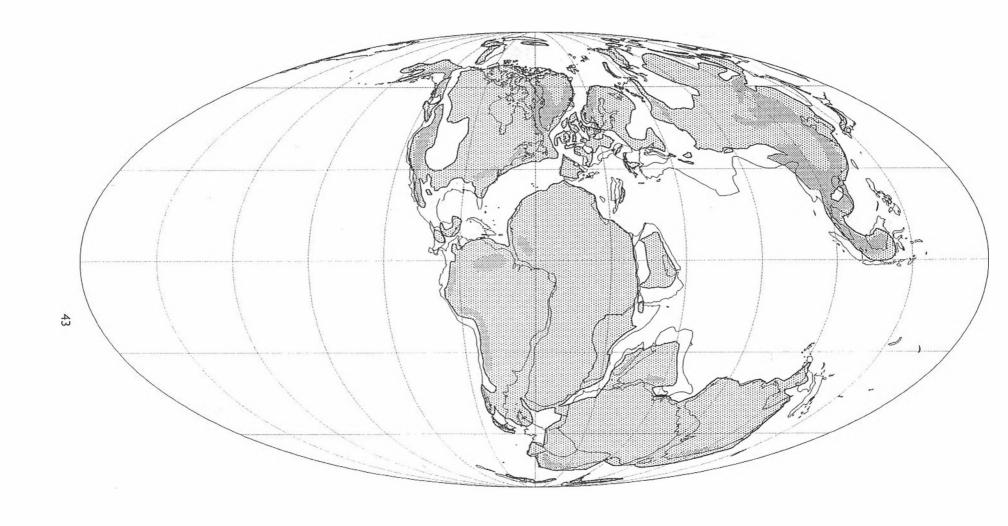
Map 17. Aptian 120 Ma paleocoastline map.



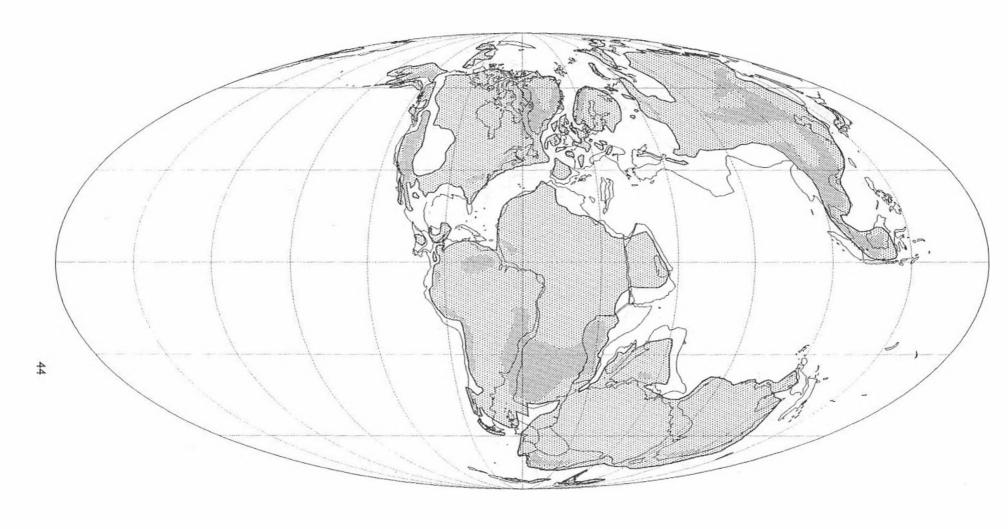
Map 18. Barremian-Hauterivian 130 Ma paleocoastline map.



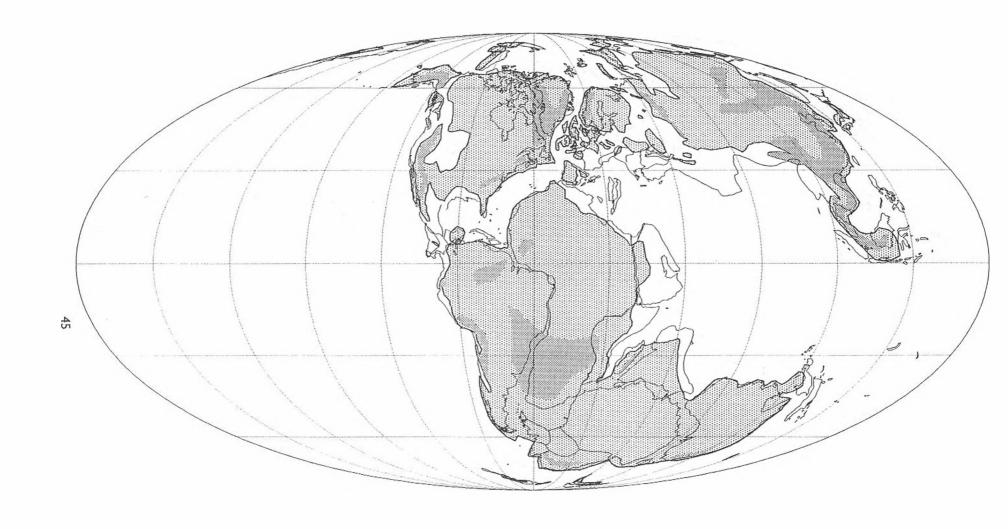
Map 19. Valanginian-Berriasian 138 Ma paleocoastline map.



