PALAEONTOGRAPHICA AMERICANA

(Founded 1917)

VOL. 9

No. 52

LOWER BAJOCIAN (JURASSIC) CEPHALOPOD FAUNAS FROM WESTERN CANADA and PROPOSED ASSEMBLAGE ZONES FOR THE LOWER BAJOCIAN OF NORTH AMERICA

By

RUSSELL L. HALL AND GERD E. G. WESTERMANN

April 3, 1980

Paleontological Research Institution

Ithaca, New York, U.S.A. 14850

Library of Congress Card Number: 80-81384

Printed in the United States of America
Arnold Printing Corporation
Ithaca, N.Y. 14850 USA

CONTENTS

LIST OF ILLUSTRATIONS

LIST OF TABLES

LOWER BAJOCIAN (JURASSIC) CEPHALOPOD FAUNAS FROM WESTERN CANADA

and

PROPOSED ASSEMBLAGE ZONES FOR THE LOWER BAJOCIAN OF NORTH AMERICA

by

RUSSELL L. HALL
Department of Geology
University of Calgary
Calgary, Alberta, Canada

and

GERD E. G. WESTERMANN
Department of Geology
McMaster University
Hamilton, Ontario, Canada

ABSTRACT

The Stephanoceratidae and Sphaeroceratidae of the northeastern Pacific margin, particularly the Queen Charlotte Islands, have been reinvestigated. The Lower Yakoun Formation along Skidegate Inlet contains three distinct late Early Bajocian ammonite faunules, probably in the following succession (from top): *Stephanoceras itinsae* - *Chondroceras oblatum* faunule on South Balch Island; *Stephanoceras skidegatense* - *Chondroceras defontii* faunule at Richardson Bay; and *Zemistephanus richardsoni* faunule at Mackenzie Bay.

Taxonomic and stratigraphic comparisons indicate close affinities to southern Alaska, mainland British Columbia, western Alberta and eastern Oregon, and lesser ones to the western interior of the United States. The following dimorphic species are recognised: *Stephanoceras itinsae* (McLearn) ♀ & ♂ [= *S. yakounense* McLearn ♀ - *Itinsaites itinsae* McLearn ♂], *S. skidegatense* (Whiteaves) ♀ & ♂ [microconch new], *Zemistephanus richardsoni* (Whiteaves) ♀ & ♂ [microconch new], *Chondroceras oblatum* (Whiteaves) ♀ & ♂ [microconch new] and *C. defontii* (McLearn) ♀ & ♂ [microconch new]. The microconchiate *Itinsaites* McLearn, 1927 becomes, therefore, synonymous with *Stephanoceras* Waagen, 1869: *Kanastephanus* McLearn, 1927 becomes synonymous with *Zemistephanus* McLearn, 1927. The new species *Zemistephanus alaskensis* is described.

The following Bajocian assemblage zones and subzones are formally distinguished for parts or all of western North America (from top):

Megasphaeroceras rotundum Zone. — *Leptosphinctes, Sphaeroceras, Lissoceras* and late Stephanoceratinae. Early Late Bajocian, *c.* Subfurcatum Chronozone.

Chondroceras oblatum Zone. — *Stephanoceras itinsae*, abundant *S. (Stemmatoceras)* and some *Teloceras*. Late Early Bajocian, late Humphriesianum Chronozone.

Stephanoceras kirschneri Zone. — late Early Bajocian.

Zemistephanus richardsoni Subzone. — *S. kirschneri, Zemistephanus, Chondroceras defontii, Poecilomorphus* and *Dorsetensia*. Early Humphriesianum Chronozone.

(Unnamed Subzone). — *S. (Skirroceras) juhlei, Asthenoceras* and early *Dorsetensia*. Late Sauzei Chronozone.

Parabigotites crassicostatus Zone. — *Emileia, Sonninia (Papilliceras), Stephanoceras* and *Arkelloceras*. Middle Early Bajocian, early Sauzei Chronozone.

Docidoceras widebayense Zone. — earliest Bajocian.

Witchellia sutneroides Subzone. — *c.* Ovalis Chronozone.

Docidoceras camachoi Subzone. — *D. (Pseudocidoceras)* and *Sonninia (Euhoploceras)*. *c.* Discites Chronozone.

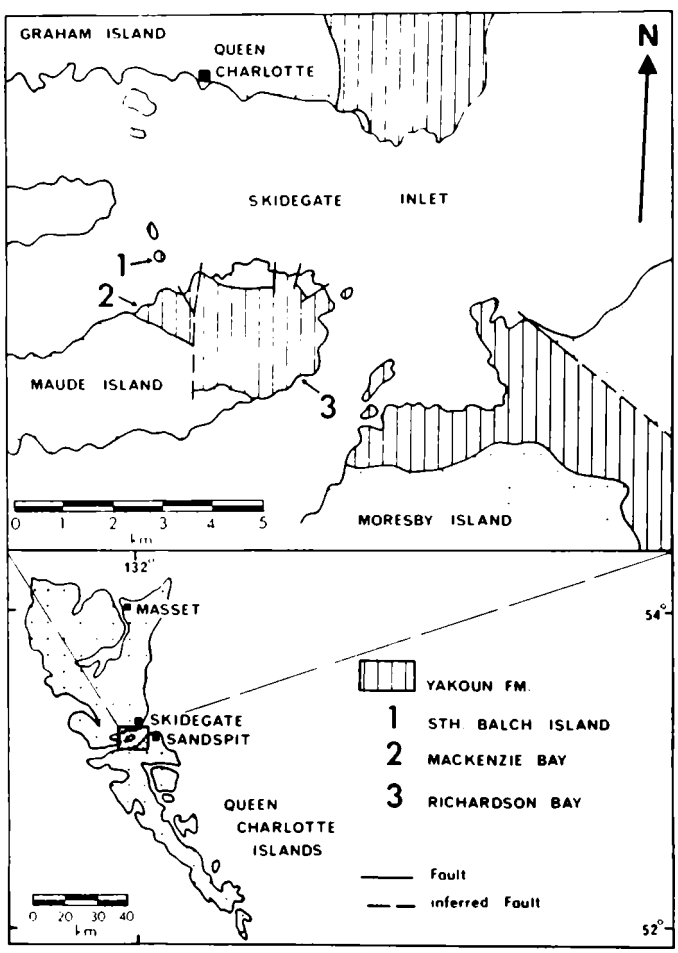
INTRODUCTION

Early failures to separate Jurassic and Cretaceous faunas from the Queen Charlotte Islands were the result of mixing of collections thought to have come from the same strata exposed on several islands in Skidegate Inlet (Whiteaves, 1876, p. 6). MacKenzie (1916) recognised that this confusion had arisen because of failure to separate sandstones of the Cretaceous Haida Formation from the underlying Middle Jurassic Yakoun Formation.

The presence of two distinct ammonite faunas within the Yakoun Formation was shown by McLearn's careful collecting from several measured stratigraphic sections (McLearn, 1927, 1949). Ammonites from three localities within the lower volcaniclastic parts of the Yakoun Formation (Richardson Bay, MacKenzie Bay and South Balch Island; see Text-fig. 1) were presumed to represent a single ammonite faunule belonging to the *Stephanoceras humphriesianum* Zone (McLearn, 1949; Imlay, 1964, p. B19). Arkell (1956, p. 542), however, suggested that the fauna from MacKenzie Bay was older and possibly represented the *Otoites sauzei* Zone; this was based on the misidentification of *Zemistephanus carlottensis* (Whiteaves) ♀ with *Pseudotoites*.

One of us [Hall] made new collections from the Queen Charlotte Islands and parts of western Alberta during the summer of 1971 to (1) clarify the faunal associations at each of the three localities in the lower parts of the Yakoun Formation where the supposed "Lower Yakoun or *Stephanoceras* fauna" was present; (2) determine the stratigraphic sequence and ages of these ammonite faunas and attempt to correlate them with similar faunas from Alaska, mainland western Canada and United States, Chile and Argentina; and (3) obtain sufficient material for biometric analysis in an attempt to demonstrate infraspecific morphological identity of the juvenile stages of several sexual dimorphs and to allow pairing at the species level.

The lower Yakoun Formation is exposed along the shores of Skidegate Inlet, Queen Charlotte Islands and on several small islands in the Inlet (Text-fig. 1). Lower Bajocian ammonites have been recorded and described from three localities: Maude Island, at Richardson Bay on the southeast shore and MacKenzie Bay on the northwest shore, and on South Balch Island. These localities are easily accessible by small boat from Queen Charlotte City on the north shore of Skidegate Inlet. Detailed measurement of each section, and *in situ* collecting of fossils, were attempts to establish the stratigraphic and geographic distributions of ammonites. Special attention was given to forms believed to represent dimorphic pairs.



Text-figure 1. — Index map for the three main fossil localities in Skidegate Inlet of Queen Charlotte Islands, British Columbia.

Stratigraphic measurement, correlation and fossil collecting were hampered by some physical features of these islands. The average tidal range of approximately 7.5 m commonly restricts work along the shore to short periods each day. Much of the exposed platform is covered by seaweed and other marine life, which makes collecting difficult. Many of the fossils, covered by water much of the time, are badly weathered and difficult to remove or transport. Although exposure on the wide, wave-cut platforms is good, sections are often interrupted by Recent beach deposits, ranging from sand and pebbles to large boulders. Thick forests of pine, cedar and hemlock with a thick undercover of mosses, ferns and fallen trees cover the islands to within a few meters of the high tide level, so that tracing beds across even the smallest islands is impossible. Streams exposing rock outcrops occur only on the largest islands. Many small faults that break up the sections cannot be traced laterally for any distance.

Access to isolated outcrops of the Yakoun Formation in the interior of the two largest islands (Graham and Moresby) is by private forestry roads. Outcrops yielding specimens of Lower Bajocian age are not easily located and are of no biostratigraphic value for the lower part of the Yakoun Formation.

Collections were made by both of us from the Rock Creek Member of the Fernie Group in Ribbon Creek, southern Alberta. This outcrop is well exposed and easily accessible from the Kananaskis-Coleman road.

Our material was supplemented by collections of A. Sutherland Brown made between 1958 and 1965, and by some undescribed specimens collected several decades ago by F. H. McLearn; this material was made available on loan from the collections of the Geological Survey of Canada in Ottawa. Two specimens of *Stephanoceras itinsae* (McLearn) ♀ known to have come from Skidegate Inlet (the lithology of the matrix is identical with that on specimens collected by Hall from South Balch Island) were loaned by the Geology Department museum at the University of British Columbia. Comparative material from southern Alaska was obtained on loan from the U.S. National Muse-

um in Washington, D.C. Specimens from the Rock Creek Member collected from several scattered localities in northern Alberta and kept in the collections of the Geology Department, University of Alberta, were examined, but most lacked information on locality, associated fauna and precise stratigraphy.