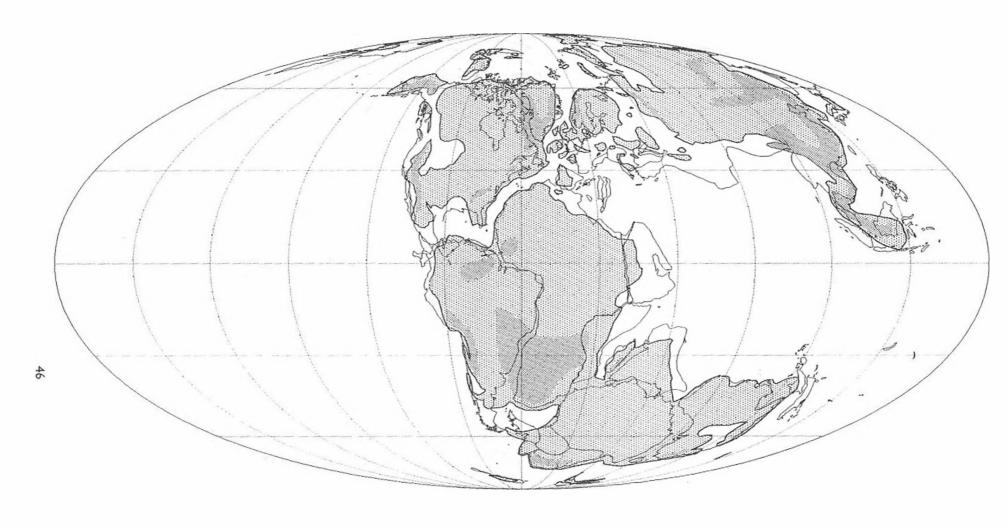
Map 20. Tithonian 148 Ma paleocoastline map.



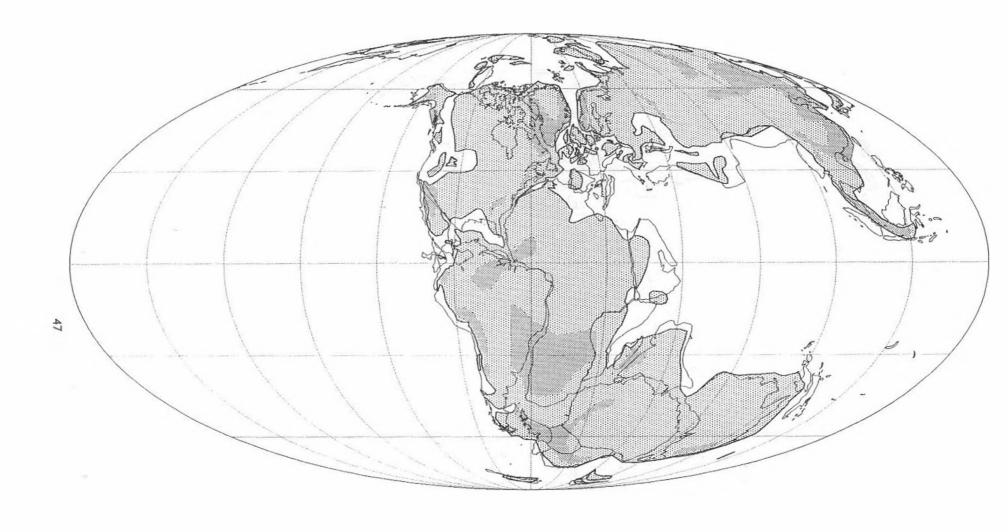
Map 21. Kimmeridgian 153 Ma paleocoastline map.



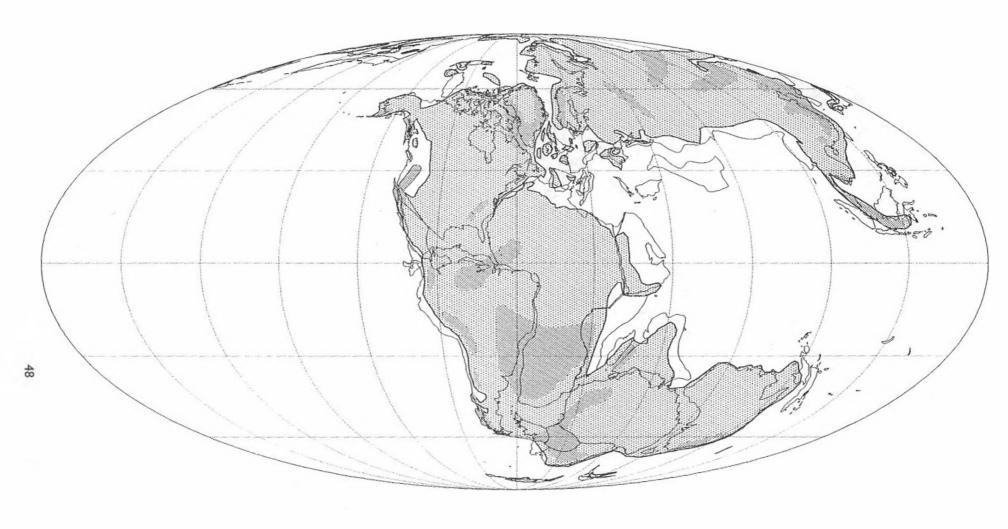
Map 22. Oxfordian 155 Ma paleocoastline map.



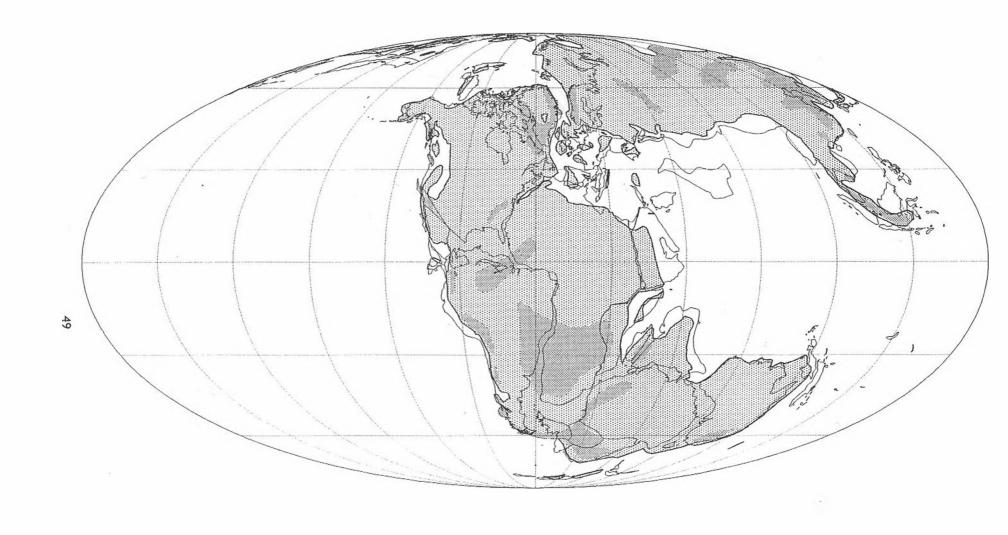
Map 23. Callovian 160 Ma paleocoastline map.



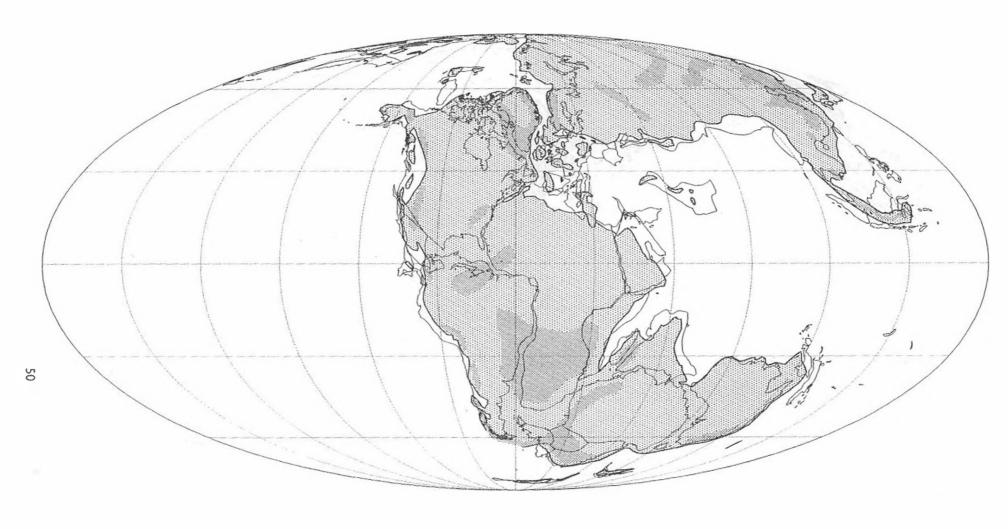
Map 24. Bajocian 170 Ma paleocoastline map. There is no Bathonian map.