The holotypes and other figured specimens of species based on material from the Queen Charlotte Islands, western Alberta and southern Alaska were re-examined; when such material was not easily accessible, plaster molds were obtained.

The Lower Bajocian includes the European Standard Zones of *Hyperlioceras discites* to *Stephanoceras humphriesianum*. Our scheme follows that most recently developed by Mouterde and others (1971, pp. 9-13) for France, by Parsons (1974, table 1, p. 154) and Morton (1975, table 1, p. 43) for Great Britain and by Sturani (1967) for Central and Western Europe.

ACKNOWLEDGEMENTS

Access to type collections was kindly provided by Dr. M. K. Howarth (British Museum of Natural History), Dr. I. Cooke (British Geological Survey Museum) and Dr. T. Bolton (Geological Survey of Canada). Dr. C. Stelck (University of Alberta), Dr. R. Imlay (United States Geological Survey), and Dr. A. Sutherland Brown (Geological Survey of British Columbia) made available collections from western Alberta, southern Alaska and British Columbia, respectively. Drs. Bolton and Stelck also provided photographs of type specimens. Helpful discussions with Dr. J. Callomon and Mr. C. Parsons, England, and Prof. J. Wiedmann, C. F. R., are also acknowledged. Mr. J. Whorwood and Mr. R. Larush assisted with the photography.

Financial assistance for this project was provided by the National Research Council of Canada and McMaster University, Hamilton, Ontario, Canada. Funds to defray some costs of publication were provided by the University of Calgary.

REPOSITORIES

Material described or figured here is kept in one of six repository institutions. Catalogue numbers of these institutions are cited in the text by means of the following abbreviations:

McM — Department of Geology paleontology collections, McMaster University, Hamilton, Ontario.

GSC — Geological Survey of Canada collections, Ottawa; includes material collected by Richardson, McLearn, Sutherland Brown and Frebold.

UBC — University of British Columbia, Geology Department museum.

USNM — U.S. National Museum, Washington, D.C.; from southern Alaska. Unfigured material from southern Alaska is identified by the U.S. Geological Survey Mesozoic locality numbers with decimal digits added by us for identification of individual specimens from each locality.

UA — University of Alberta, Edmonton; Geology Department.

STRATIGRAPHY

QUEEN CHARLOTTE ISLANDS

Previous Work

Early geological exploration of the Mesozoic formations of Skidegate Inlet (Pl. 16, fig. 1) was prompted by commercial interest in the coal-bearing rocks of the area. J. Richardson first examined the broad synclinal structure along the Inlet and nearby islands in 1873 and divided the rocks into three horizons (Table 1). Believing the shales on several islands to represent the same geological horizon he did not think it necessary to maintain separate fossil collections (Whiteaves, 1876, p. 6). In describing these fossils Billings (1873) inferred that more than one fauna was present. Whiteaves (1876, pp. 87, 88) also noted "an apparent mixture of Oolitic and Cretaceous types". He recognised that *Ammonites Richardsoni*, *A. Carlottensis*, *A. Skidegatensis* and *A. Loganianus* 'forms A and B' resembled "Oolitic" forms from England (Middle - Upper Jurassic), but added (p. 91) that this similarity "is often of a very general character and can scarcely in any case be shown to amount to actual specific identity." Among Richardson's specimens were several species known to be Cretaceous and so Whiteaves (p. 9) concluded that only one fauna was present and that it represented "a blending of the life of the Cretaceous period with that of the Jurassic".

Richardson's failure to separate Cretaceous rocks (now Haida Formation) from Jurassic rocks (now Yakoun Formation) and the continued insistence by Whiteaves (1876, 1884, 1900) that only one fauna was present confused later workers.

More extensive work by Dawson (1880) resulted in division of the "Cretaceous" sequence into five units (Table 1). Whiteaves (1884) described the fossils collected by Dawson and again noted four ammonites from the "lower

RICHARDSON 1873	DAWSON 1880 1889	ELLS 1906	CLAPP 1914	MACKENZIE 1916	SUTHERLAND BROWN 1968	AGE
upper shale and sandstone	upper shales and sandstone 1500 A		Skidegate member		Skidegate Formation 2000' (Queen Charlotte Group)	
coarse conglomerate	conglomerate 2000 B		Honna member (Queen Charlotte Series)		Honna Formation 1300' – 1400' (Queen Charlotte Group)	Turonian ? Cenomanian
lower shale with coal and iron ore	Queen Charlotte Islands formation: lower shales and sandstones with coal 5000' C		Haida member (Queen Charlotte Series)		Haida Formation 3800' (Queen Charlotte Group)	Cenomanian Albian
					Longarm Formation 4000' (Queen Charlotte Group)	Barremian Hauterivian late Valanginian
	agglomerates 3500' D		Image member (Queen Charlotte Series)	Yakoun Formation	Yakoun Formation 3000'-6000' (Vancouver Group)	early Callovian early Bajocian
	Queen Charlotte Islands formation: lower sandstone 1000' E		Vancouver Group	Maude Formation	Maude Formation 600' (Vancouver Group)	Toarcian Pliensbachian late Sinemurian

Table 1. — History of the use of lithostratigraphic units for the Jurassic - Cretaceous rocks of the Skidegate Inlet area, Queen Charlotte Islands. Heavy line represents position of the Jurassic - Cretaceous boundary.

shales" (Unit C of Dawson) that he stated would be regarded as Jurassic in Europe: *Ammonites Richardsoni*, *Stephanoceras oblatum* [= *Ammonites Loganianus* form A], *S. cepoides* [= *Ammonites Loganianus* form B] and *Perisphinctes Skidegatensis*. However, the association in the collection of these and known Cretaceous forms again led Whiteaves (1884) to suggest a lower Middle Cretaceous age. Stanton and Martin (1905) noted the resemblance of early Middle Jurassic fossils from southern Alaska to certain forms in the "lower shales" of the Queen Charlotte Islands; they believed the latter to have no connection with the Cretaceous faunas supposed to occur in the same formation.

In 1889 Dawson distinguished his Units C, D and E as the Queen Charlotte Islands Formation, believing them to represent a continuous sequence which rested unconformably on older (probably Triassic) rocks. Unit D was separated

from the overlying coal-bearing beds by Ells (1906) and given a pre-Cretaceous age. However, Clapp (1914) extended the Queen Charlotte Series to include Unit D, believing it to be a basal conglomerate of local development, conformable with the overlying units, but resting unconformably on "metamorphic volcanic rocks which seem to belong to the Vancouver Group" (p. 12). Argillites and sandstones, shown by field relations to lie unconformably below the Queen Charlotte Series, were presumed to be of Jurassic or Triassic age. Clapp formally named the Skidegate, Honna and Haida Members for Dawson's subdivisions A, B and C.

The Jurassic of Skidegate Inlet was finally divided into two formations by MacKenzie (1916): the lower Maude Formation (Dawson's Unit E) and the overlying Yakoun Formation (Dawson's Unit D and those parts of Unit C which contained Jurassic fossils). MacKenzie recognised that the strong lithological similarities between the Cretaceous Haida Formation and parts of the Jurassic Yakoun Formation had been responsible for Dawson's failure to separate them and the faunas they contained. He stressed the importance of the unconformity between these two formations.

Detailed mapping and careful collecting with attention to stratigraphy by McLearn in 1921 led to the recognition of several faunal horizons within the Jurassic formations of Skidegate Inlet. His descriptions of the stratigraphy and ammonite faunas of the Yakoun Formation (McLearn, 1927, 1929, 1932a, 1949) led to the distinction of two faunas: the Lower Yakoun or *Stephanoceras* fauna of early Bajocian age and the Upper Yakoun or "*Seymourites*" fauna of early Callovian age, separated by a thick sequence of unfossiliferous volcanic agglomerates and tuffs. A section along the southern shore of Maude Island from Richardson Bay to Robber Point was recorded in detail; it has recently been remeasured and designated the type section of the Yakoun Formation by Sutherland Brown (1968, pp. 68, 72, 73).

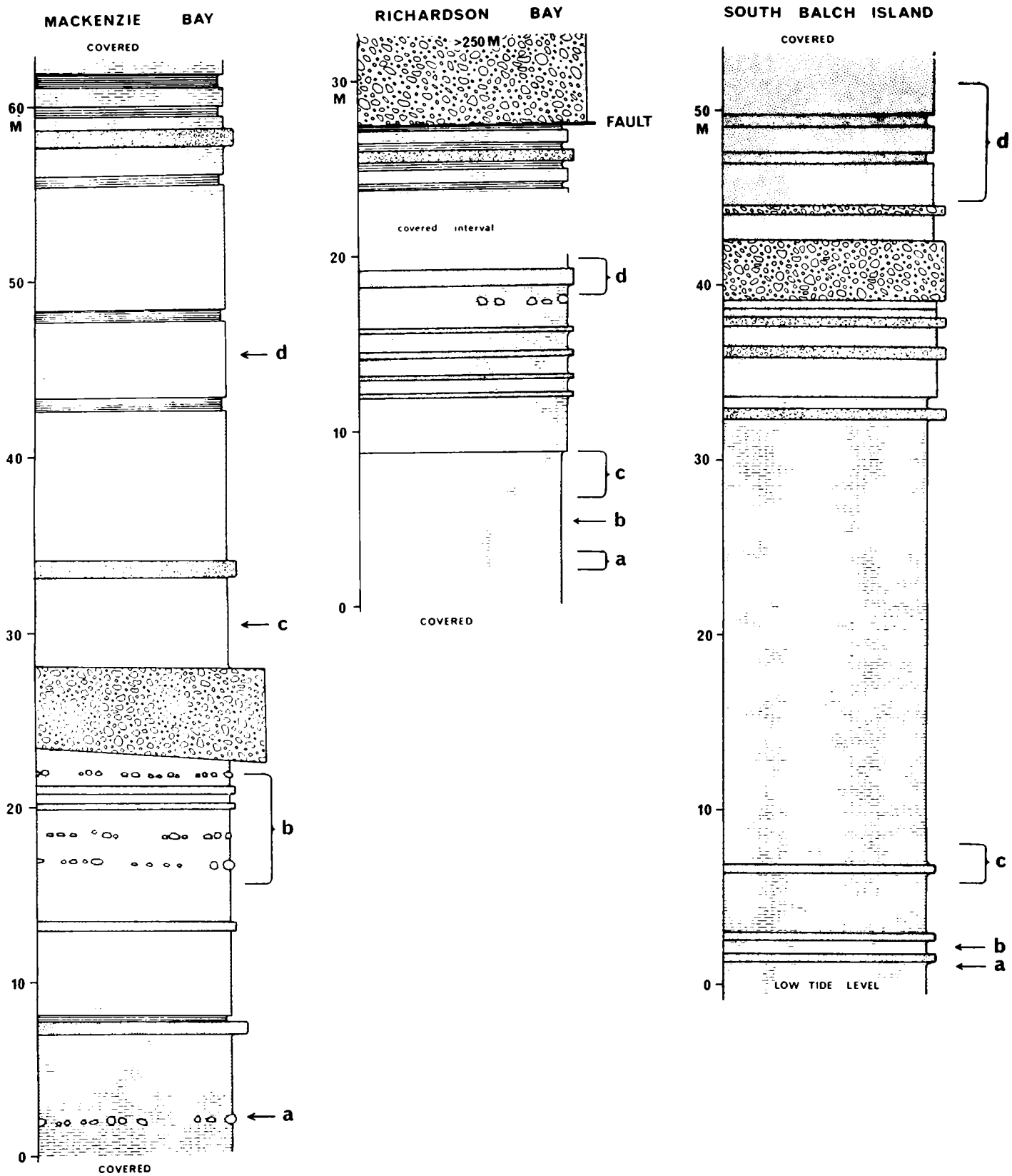
Yakoun Formation

The Yakoun Formation is comprised of rocks ranging from massive volcanic agglomerates and tuffs to volcanic sandstones, shales and siltstones. The type section on the southeastern shore of Maude Island is about 915 m thick and has been subdivided into five members (Sutherland Brown, 1968):

E member	140 m, interbedded massive greenish volcanic sandstone with pebbly beds and grey shales and siltstones. — Callovian
D member	240 m, dominantly finely crystalline lithic tuffs with minor volcanic sandstone.
C member	290 m, coarse, porphyritic andesite and conglomerate.
B member	30 m, interbedded shales and tuffaceous sandstones. — Bajocian.
A member	200 m, fine porphyritic andesite agglomerate and tuffs with white calcareous cement.

The rocks present in any one section of the Yakoun Formation vary considerably. Sutherland Brown (1968) has delineated a "facies front" which marks the zone of transition from dominantly volcanic agglomerate in the east to tuffs, volcanic sandstones and shales in the west. The type section is in this zone of transition so that all lithologies are present in abundance. Clastic rocks in the sequence are composed predominantly of fragments derived from porphyritic andesite. The agglomerates, tuffs and lapilli tuffs contain mainly blocks and fragments of porphyritic andesite, about 20 percent crystal fragments and 20 percent fine matrix that commonly is considerably altered. Related conglomerates are also composed almost entirely of rounded volcanic rock fragments. The volcanic sandstones are made up of subangular fragments of porphyritic andesite and angular crystal fragments in a chloritic matrix. Sutherland Brown envisaged the agglomerates and tuffs originating from a series of vents along a line on the eastern edge of the islands. Pulses of volcanic activity were separated by periods of marine sedimentation with redistribution of volcanic detritus in a shallow, neighbouring sea. The abundance of leaves, fruit and wood fragments in some of these marine strata suggests the proximity of a vegetation-covered land surface. A period of non-marine accumulation resulted in the formation of coal beds within the volcanic sandstones of the Yakoun River valley.

MacKenzie Bay (Pl. 16, fig. 2). — Both the upper and lower boundaries of the B member are covered by Recent beach deposits. The basal part of the exposed section consists of just over 7 m of fine, dark shales, highly fragmented and with many small (2 to 6 cm), hard, rounded or elongated calcareous concretions. Several thin bands of grey-green sandstone show graded bedding. From one horizon (*a* in Text-fig. 2) near the base (exposed only at low tide) come large, rounded calcareous concretions (10 to 15 cm), about half of which contain *Zemistephanus richardsoni* (Whiteaves) ♀ & ♂ and *Z. crickmayi* (McLearn) ♂. *Chondroceras* sp., a few bivalves and belemnites were collected from the associated shales. Abruptly overlying the shales is a thin but very prominent bed (60 cm thick) of buff-coloured volcanic sandstone with some minor bands of angular lithic fragments. Scour-and-fill structures are abundant, and there is some cross-bedding. Load structures occur along the base of the unit.



Text-figure 2. — Measured sections and fossil occurrences in the lower Yakoun Formation of Skidegate Inlet, Queen Charlotte Islands; see Table 2 for listing of fossil assemblages.

The succeeding 15 m of grey shale with sandy interbeds (*b* in Text-fig. 2) are sparsely fossiliferous, having yielded single specimens of *Chondroceras* sp., *Zemistephanus* ♀ sp., together with *Stephanoceras*? (*Stemmatoceras*?) ex gr. *S. acuticostatum* (Weisert) ["*Teloceras itinsae*" McLearn], a few belemnites, bivalves and brachiopods. This unit is well bedded with conspicuous bands of elongated calcareous concretions. Cutting unconformably across the top of these beds is a massive, wedge-shaped unit, 5 m thick, of green volcanic sandstone and agglomerate.

In the overlying 35 m of poorly-bedded grey, silty shale a few fossils were found at two horizons. From just 2 m above the agglomerate (*c* in Text-fig. 2) come several *Zemistephanus alaskensis* n. sp. ♀ and 15 m higher (*d* in Text-fig. 2) additional *Z. richardsoni* (Whiteaves) ♀ & ♂, together with rare bivalves, belemnites and some worm borings. Some grey-black argillite bands give the only evidence of bedding.

Table 2. — Ammonite species found in the stratigraphic sections of Skidegate Inlet, Queen Charlotte Islands (Text-fig. 2). Dimorph is indicated, and number of specimens appears in ().

Locality	Bed	Ammonite species
Richardson Bay	*d*	*Stephanoceras skidegatense* (Whiteaves) ♂ (7) & ♀ (7)
		Stephanoceras sp. aff. *S. skidegatense* ♂ (1)
		Chondroceras defontii (McLearn) ♂ (1)
	c	*Stephanoceras skidegatense* (Whiteaves) ♂ (1) & ♀ (3)
		Stephanoceras sp. aff. *S. skidegatense* (Whiteaves) ♂ (1)
		Chondroceras defontii (McLearn) ♂ (1) & ♀ (7)
	b	*Stephanoceras skidegatense* (Whiteaves) ♀ (1)
	a	*Stephanoceras skidegatense* (Whiteaves) ♂ (1) & ♀ (1)
		Chondroceras defontii (McLearn) ♀ (1)
South Balch Island	*d*	*Stephanoceras itinsae* (McLearn) ♂ (3) & ♀ (15)
		Chondroceras oblatum (Whiteaves) ♂ (1) & ♀ (3)
		Cenoceras sp. (2)
		Calliphylloceras sp. (1)
	c	*Stephanoceras itinsae* (McLearn) ♂ (4) & ♀ (10)
		Chondroceras oblatum (Whiteaves) ♀ (1)
	b	*Stephanoceras itinsae* (McLearn) ♀ (5) & juv. (5)
		Chondroceras oblatum (Whiteaves) ♀ (2)
		Calliphylloceras sp. (1)
	a	*Stephanoceras itinsae* (McLearn) ♀ (1)
		Chondroceras oblatum (Whiteaves) ♀ (2)
MacKenzie Bay	*d*	*Zemistephanus richardsoni* (Whiteaves) ♂ (1) & ♀ (3)
		Zemistephanus sp. indet. juv. (1)
	c	*Zemistephanus alaskensis* n. sp. ♀ (2)
	b	*Zemistephanus* sp. indet. ♀ (1) & juv. (1)
		Chondroceras n. sp. indet. ♀ (1)
		"*Teloceras itinsae*" McLearn juv. (2)
	a	*Zemistephanus richardsoni* (Whiteaves) ♂ (4) & ♀ (15)
		Zemistephanus crickmayi (McLearn) ♂ (7)
		Chondroceras n. sp. indet. ♀ (2)

Richardson Bay (Pl. 16, fig. 3). — The base of the section is covered by Recent beach deposits. The lower 9 m consist of massive, poorly-bedded shales with some white feldspar grains. From several broad and poorly-defined horizons (*a-c*) come specimens of *Stephanoceras skidegatense* (Whiteaves) ♀ & ♂ and abundant *Chondroceras defontii* (McLearn) ♀ & ♂. Above these are 7.5 m of grey, highly-fractured sandy shale with occasional thin bands of pebbly tuff and some lightly contorted bands of finer, yellowish-green tuff.

From the succeeding 3.5 m of dark, grey brown shale (*d*) additional specimens of *S. skidegatense* (Whiteaves) ♀ & ♂ and *C. defontii* (McLearn) ♂, some bivalves, gastropods and belemnites were collected. This unit also contains abundant hard, calcareous concretions of varying size and fragments of carbonized logs and branches.

Above a covered beach interval of 4 m is a thin succession of bands of alternating brown-grey sandy shales, blue-grey cherty argillite and agglomerates which are highly contorted and broken apart. These are faulted against massive brown, volcanic agglomerate containing some large blocks of the underlying contorted sediments. This fault represents the base of the thick, massive agglomerate designated the C member by Sutherland Brown (1968, p. 73).

South Balch Island (Pl. 16, fig. 4). — The western shore of this small island has an apparently continuous exposure of about 51 m through the B member. Beginning on the southern shore is a sequence of 6 m of fine, grey shales with three prominent beds of grey-green sandstone all showing minor fault displacements. Fossils in this sequence (*a-c*) include abundant *Stephanoceras itinsae* (McLearn) ♀ & ♂, some *Chondroceras oblatum* (Whiteaves) ♀ & ♂, bivalves, a few brachiopods and fossil fruit.

Next comes a barren, monotonous sequence of grey shales 25.5 m thick extending along the western shoreline and largely covered by loose boulders. It has yielded several bivalves and gastropods and some fossil fruit. Towards the top of the sequence massive volcanic agglomerate and sandstone become the dominant lithologies with a few interbeds of sandy shales, one of which (*d*) yielded the *S. itinsae* - *C. oblatum* fauna, along with single specimens of ?*Eutrephoceras* sp., *Cenoceras* sp. and *Calliphylloceras* sp. and rare bivalves.