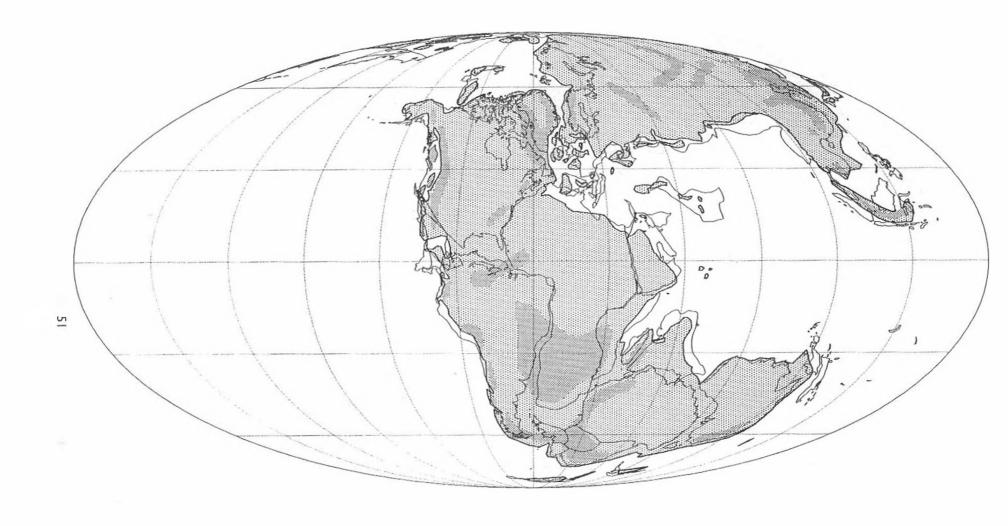
Map 25. Toarcian 180 Ma paleocoastline map. There is no Aalenian map.



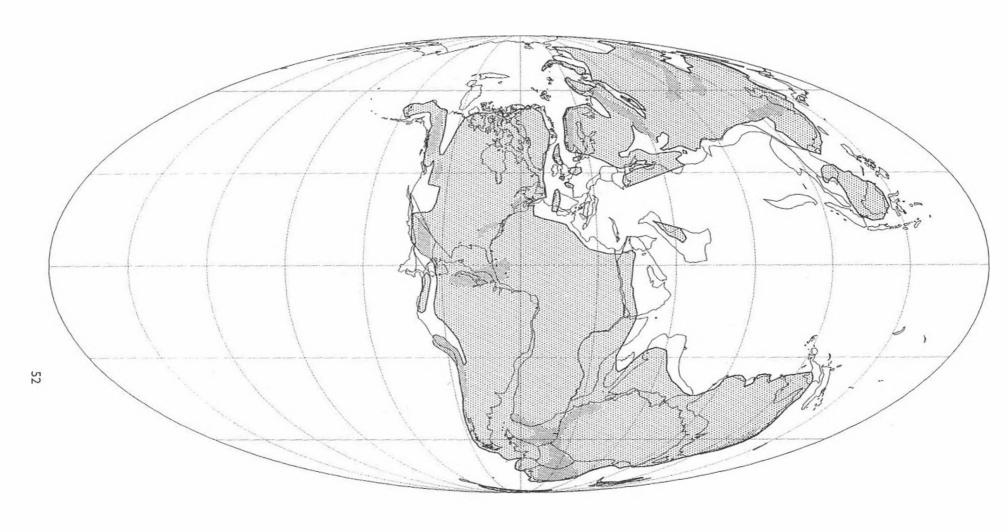
Map 26. Pliensbachian 190 Ma paleocoastline map.