The eastern side of the island is a large fault block cut by several smaller faults. It is composed mostly of massive, coarse volcanic sandstones, grey-green and brown in colour. There are minor dark grey shales. In addition to scattered specimens of *S. itinsae* (McLearn) ♀ & ♂, *C. oblatum* (Whiteaves) ♀ & ♂ and a single *Calliphylloceras* sp., these

PREVIOUSLY IDENTIFIED TAXA		REVISED AND NEW TAXA
1. MacKenzie Bay		
Zemistephanus richardsoni (Whiteaves)		*Zemistephanus richardsoni* (Whiteaves) ♀
Z. funteri McLearn		
		Z. richardsoni (Whiteaves) ♂ **
Z. carlottensis (Whiteaves)		*Z. carlottensis* (Whiteaves) ♀
Z. vancouveri McLearn		
		Z. alaskensis n. sp. ♀ **
Normannites (Kanastephanus) crickmayi (McLearn)		
N. (K.) canadensis (McLearn)		*Z. crickmayi* (McLearn) ♂
N. (K.) altus (McLearn)		
N. (K.) mackenzii (McLearn)		
Teloceras itinsae McLearn		*Stephanoceras* ex gr. *acuticostatum* Weisert ♀
		Chondroceras n. sp. indet. **
2. Richardson Bay		
Stephanoceras skidegatense (Whiteaves)	.....	*Stephanoceras skidegatense* (Whiteaves) ♀
S. skidegatense var. *laperousii* McLearn		
		S. skidegatense (Whiteaves) ♂ **
		S. sp. ♂ aff. *S. skidegatense* (Whiteaves) ♂ **
Chondroceras (Defonticeras) defontii (McLearn)		
C. (D.) colnetti (McLearn)	.....	*Chondroceras defontii* (McLearn) ♀
C. (D.) ellsi (McLearn)		
C. (D.) marchandi (McLearn)		
		C. defontii (McLearn) ♂ **
C. (D.) maudense (McLearn)		*C. maudense* (McLearn) ♀
3. South Balch Island		
Stephanoceras yakounense McLearn	.....	*Stephanoceras itinsae* (McLearn) ♀
Normannites (Itinsaites) itinsae (McLearn)	.....	*S. itinsae* (McLearn) ♂
Stephanoceras caamanoi McLearn	.....	*S. caamanoi* McLearn ♀
	.....	*Calliphylloceras* sp. indet. **
	.....	?*Eutrephoceras* sp. indet. **
	.....	*Cenoceras* sp. indet. **
Chondroceras (Defonticeras) oblatum (Whiteaves)	.....	*Chondroceras oblatum* (Whiteaves) ♀
	.....	*C. oblatum* (Whiteaves) ♂ **

Table 3. — List of previously recognized, revised and new taxa from the lower part of the Yakoun Formation (** = forms described here for the first time).

sandstones contain many large concretions with concentric internal layering, leaf impressions, fossil fruit, carbonized stems and wood fragments. Scattered throughout the strata on both sides of the island are irregular accumulations, 10-30 cm across, of darker shale containing a concentration of broken ammonite fragments, small bivalve shells and gastropods. Their origin remains unknown.

Ammonite Faunas and Correlations

It is clear from the faunal lists of the three fossil localities in the lower Yakoun Formation (Table 3) that three distinct and separate faunas are present; the problem of their relative ages now arises. Although the section at each locality is thick and apparently continuous, no specimen representing the other two faunas has been found in stratigraphic succession at any locality. McLearn (1949) believed that these three faunas were "apparently all of one fauna, the lower Yakoun or *Stephanoceras* fauna" (p. 16). Imlay (1964, p. B19) also held that the "various species of *Chondroceras, Normannites* and *Teloceras* that were obtained from these localities on Maude Island occur together in the Fitz Creek Siltstone in Alaska and belong to a single faunule". However, as discussed under *South Alaska* (*q.v.*) species of these genera are segregated stratigraphically in the Fitz Creek Siltstone. Arkell (1956, p. 542) suggested on paleontological grounds that the fauna from MacKenzie Bay was distinct from, and older than, the Richardson Bay fauna.

MacKenzie Bay. — Although the holotype of "*Teloceras itinsae*" McLearn from this locality was collected as float, one new larger specimen and one fragment were collected *in situ* in the lower third of the section (*i.e.*, within the range zone of *Zemistephanus richardsoni*). This poorly known species is either a *S. (Stemmatoceras)* of the *acuticostatum* group or, perhaps, a *Teloceras*, and clearly indicates the *S. humphriesianum* Zone.

The *Zemistephanus* fauna from this locality corresponds closely in composition to that from the Fitz Creek Siltstone in southern Alaska (see discussion under *South Alaska*). Species in common are *Z. richardsoni* (Whiteaves) ♀ & ♂, *Z. carlottensis* (Whiteaves) ♀, *Z. crickmayi* (McLearn) ♂ ["*Normannites (Kanastephanus)* spp."], and *Z. alaskensis* n. sp. ♀. Species of *Chondroceras* are also present in both faunas.

Because of discontinuity of outcrop the lithostratigraphic position of these beds relative to those at Richardson Bay and on South Balch Island cannot be established. The only clue to the relative ages of the contained faunas is a small ammonite nucleus (GSC 13636; Pl. 9, fig. 5) that, according to McLearn's handwritten label, was collected by him between MacKenzie and Maude Bays on the north shore of Maude Island. This locality description corresponds to the uppermost part of the section shown for MacKenzie Bay in Text-figure 2, *i.e.* stratigraphically above the *Zemistephanus* horizons. No other material has been recovered from this part of the section. This nucleus bears a very strong resemblance to *Stephanoceras skidegatense* (Whiteaves) (McM J1802f; Pl. 9, fig. 4) collected about 20 m above the exposed base of the section at Richardson Bay. A detailed analysis of similar faunas within the Fitz Creek Siltstone and Cynthia Falls Sandstone of southern Alaska (see discussion under that heading) also suggests that the *Zemistephanus* fauna is older than those from Richardson Bay and South Balch Island which contain abundant *Stephanoceras* but no *Zemistephanus*.

Richardson Bay. — The fauna is characterized by abundant *Chondroceras defontii* (McLearn) ♀ & ♂ and *Stephanoceras skidegatense* (Whiteaves) ♀ & ♂, while *Zemistephanus* is unknown. In southern Alaska *C. defontii* ♀ ranges through most of the Fitz Creek Siltstone where it is associated with *Zemistephanus* and persists into the overlying Cynthia Falls Sandstone where *Z. richardsoni* is absent and *C. oblatum* ♀ first appears. We suggest that the *S. skidegatense* faunule is probably younger than the *Z. richardsoni* faunule.

South Balch Island. — The *Stephanoceras (Stephanoceras) itinsae* (McLearn) ♀ from this locality is very similar to European species group of *S. umbilicum* and *S. mutabile;* these forms in western Europe first appear and reach maximum numbers in the middle parts of the *S. humphriesianum* Zone (Schmidtill and Krumbeck, 1938; Mouterde *et al,* 1971). *S. itinsae* ♀, which is so abundant in the Bajocian faunas of the Queen Charlotte Islands and Alberta, is not known from the thick sequences of southern Alaska where the *Zemistephanus* faunule is dominant. *Stephanoceras obesum* Imlay from the Tuxedni Formation in the Talkeetna Mountains of southern Alaska (Imlay, 1964, pp. B45, 46; pl. 18, figs. 5-11) is a similar but apparently older form (see discussion under subgenus *Stephanoceras* Waagen, 1869). In the Cynthia Falls Sandstone of southern Alaska, however, *Chondroceras oblatum* (Whiteaves) occurs with an undescribed species of *Stephanoceras* (Imlay, 1964, table 7). *Chondroceras* cf. *oblatum* (Whiteaves) and *C. defontii* (McLearn) are listed from the Cynthia Falls Sandstone (table 7), but do not occur at the same locality, a feature also noted in the lower Yakoun Formation. Whether such faunal segregation is due to some ecological factor or difference in age is not clear.

Conclusions. — It is tentatively concluded that the *S. itinsae* - *C. oblatum* faunule from South Balch Island is younger than the *Zemistephanus* faunule at MacKenzie Bay and that they correlate respectively, to the middle to upper and lower *S. humphriesianum* Standard Zone of Europe (Table 4).

MAINLAND BRITISH COLUMBIA

Stephanoceratid and sphaeroceratid ammonites have been recorded from the thick volcaniclastic sequences comprising the Laberge Group in northern British Columbia and the Hazelton Group in central and central-western British Columbia. Knowledge of the systematics and stratigraphic occurrence of these forms is, however, minimal because of poor preservation and widely scattered localities; much work remains to be done on the Middle Jurassic ammonites of this region.

In southwestern British Columbia, *S. (Skirroceras)* cf. *kirschneri* Imlay occurs in the Taseko Lakes area, but cannot be dated independently since the associated "*Witchellia* sp." cannot be identified at the generic level (Frebold *et al.*, 1969, pl. 4, fig. 7). That fragment could, among others be a *Sonninia (Papilliceras)* of the European *O. sauzei* Zone or a *Dorsetensia* of the upper *O. sauzei* to *S. humphriesianum* Zones. Better evidence for the age of *Zemistephanus* was furnished by the Lookout Section in Manning Park, 180 km due east of Vancouver (Frebold *et al.*, 1969, p. 18). The isolated occurrence of *Chondroceras "marshalli"* (McLearn) [= *C. oblatum*] in the same area indicated the *S. humphriesianum* Zone.

Lookout Section. — Frebold *et al.* (1969, p. 19) described a thick sequence of deep-water turbidites which yielded, from the base (1) "Otoitidae", (2) a faunule with "*Graphoceras crickmayi* n. sp." and *Zemistephanus richardsoni,* (3) *Chondroceras* spp. indet. and unidentified ammonites, and (4) an upper faunule with *Chondroceras* aff.

ellsi (McLearn) and *Stephanoceras* cf. *S. caamanoi* (McLearn). While the large figured specimen of *Z. richardsoni* was correctly identified, the associated alleged new species of the late Aalenian *Graphoceras* was based on 3 fragmentary specimens with unknown whorl section, *i.e.*, two one-sided impressions (including the holotype) and one crushed specimen.

Imlay (1973, p. 75) had already noted the close similarity of "*Graphoceras crickmayi*" Frebold *et al.* (1969) to *Poecilomorphus* "*varius*" Imlay (1973), stating that they differ only in the slightly concave flanks of the former. However, this feature may well be the result of partial crushing, and the species erected by Frebold and Imlay cannot be discriminated on objective grounds. However, the type-material of "*P. varius*" Imlay is very poor and both "species" are a close match to, and may well be conspecific with, *P. cycloides* (d'Orbigny). An excellent series of this variable species, a subzonal index in the basal *S. humphriesianum* Zone from the Alps to England, was figured by Sturani (1964, pl. 8). Whether "*Graphoceras crickmayi*" is conspecific or closely affiliated with *P. cycloides* makes little difference for the time-correlation, since the entire genus *Poecilomorphus* (*sensu* Sturani) is restricted to that subzone. This therefore strongly indicates the age of the occurrence of *Z. richardsoni* as *P. cycloides* subzone [= *D. romani* Subzone], *i.e.*, approximating the age of *Zemistephanus* in eastern Oregon, as discussed below.

ALBERTA

Jurassic rocks ranging in age from Sinemurian to Upper Portlandian comprise the Fernie Group which has been mapped from southernmost Alberta over 1,000 km to the north in the Peace River area (Frebold, 1957, pp. 1, 2; Frebold and Tipper, 1970, pp. 9, 10). The section is incomplete, with many hiatuses indicated by the absence of index fossils.

Warren (1934) proposed the name Rock Creek Member for a calcareous sandstone bed, 1.5 to 9 m thick and 15 to 45 m above the base of the Fernie Shale in southern Alberta. No identifiable ammonites have been found in this unit at its type locality at Rock Creek, but interbedded shales and calcareous sandstones in many other parts of western Alberta with *Stephanoceras*, *S. (Stemmatoceras)*, *Teloceras* and *Chondroceras* spp. and belemnites of the "*Teloceras* fauna" (Warren, 1934) [*S. humphriesianum* Chronozone] are considered correlative (Frebold, 1976). This rock sequence continues downward, containing earlier ammonite assemblages in some places. Frebold (1976, p. 4) has recently proposed extending the Rock Creek Member to include all strata between the "Toarcian paper shale" and Callovian "grey beds"; ammonites representing both the "*S. sowerbyi*" and *S. humphriesianum* Chronozones are known from this interval. This extension of the Rock Creek Member from a single sandstone bed, as originally defined by Warren (1934), to include a sequence of calcareous sandstones, shales and limestones is not desirable.

While the "*Teloceras* fauna" has been recorded from numerous localities throughout western Alberta (Warren, 1947; McLearn, 1927, 1928, 1930, 1932b; Frebold, 1957), the European *O. sauzei* Chronozone is rarely indicated, as it is at the Snake Indian River locality (below).

Ribbon Creek.—New collections of the "*Teloceras* fauna" were made separately by both of us at the Ribbon Creek locality of the Rock Creek Member proper; details of this section were given by Frebold (1957, pp. 81, 28).

The best preserved and most abundant forms at this locality are *Chondroceras allani* (McLearn) ♀ and *C. oblatum* (Whiteaves) ♀ ["*C. marshalli* (McLearn)" ♀]. "*Zemistephanus*" *crickmayi* Frebold (1957, pp. 52, 53; pl. XXV; pl. XXVI; pl. XXVII) is here referred to *Teloceras* [*Teloceras crickmayi* (Frebold, 1957) ♀ (*q.v.*)]; we have found other larger, fragmentary specimens of this species which may be the only North American true *Teloceras*. The holotype of "*Teloceras*" [*S. (Stemmatoceras)*] *dowlingi* McLearn is also assumed by Frebold (1957, p. 52) to have come from this locality. "*Stemmatoceras albertense*" McLearn of Frebold (1957, p. 50) is probably another *Stephanoceras itinsae* (McLearn) ♀ [*Stephanoceras (Stephanoceras) itinsae* (McLearn, 1927) ♀ & ♂ (*q.v.*)]; other fragmentary material from this locality is also identified with this species (Pl. 8, fig. 1). The corresponding microconch, "*Itinsaites itinsae*" McLearn, has also been collected from this locality (McM J1838; Pl. 8, fig. 7), along with numerous possibly conspecific stephanoceratid nuclei.

The association of *Stephanoceras itinsae* ♀ & ♂ with *Chondroceras oblatum* ♀ indicates that the Rock Creek Member exposed at Ribbon Creek is correlative with the lower Yakoun Formation of South Balch Island (see discussion under *Queen Charlotte Islands*). The presence of *Teloceras*, together with several *S. (Stemmatoceras)*, and the absence of *Zemistephanus*, suggest correlation with the upper parts of the *S. humphriesianum* Standard Zone of western Europe. Similar *S. (Stemmatoceras)* with rather narrow coronate inner whorls, are reported from just below the *T. blagdeni* Subzone in France (Mouterde *et al.*, 1971, pp. 11, 12) and in England (Parsons, *pers. comm.*). Only

Teloceras crickmayi (Frebold) closely resembles typical *Teloceras* of the *T. blagdeni* Subzone.

Snake Indian River. — Two faunules from two horizons 18 m apart are present in the fine section of the Rock Creek Member (*s.l.*) exposed in Snake Indian River valley (Frebold, 1957, pl. 7A). The "*Stephanoceras* ex. gr. *skidegatense*" described by Frebold (1957, p. 49; pl. 22, fig. 2; pl. 25, fig. 2) is probably a *S. itinsae* (McLearn) ♀ [*Stephanoceras (Stephanoceras) itinsae* (McLearn, 1927) ♀ & ♂ (*q.v.*)]. Unfortunately, its exact stratigraphic horizon is not recorded but is assumed to be the upper horizon, the "Rock Creek Member *sensu stricto*" with a belemnite-bivalve assemblage; "stephanoceratids" occur at several levels in this section (Frebold, 1957, p. 92). Westermann (1964b) collected from the sparse lower faunule (base of sandstone with large concretions) a single *Arkelloceras* n. sp. aff. *A. mclearni*, together with fragments identified with *S. (Stemmatoceras)* aff. *S. frechi* (Renz) and "*Kumatostephanus*" cf. *turgidulum* (Quenstedt), resembling the "*Stephanoceras* sp. indet." of Frebold (1957, p. 50, pl. 22, fig. 1). He dated this assemblage as *O. sauzei* Chronozone. This age for *Arkelloceras* was independently demonstrated by Imlay (1964, pp. B53, 54) in the Kialagvik Formation of the Alaska Peninsula. The occurrence of *Arkelloceras* in the Kingak Shale of north Alaska above the last *Pseudolioceras* also supports this age (Imlay, 1976). The "*Kumatostephanus*" specimens from Snake Indian River consist only of body chambers and closely resemble *Stephanoceras (Skirroceras?) nelchinanum* (Imlay) from the *P. crassicostatus* Zone of southern Alaska.

SOUTH ALASKA

Along the Alaskan Peninsula, early Bajocian faunas correlated with the "*S. sowerbyi* Zone" have been described from the Kialagvik Formation by Westermann (1969b). This is followed by the same *Parabigotites* faunule as at the top of the Red Glacier Formation in the Cook Inlet region (Imlay, 1964).

A richly fossiliferous upper Lower Bajocian sequence is present in the Tuxedni Group of the Cook Inlet region and Talkeetna Mountains, South Alaska. Correlation with the western European Standard Zones, however, is hampered by the scarcity of Sonniniidae (especially *Dorsetensia, Sonninia* and *Poecilomorphus*) above the Red Glacier Formation and of the familiar Stephanoceratidae so common in the *O. sauzei* and *S. humphriesianum* Zones of Europe, i.e., *Stephanoceras* s.s., *S. (Stemmatoceras)*, "*Normannites*" and *Teloceras*. The great and varying thicknesses (180-1200 m) of these sequences of massive silty shales, sandstones and greywackes with some massive conglomerates and the presence of major thrust faults make detailed litho- and biostratigraphy difficult (Imlay, 1964).

The upper parts of the Red Glacier Formation were correlated with the upper *O. sauzei* Zone because "*Normannites*" and *Chondroceras* are not known from below this level (Imlay, 1964, p. B12). The specimens of *Chondroceras* cf. *defontii* (McLearn) listed as occurring in the Red Glacier Formation, however, were *ex situ* and the reported species of "*Normannites*" are now referred to other genera. "*N. kialagvikensis*" Imlay is the microconch of *Parabigotites crassicostatus* Imlay (cf. Imlay, 1973, pp. 32, 85, 95); both dimorphs are restricted to the Red Glacier formation where they commonly occur together. "*N. (Itinsaites) crickmayi*" (McLearn) is a microconch form of *Zemistephanus* ♀ [*Zemistephanus crickmayi* (McLearn, 1927) ♂ (*q.v.*)]. Both macroconch and microconch samples of *Chondroceras* have been reported recently from the supposed *W. laeviuscula* Zone of southern England (Parsons, 1974, p. 167) and the genus occurs at a similar level in Argentina (Westermann and Riccardi, unpublished information). This removes the constraints which led Imlay to suggest correlation of this fauna with the upper part of the *O. sauzei* Zone.

In the highest parts of the Red Glacier Formation, *Stephanoceras (Skirroceras?) kirschneri* Imlay ♀ is the dominant species; it is not associated at any of the listed localities with *P. crassicostatus* Imlay ♀ and ♂. *S. kirschneri* is intermediate in coiling between typical *Skirroceras*, mainly of the *O. sauzei* Zone, and the slightly more involute *Stephanoceras* s.s. of the early *S. humphriesianum* Zone; the inner whorls resemble *S. pyritosum* (Quenstedt) known from Europe and from similar levels in the Andes (Westermann, unpublished information). This would suggest that the *S. kirschneri* range zone belongs in the late *O. sauzei* — early *S. humphriesianum* Chronozones, *i.e.*, intermediate between the ranges assumed by Imlay (1964, 1973).

The overlying Gaikema Sandstone, of which only the basal part is fossiliferous, contains *Emileia constricta* Imlay (a close relative of *E. polyschides*, Waagen), *Bradfordia? cariboueensis* Imlay, "*Witchellia*? sp." and *Sonninia (Papilliceras)* cf. *arenata* (Quenstedt). This fauna was originally correctly correlated with the *O. sauzei* Zone (Imlay, 1964). *S. (Papilliceras)* ranges from the upper "Sowerbyi Zone" through the *O. sauzei* Zone in western Europe (Mouterde *et al.*, 1971, p. 11; Parsons, 1974), Chile and Argentina (Westermann and Riccardi, 1972, pp. 73-77). *E. polyschides* has a similar range in Europe (Parsons, 1976), while the other taxa are of dubious affinity and age.

Due to the supposed presence of *"Normannites"* and *Teloceras,* the fauna of the Fitz Creek Siltstone was correlated with the *S. humphriesianum* Zone (Imlay, 1964, p. B12). But none of the figured *"Normannites (Itinsaites) crickmayi"* (McLearn) and *"N. (I.) itinsae"* (McLearn) from this unit is a true *"Normannites"* or *"Itinsaites"*, i.e., *Stephanoceras* microconch; *"Kanastephanus" crickmayi* and its allied forms are microconch *Zemistephanus,* as is the *"N. (I.) itinsae"*. They show neither the rounded whorl section nor the persistence of dense secondary ribbing (three to each primary) right to the aperture, two features which characterize *Stephanoceras itinsae* ♂. Furthermore, the material identified with the poorly known *"Teloceras itinsae"* McLearn (Imlay, 1964, p. B50) belongs to *Zemistephanus alaskensis* n. sp.

The dominant ammonite species in the Fitz Creek Siltstone is *Zemistephanus richardsoni* (Whiteaves). It is restricted to this unit and here proposed as an appropriate guide fossil to replace Imlay's *"Teloceras itinsae". Stephanoceras (Skirroceras) kirschneri* Imlay appears to range to the top of the *Z. richardsoni* range zone (Imlay, 1964, p. B47, Table 8, USGS Mesoz. locs. 26599, 21274).