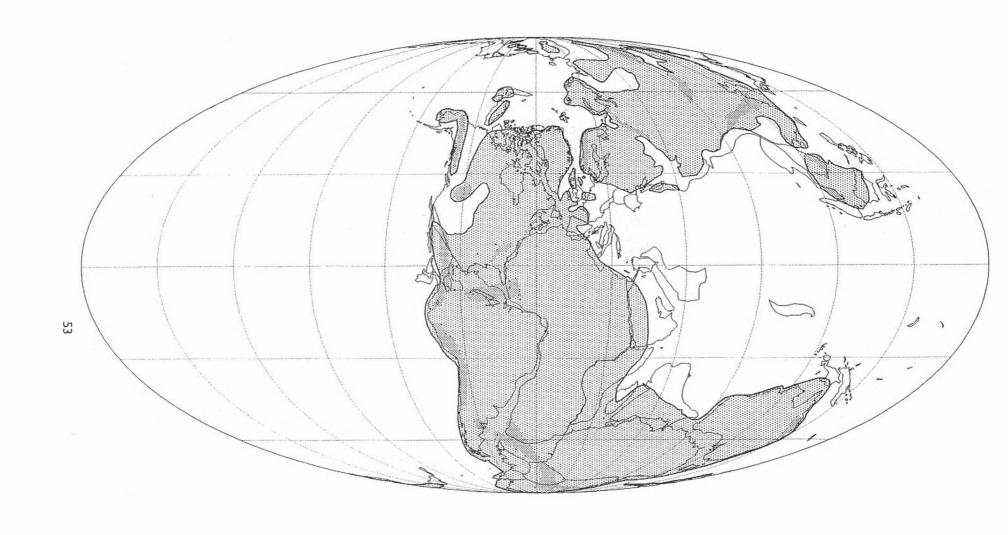
Map 27. Sinemurian 200 Ma paleocoastline map.



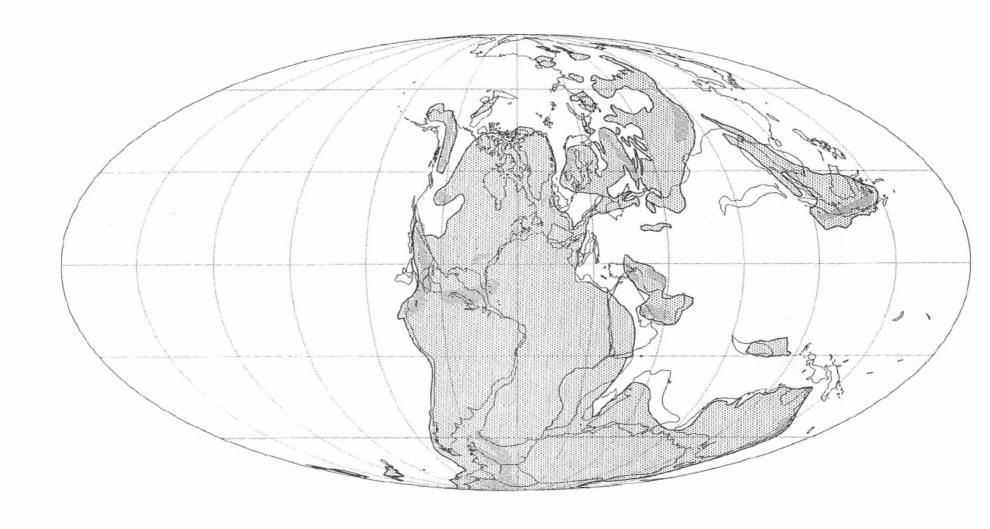
Map 28. Hettangian 205 Ma paleocoastline map.



Map 29. Late Triassic (Rhactian-Norian-Carnian) 220 Ma paleocoastline map.



Map 30. Middle Triassic (Ladinian-Anisian) 238 Ma paleocoastline map.



Triassic Scythian Spathian-Nam riesbachian) paleocoast

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The inferred positions of global paleoshorelines through the 240 million years of the Mesozoic and Cenozoic are presented within this atlas. Thirty-one maps, generally corresponding to stratigraphic stages, provide a snapshot of the continents at approximately 8 million year intervals. The maps provide a representation of the gross changes in the distribution of land and sea throughout the Mesozoic and Cenozoic plotted on Mollweide projections of paleocontinental reconstruction. They do not distinguish between well and poorly defined shorelines, but the information sources are set out in a bibliography numbering more than 2000 primary paleogeographic references. Work up to 1980 was used in the initial compilation, and the final maps use additional unreferenced and unpublished data up to about 1987.



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