Chondroceras defontii (McLearn), which forms a strong association with *Zemistephanus* throughout the Fitz Creek Siltstone, does not usually occur with *C. allani* (McLearn); they are reported together only at one locality (USGS Mesoz. loc. 21276) which covers a stratigraphic interval of 10.7 m (Imlay, 1964, table 6). *C. allani* does occur together with *C.* cf. *oblatum* (Whiteaves) which first appears at the base of the overlying Cynthia Falls Sandstone (Imlay, 1964, table 8) and also with *C. "marshalli"* [= *C. oblatum*] in the Tuxedni Formation (Imlay, 1964, table 12). *C. allani* and *C. oblatum* are found associated with undescribed *Stephanoceras* s.s. and *S. (Stemmatoceras),* subgenera with which they commonly occur in the Queen Charlotte Islands and in the *"Teloceras* fauna" of western Alberta (see discussion under *Alberta*). The evidence indicates that the *C. allani - C. oblatum - Stephanoceras* faunule represents a higher horizon than the *Z. richardsoni - C. defontii* faunule and is equivalent to middle and upper parts of the *S. humphriesianum* Zone.

OREGON

Bajocian rocks in east-central Oregon (Imlay, 1973, pp. 8-9, fig. 1) consist of sequences of marine clastics ranging from coarse conglomerates to siltstones and mudstones. Clasts are of volcanic origin and locally there are andesitic lavas and volcanic breccias. These rocks vary greatly in thickness so that lithic units cannot be traced laterally for any great distance. Dickinson and Vigrass (1965) therefore extended the Snowshoe Formation (Lupher, 1941) to include a number of previously named "Formations". In the Suplee area, it is divided, from below, into the Weberg, Warm Springs and Basey Members. Ammonite faunas and lateral continuity of some beds allow correlation of these units with the Snowshoe Formation in the Izee area where it is much thicker and includes strata both older and younger than in the Suplee area.

A westward onlapping sequence of hard, sandy limestone and calcareous sandstone, 15 to 75 m thick, comprises the Weberg Member in the Suplee area (Imlay, 1973, pp. 20-22). The association of *Tmetoceras scissum* (Benecke), *Praestrigites* and *Eudmetoceras* near the base indicates correlation with the European *G. concavum* Zone and the Alaskan *E. howelli* Zone; the middle and upper parts of the Member yield *Sonninia (Euhoploceras) adicra modesta* Buckman, *Docidoceras lupheri* Imlay and the evolute *Witchellia (Latiwitchellia)* spp. indicating the *H. discites* and *S. ovalis* Zones (Imlay, 1973, pp. 20-22).

The overlying Warm Springs Member consists of soft, calcareous, thinly laminated claystones and mudstones, 30 to 90 m thick, and on the basis of faunal similarity may be correlated with the upper parts of the unnamed "Lower Member" in the Izee area. The Lower Member there consists of 150 to 230 m of black and dark grey siltstone, claystone and mudstone with local developments of concretions and thin limestone beds (Imlay, 1973, p. 14). The correlation of the faunas with the middle and upper parts of the "*S. Sowerbyi*" Zone [*S. ovalis* and *W. laeviuscula* Standard Zones] and the *O. sauzei* Standard Zone is based on

> . . . the presence of *Witchellia* throughout the member, the end of occurrences of *Witchellia* and *Papilliceras* at its top, the presence of *Emileia* and *Dorsetensia* in its upper two thirds, the presence of *Normannites* near its top, and the presence of *Fontannesia* near its base (Imlay, 1973, p. 22).

In the upper parts of the Warm Springs Member is an assemblage including the diagnostic ammonites *Parabigotites crassicostatus* Imlay, *Stephanoceras (Skirroceras) juhlei* Imlay, *Witchellia connata* (Buckman), *"Otoites contractus"* (Sowerby) and *Emileia buddenhageni* Imlay, all of which occur together at a single locality (Imlay, 1973, table 6, loc. 78). This fauna correlates with the *Parabigotites* faunule found in the upper parts of the Red Glacier Formation in southern Alaska and has been correctly placed in the *O. sauzei* Chronozone (Imlay, 1964, 1973); it could, however, begin in the *W. laeviuscula* Chronozone. Significantly, *E. buddenhageni* is closely affiliated to *E. giebeli submicro-*

stoma (Gottsche) from the *W. laeviuscula* to lower *O. sauzei* Chronozones of the Andes (Westermann and Riccardi, 1979).

The overlying Basey Member in the Suplee area consists of massive units of volcanic fragmental rocks and interbedded andesitic lavas. It is 760 m thick but thins rapidly to the east, grading into finer clastic rocks that are placed in the Middle Member of the Snowshoe Formation in the Izee area. Here this unit is 305 m thick and consists of thinly bedded, dark grey and black calcareous siltstone and claystone with alternating grey and green sandy siltstones and fine sandstones of volcanic origin. Further east these beds intertongue with conglomeratic volcanic rocks of the Silvies Member (Imlay, 1973, p. 14). According to Imlay the lower fauna resembles that of the underlying Warm Springs Member, *i.e.*, (?) *Dorsetensia* cf. *subtecta* Buckman, *Pelekodites* (?) *silviesensis* Imlay, *S. (Skirroceras) juhlei* Imlay and *Asthenoceras delicatum* Imlay, together with the first *S. (Skirroceras) kirschneri* Imlay and *Dorsetensia* (?) *oregonensis* Imlay. The middle part of the unit has yielded few ammonites but in the upper part *Poecilomorphus* cf. "*varius*" Imlay [= *P. cycloides*?], *Chondroceras allani* (McLearn), "*Normannites*" *orbignyi* Buckman, "*N. (Itinsaites) crickmayi* (McLearn)", [= *Zemistephanus* ♂] and "*Stephanoceras* cf. *nodosum* (Quenstedt)" are said to make their first appearance. The lower fauna was correlated with the *O. sauzei* Zone [here, lower *S. kirschneri* Zone] and the upper fauna with the *S. humphriesianum* Zone (Imlay, 1973, p. 24) [here, *Z. richardsoni* Subzone of *S. kirschneri* Zone]. Although the critical septal sutures of the *Dorsetensia* have not been described, at least some of the specimens appear to be true *Dorsetensia*, indicating an age not older than *O. sauzei* Chronozone (Westermann and Riccardi, 1972; Morton 1972, pers. comm.). The occurrence of late *S. kirschneri* together with *Zemistephanus* near the top of the Basey Member suggests correlation with the upper Red Glacier Formation or lowermost Fitz Creek Siltstone in southern Alaska (Imlay, 1964, pp. B28-30). The associated *Poecilomorphus* ex gr. *P. cycloides* (d'Orbigny) is of time-stratigraphic significance because the entire genus (*s.s.*) is restricted in western Europe to the lower part of the *S. humphriesianum* Zone, marking the "*P. cycloides* Subzone" of Sturani (1971) [= *Dorsetensia romani* Subzone]. Imlay omitted the comparison of his alleged new species to *P. cycloides*, with which Sturani (1971, p. 100) united the "species" of Buckman mentioned by Imlay as well as all other Buckman "species". This upper fauna containing *Zemistephanus* thus is correlated with the lower, rather than with the upper, *S. humphriesianum* Standard Zone as was suggested by Imlay (1973, p. 24).

At a number of localities in the unnamed "Middle Member" of the Snowshoe Formation in the Seneca and Emigrant Creek areas, ammonite faunas occur which are very similar in generic and even specific composition to those from the Fitz Creek Siltstone in southern Alaska (Imlay, 1973, table 9). *Stephanoceras* (s.s.), *S. (Stemmatoceras)* and *Teloceras* which are so abundant in British Columbia and in the "*Teloceras* fauna" of Alberta are absent from the associations of *Stephanoceras kirschneri* - *S. juhlei* and of "*Normannites crickmayi*" [= *Zemistephanus* ♂] - *Zemistephanus richardsoni* - "*Teloceras itinsae*" [= *Zemistephanus* ♀] - *Chondroceras*. This again suggests that these associations are older than the British Columbia and Alberta faunas mentioned.

Specimens identified with the poorly known "*Teloceras itinsae*" McLearn by Imlay (1973, p. 90; pl. 46, figs. 10, 11, 13) are too poorly preserved and fragmentary to allow confident placement in *Teloceras*; indeed, the nature of the primary ribbing and nodes is very similar to that seen on the phragmocone whorls of *Zemistephanus*. Microconchs described under "*Normannites (Itinsaites) crickmayi* (McLearn)" (Imlay, 1973, pp. 83, 84; pl. 41, figs. 2-5), all laterally crushed, have ribbing and coiling indicative of *Zemistephanus* ♂ as discussed here under the genus *Zemistephanus* McLearn, 1927 (*q.v.*). The specimens identified with "*N. orbignyi* Buckman" (Imlay, 1973, pp. 82, 83; pl. 41, figs. 9, 10, 18, 20), though apparently too fragmentary to allow detailed measurement, also closely resemble *Zemistephanus* ♂. Three specimens from a single locality (Lupher's loc. 272) were compared with "*Itinsaites*" [*Stephanoceras*] *itinsae* McLearn ♂ (Imlay, 1973, p. 48; pl. 41, figs. 6-8) but are much smaller than that species and have stronger, more rounded nodes. *Stephanoceras itinsae* has not been found elsewhere associated with the *S. kirschneri* - *S. juhlei* assemblage and *Dorsetensia* (?) *oregonensis* Imlay, *Pelekodites silviesensis* Imlay and *Sonninia* cf. *alsatica* (Haug) ["*S.* cf. *nodatipingius*"] which comprise the rest of the fauna at this locality, suggest correlation with the upper *O. sauzei* or lower *S. humphriesianum* Standard Zones. In British Columbia and Alberta, the true *Stephanoceras itinsae* is associated with *S. (Stemmatoceras)*, *Teloceras* and *C. oblatum* (Whiteaves) and is placed in the middle to late *S. humphriesianum* Chronozone.

The virtual absence in Oregon of the usually rich "*Teloceras* fauna" of Alberta and of similar associations of British Columbia, is noteworthy. The "*Stemmatoceras* aff. *S.*

albertense McLearn" of Imlay (1973, pp. 89, 90; pl. 46, figs. 1-9) is very similar to *Stephanoceras itinsae* (McLearn) in whorl shape, ornamentation and suture pattern but has more closely spaced primary ribs. Imlay compared this material with a specimen from Ribbon Creek, Alberta, described by Frebold (1957; pl. 21, fig. 2a, b; pl. 23, figs. 1b, c) as "*Stemmatoceras albertense*" and herein referred to *S. itinsae* ♀ [see discussion under *Stephanoceras (Stephanoceras) itinsae* (McLearn, 1927) ♀ & ♂]. The Oregon specimens, however, all come from a conglomeratic sandstone (Lupher's loc. 8) containing fossils that have been derived from underlying parts of the Snowshoe Formation (Imlay, 1973, pp. 29, 30) and so have little biostratigraphic significance. Fragments representing four unnamed species of *Stephanoceras* are known from localities within the undifferentiated Snowshoe Formation of the Seneca area. Of these, "*Stephanoceras* sp. C" and "*Stephanoceras* sp. D" are associated with *Spiroceras,* suggesting Upper Bajocian, and appear to be macroconchs of *Lupherites* Imlay; the similar "*Stephanoceras* sp. B", however, is said to come from a fauna (Lupher's loc. 57) indicative of the *O. sauzei* Chronozone.

WYOMING, IDAHO AND UTAH

Thick sequences of grey, shaly limestones with minor sandstone, red beds and massive limestone found along the Wyoming-Idaho border and in north-central Utah have been designated the Twin Creek Limestone. The original definition and subsequent use of this stratigraphic unit have been reviewed by Imlay (1967, p. 2; table 1) who distinguished seven members and described their ammonite and bivalve faunas.

The lowest unit, the Gypsum Spring Member, varies greatly in thickness and consists mainly of soft, brownish red to yellow siltstones and claystones which are poorly exposed, with interbedded brecciated and chert-bearing limestones (Imlay, 1967, p. 17). Thick masses of gypsum also occur in some areas. An early Bajocian age (*S. humphriesianum* Zone) was suggested for this poorly fossiliferous member because the superposed Sliderock Member contains earliest Late Bajocian ammonites.

The basal oolitic limestone of the Sliderock Member varies similarly in thickness. Most of its upper parts consist of dark, fossiliferous, bedded limestones, that have yielded an interesting ammonite assemblage at numerous localities. Stephanoceratids known from the *S. humphriesianum* Chronozone of western Canada are said to be associated with *Megasphaeroceras rotundum* Imlay and *Eocephalites primus* Imlay (Imlay, 1967, table 3). *M. rotundum* is well known from the early Late Bajocian of southern Alaska where it is associated with "*Dettermanites*" [= "*Normannites*" = *Stephanoceras* ♂] *vigorosus* (Imlay), *Leptosphinctes cliffensis* Imlay, and *L. (Prorsisphinctes) delicatus* (Imlay, 1961, 1962, 1964). A single specimen described as *Stephanoceras* cf. *skidegatense* (Whiteaves) from 33 m below the top of the Sliderock Member (Imlay, 1967, p. 89; pl. 6, fig. 10) differs markedly from that species in the Queen Charlotte Islands in both strength and density of the body chamber ornamentation and is here identified with *S. itinsae* (McLearn) ♀. This occurrence is presumably lower than that of the numerous fragments of "*Stemmatoceras* aff. *S. albertense* McLearn" found with "*Stemmatoceras*" [*Stephanoceras* s.s.] *arcicostum* Imlay and the *Megasphaeroceras* faunule in the upper 15 m of the Sliderock Member. Also associated with this fauna are four corroded fragments of "*Normannites?* cf. *N. crickmayi* (McLearn)" (Imlay, 1967; pl. 12, figs. 1-4). These specimens have, however, much coarser secondary ribbing than *Zemistephanus crickmayi* ♂ or related species. The uppermost Sliderock Member belongs in the *Megasphaeroceras rotundum* Zone (defined below) of the basal Upper Bajocian which, in the *S. subfurcatum* Standard Zone of Europe, yields a similar mixed fauna with surviving Stephanoceratinae (*cf.* Mouterde *et al.*, 1971).

ASSEMBLAGE ZONES

Table 4 shows the scheme of assemblage zones for the Lower Bajocian of western North America here proposed, and their suggested correlation with European Standard Zones. Correlation in the lowermost Bajocian is uncertain because of the recent fine subdivision of the former European "*Sonninia sowerbyi* Zone" into three Standard Zones, *i.e.*, those of *Hyperlioceras discites* (oldest), *Sonninia ovalis,* and *Witchellia laeviuscula.*

The definition and characteristic fauna, occurrence and correlation of these zones follows, beginning with the oldest:

1. — Assemblage Zone of **Docidoceras widebayense**

Definition. — The assemblage is characterized by abundant *Docidoceras (Pseudocidoceras) widebayense* Westermann and other species, together with late *Pseudolioceras*, early *Witchellia* and the *Sonninia* subgenera *Euhoploceras* and *Alaskinia;* associates are *Pseudotoites, Asthenoceras* and *Eudmetoceras.* Two assemblage subzones can be recognized within this Zone:

a. — The *Docidoceras camachoi* Subzone in the lower

		EUROPE		SOUTH ALASKA		BRITISH COLUMBIA		SW. ALBERTA	OREGON & W. INTERIOR U.S.A.	
BAJOCIAN	Upper	SUBFURCATUM		Megasphaeroceras rotundum		?		?	Megasphaeroceras rotundum	
	Lower	HUMPHRIES-IANUM	BLAGDENI	? Chondroceras oblatum		Chondroceras oblatum		Chondroceras oblatum	Chondroceras oblatum	
			"HUMPHRIES"							
			ROMANI	Stephanoceras kirschneri	Zemistephanus richardsoni	Stephanoceras kirschneri	Zemistephanus richardsoni	? with Arkelloceras	Stephanoceras kirschneri	? Zemistephanus richardsoni
		SAUZEI	HEBRIDICA		?		?			S. juhlei
				Parabigotites crassicostatus		? with S. gracilis & Euhoploceras		with Sonninia gracilis & Euhoploceras	Parabigotites crassicostatus	
		LAEVIUSCULA		?					with Witchellia Latiwitchellia Fontannesia Docidoceras Euhoploceras	
		OVALIS		Docidoceras widebayense	Witchellia sutneroides					
		DISCITES			Docidoceras camachoi					
				Eudmetoceras amplectens						
AAL.		CONCAVUM		Erycitoides howelli	Tmet. tenue / Ery. teres	Erycitoides howelli		?	with Tmetoceras, Praestriqites	

Table 4. — Proposed Bajocian assemblage zones for western North America and their correlation with the European standard zones.

part, containing *D. (Pseudocidoceras) camachoi* Westermann, *D. (P.) widebayense* Westermann and *Sonninia (Euhoploceras) bifurcata* Westermann; approximately *H. discites* Chronozone.

b. — The superjacent *Witchellia sutneroides* Subzone, with *W. sutneroides* Westermann and *D. (P.) widebayense;* approximately *S. ovalis* Chronozone.

Occurrence. — The best known assemblage and faunal sequence of this zone is at Wide Bay, Alaska Peninsula (Westermann, 1969b), here designated the type area. The southeastern Alaskan faunas include part of this assemblage but may differ in age (Imlay, 1964; Westermann, 1969b). A similar, related assemblage occurs in the middle and upper parts of the Weberg Member, Snowshoe Formation in Eastern Oregon; abundant *Sonninia (Euhoploceras), Docidoceras* and *Asthenoceras* with *Witchellia (Latiwitchellia)* and *Fontannesia* indicate time-equivalence to this zone (Imlay, 1973).

Correlation. — This zone is approximately equivalent to the *H. discites* and *S. ovalis* Chronozones.

2. — Assemblage Zone of **Parabigotites crassicostatus**

Definition. — This zone utilizes one of Imlay's (1964) suggested "guide fossils" for southern Alaska. The characteristic fauna there consists of *Parabigotites crassicostatus* Imlay ♀ and ♂, *Sonninia tuxedniensis* Imlay, *Witchellia adnata* Imlay [= *Dorsetensia*], *Stephanoceras* cf. *triptolemus* (Morris and Lycett), *Emileia constricta* Imlay ♀ and ♂ and *Sonninia (Papilliceras)* cf. *arenata* (Quenstedt) recorded at USGS Mesoz. locs. 21296, 21261, 21263 and 3009. In eastern Oregon the following fauna (at USGS Mesoz. loc. 29241, 29395 and Loc. no. 69) from the middle and upper parts of the Warm Springs Member, Snowshoe Formation also represents this Zone (Imlay, 1973): *Parabigotites crassicostatus* Imlay ♀, *"Otoites contractus"* (Sowerby) [= *Emileia constricta* Imlay ♂], *Emileia buddenhageni* Imlay, *Witchellia connata* (Buckman), *Sonninia (Papilliceras) stantoni* (Crickmay), *Stephanoceras mowichense* Im-

lay, "*Normannites* cf. *crickmayi*" (McLearn), *Dorsetensia* cf. *pulchra* Buckman and *Stephanoceras (Skirroceras) juhlei* Imlay. *Arkelloceras* Frebold also appears to belong to this Zone (see discussion under *Alberta*).

Occurrence. — As well as the type locality in the middle parts of the Red Glacier Formation, Cook Inlet, southern Alaska (Imlay, 1964) and the middle and upper parts of the Warm Springs Member, Snowshoe Formation, eastern Oregon (Imlay, 1973) occurrences of this fauna have been recorded from the upper parts of the Kialagvik Formation at Wide Bay, Alaska Peninsula (Westermann, 1969b) and from the Rock Creek Member, Fernie Group in the Snake Indian River valley, western Alberta (by the presence of *Arkelloceras;* Imlay, 1964; Westermann, 1964b).

Correlation. — The *P. crassicostatus* Zone is placed in the early *O. sauzei* Chronozone (? and latest *W. laeviuscula* Chronozone).

3. — Assemblage Zone of **Stephanoceras kirschneri**

Definition. — In southern Alaska *Stephanoceras (Skirroceras) kirschneri* Imlay first appears in the upper parts of the Red Glacier Formation (USGS Mesoz. loc. 21267), associated with *Holcophylloceras costisparsum* Imlay and *Sonninia?* sp. juv. (Imlay, 1964); it occurs alone at Mesoz. locs. 21264 and 3013. The species ranges almost to the top of the overlying Fitz Creek Siltstone (Mesoz. locs. 26599 and 21274) where it is associated with species representing the subzone of *Zemistephanus richardsoni* (discussion below). In eastern Oregon *S. kirschneri* occurs near the base of the Basey Member, Snowshoe Formation, associated with *Dorsetensia oregonensis* Imlay, *D.* cf. *subtexta* Buckman, *D.* cf. *edouardiana* (d'Orbigny), *Asthenoceras delicatum* Imlay, *Pelekodites silviesensis* Imlay and *S. (Skirroceras) juhlei* Imlay.

The Subzone of *Zemistephanus richardsoni* represents the upper part of the *S. kirschneri* Zone; *Z. richardsoni* (Whiteaves) replaces "*Teloceras itinsae*" of Imlay [= *Zemistephanus alaskensis* n. sp.] as the "guide fossil" previously suggested for this fauna by Imlay (1964). In southern Alaska *Z. richardsoni* appears at the base of the Fitz Creek Siltstone and occurs throughout this unit with the following fauna: *Z. carlottensis* (Whiteaves), *Z. alaskensis* n. sp., *Chondroceras defontii* (McLearn), *Stephanoceras (Skirroceras) kirschneri* Imlay and *S. (Stemmatoceras?)* cf. *palliseri* (McLearn). The upper parts of the Basey Member, Snowshoe Formation, eastern Oregon have yielded a fauna which includes *Poecilomorphus 'varius'* Imlay, *Chondroceras allani*, McLearn, *Stephanoceras* cf. *nodosum* (Quenstedt), '*Normannites*' *orbignyi* Buckman, "*N. crickmayi* (McLearn)", *S. (Skirroceras) kirschneri* Imlay, *Dorsetensia oregonensis* Imlay, *Pelekodites dobsonensis* Imlay and *Sphaeroceras* sp.

The Upper Red Glacier Formation and the Fitz Creek Siltstone at Cook Inlet, southern Alaska are designated the type area for both the *S. kirschneri* Zone and its *Z. richardsoni* Subzone.

Occurrence. — Outside the type area, the *S. kirschneri* Zone is represented in the Taseko Lakes area of British Columbia where *S. (Skirroceras)* cf. *kirschneri* has been recorded (Frebold *et al.*, 1969). Faunas representing the *Z. richardsoni* Subzone are more widely known, occurring at MacKenzie Bay, Queen Charlotte Islands (with *Z. richardsoni* ♀ & ♂, *Z. alaskensis* and *Chondroceras* sp.), in Manning Park, southern British Columbia (*Z. richardsoni* with "*Graphoceras crickmayi*" Frebold = *Poecilomorphus;* Frebold *et al.*, 1969) and in the unnamed Middle Member of the Snowshoe Formation in the Emigrant Creek area, eastern Oregon (*Z.?* cf. *richardsoni* with *S. kirschneri* and "*Teloceras itinsae*" McLearn = *Zemistephanus alaskensis* n. sp.; Imlay, 1973).

Correlation. — The entire *S. kirschneri* Zone is placed in the late *O. sauzei* Chronozone, *Dorsetensia hebridica* Subzone (Morton, 1975, 1976), and in the early *S. humphriesianum* Chronozone, *D. romani* Subzone. The association of *Poecilomorphus* with the faunas of the *Z. richardsoni* Subzone indicates the *D. romani* Chrono-Subzone [= "*P. cycloides* subzone"].

4. — Assemblage Zone of **Chondroceras oblatum**

Definition. — The fauna characterizing this zone is best known from the Fernie Group at Ribbon Creek, southern Alberta where *C. oblatum* (Whiteaves) occurs with *C. allani* (McLearn), *Stephanoceras itinsae* (McLearn) ♀ and ♂ [= "*Itinsaites itinsae*" (McLearn)], *S. (Stemmatoceras) dowlingi* (McLearn), and *Teloceras crickmayi* (Frebold). This is designated the type area.

Occurrence. — The so-called "*Teloceras* fauna" is known from numerous other localities within the Fernie Group of western Alberta (Frebold, 1957), in the lower part of the Yakoun Formation on South Balch Island, Queen Charlotte Islands (*S. itinsae* ♀ and ♂ with *C. oblatum* ♀ and ♂) and at the base of the Cynthia Falls Sandstone at Tuxedni Bay, southern Alaska where *C. oblatum* first appears and is associated with *C. allani* and *Stephanoceras* sp.

Correlation. — Due to the presence of *Teloceras* together with typical *Stephanoceras*, this zone is dated as middle and late *S. humphriesianum* Chronozone, approximately "*S. humphriesianum*" and *T. blagdeni* Subzones.

5. — Assemblage Zone of **Megasphaeroceras rotundum**

Definition. — From the Twist Creek Siltstone, Cook Inlet, southern Alaska comes an ammonite fauna consisting of *M. rotundum* associated with *Macrophylloceras* cf. *grossicostatum, Calliphylloceras* sp., *Lytoceras* sp., *Spiroceras?* sp., *Lissoceras bakeri, Oppelia (Liroxyites) kellumi, Sphaeroceras talkeetnanum, 'Dettermanites' vigorosus, Leptosphinctes cliffensis* and *L. (Prorsisphinctes?) delicatus,* all Imlay spp. (1962, 1964). This is designated the type area.

Occurrence. — Outside southern Alaska faunas belonging to the *M. rotundum* Zone have been recognized in the Sliderock Member of the Twin Creek Limestone, Western Interior of the United States [where *M. rotundum* occurs with *Stephanoceras* cf. *itinsae* (McLearn), *S. (Stemmatoceras) arcicostum* Imlay, *S. (Stemmatoceras)* cf. *albertense* (McLearn), "*Normannites* cf. *crickmayi*" (McLearn) and *Eocephalites primus* Imlay (Imlay, 1967)] and in an undifferentiated part of the Snowshoe Formation at Emigrant Creek, eastern Oregon. However, the latter clearly represents a mixed fauna collected from 65 to 75 m of section (Imlay, 1973, p. 51). The zone may also be represented in the Hazelton Group of British Columbia where Frebold and Tipper (1973) have tentatively identified deformed specimens as *Megasphaeroceras* ? aff. *M. rotundum.* During fieldwork in 1978 in south-western Alberta, one of us [Hall] located in the Fernie Group a fauna comprising *Spiroceras* sp., *Megasphaeroceras* ? sp. and various stephanoceratids, which is the first recorded occurrence of this zone in Fernie strata. The geographic distribution of this zone is currently being extended to the Andes of northern and central Chile and, possibly, to Mendoza province, Argentina (Westermann and Riccardi, unpublished information; their field work in September 1978 suggests, however, that *Megasphaeroceras* is congeneric with *Eurycephalites* and that the species *rotundum* may have a much extended vertical range).

Correlation. — The association of *Leptosphinctes* with late stephanoceratids in this zone parallels faunas described from the base of the European *S. subfurcatum* Standard Zone (Pavia and Sturani, 1968; Pavia, 1969). The upper boundary of this Zone in North America is unknown.

THE SUPPOSED PACIFIC FAUNAL REALM

Increasing knowledge of Jurassic ammonoid faunas has necessitated constant revision of the concept of faunal realms first discussed by Neumayr (1883) and further developed by Uhlig (1911). As pointed out by Arkell (1956) these early schemes attempted to deal with the whole of Jurassic time, leading to many discrepancies as new faunas became known. Such comprehensive treatment was impossible because, as is now clear, the differentiation of Jurassic ammonoid faunas was both spasmodic and progressive, reaching a maximum in the Upper Jurassic (Gordon, 1976). Arkell still believed that the Lower Jurassic ammonoid faunas were universal in their distribution (Arkell, 1956, p. 609) but more recent work has shown the existence of provinciality throughout most of the Lower Jurassic (Dean *et al.*, 1961; Hallam, 1971; Sapunov, 1971; Howarth, 1973).

Arkell (1956) recognised three Jurassic ammonoid Faunal Realms, the Boreal, Tethyan and Pacific, distinguished mainly on the basis of restricted occurrences of certain ammonite families. The Boreal and Tethyan Realms are clearly discerned on this basis for most of the Jurassic, though the boundary between them frequently fluctuated north and south. The causes of initiation, maintenance and periodic expansions of these two Realms have been the subject of much recent discussion (Imlay, 1965; Stevens, 1971; Stevens and Clayton, 1971; Hallam, 1969, 1971).

Development of the Pacific Realm, beginning in the Lower Bajocian with the appearance of the "peculiar Pacific genera *Pseudotoites* and *Zemistephanus*" was proposed by Arkell (1956, p. 609). These genera were supposedly restricted to Western Australia, Indonesia and the western seas of North and South America. Records of *Zemistephanus* from Western Australia are now believed to be misidentifications (see discussion under *Zemistephanus* McLearn, 1927) and it seems this genus was endemic to the North American Cordilleran geosyncline. It is definitely known only from the Queen Charlotte Islands and southern Alaska, where it occurs in beds of the uppermost *O. sauzei* or lower *S. humphriesianum* Chronozones. *Pseudotoites* is well known from the "*S. sowerbyi* Zone" of southern Alaska (Westermann, 1969b), and Chile and west-central Argentina (Westermann and Riccardi, 1972). The transfer of *Zemistephanus carlottensis* (Whiteaves) to *Pseudotoites* by Arkell (1954) has been shown to be unwarranted (Imlay, 1964; Westermann, 1964a). *Pseudotoites* is also known from Western Australia (Arkell, 1954) and some islands of the Indonesian Archipelago (Westermann and Getty, 1970). On the other hand, the distribution of the overwhelming number of Lower Bajocian genera is cosmopolitan: *Emileia* (including *Otoites*), *Sonninia* and *Stephanoceras (Skirroceras)* in the *O. sauzei* Chronozone; *Stephanoceras* s.s. (including *Normannites*), *Teloceras* and *Chondroceras* in the *S. humphriesianum* Chronozone. Other genera do seem to be endemic to the western Cordilleran seas of North America, e.g., *Parabigotites* and *Zemistephanus.* The restricted occurrence of

these few genera in faunas otherwise composed of predominantly cosmopolitan families indicates some endemism but does not justify separate Realm status.

Also, *Pseudotoites* and *Zemistephanus* are of different age, further diminishing the number of supposedly "unique" forms defining the Pacific Realm at any one time. Different parts of the supposed Pacific Realm were inhabited at various times by ammonite genera that migrated from the Boreal, eastern Tethyan and western Tethyan Realms. While it is clear that some genera were able to migrate across the Pacific basin (Westermann and Riccardi, 1976) others retained a restricted distribution (Westermann, 1969b; Khudoley, 1974).

Of greater significance may be the geographic distribution of *Arkelloceras* which is known in relative abundance from strata of uncertain age in the North West Territory and Arctic Canada (Frebold, 1961; Frebold *et al.*, 1967), as a specimen from the *O. sauzei* Chronozone of northern Alaska (Imlay, 1976), and also from southern Alaska (Imlay, 1964) and western Alberta (Westermann, 1964b). The dominance of this genus in the Arctic may represent the initiation of the more strongly defined Boreal faunas with *Cranocephalites* in the Upper Bajocian and Lower Bathonian (Frebold, 1961, p. 36).

SYSTEMATIC PALEONTOLOGY

MEASUREMENTS

All measurements of specimens are given in mm. Where possible whorl dimensions were measured on cut and polished cross-sections or on broken specimens obtained during dissection. Most measurements were made on internal molds. The following abbreviations are used throughout the text and in text-figures:

- D = shell diameter.
- W = maximum whorl width measured between ribs or tubercles.
- H = height of the whorl measured from the umbilical seam to the venter.
- U = diameter of the umbilicus measured between the umbilical seams.
- P = Number of primary ribs per half-whorl, apicad from the stated shell (umbilical) diameter.
- S = Number of secondary ribs per half-whorl, counted as for primary ribs.
- PL = length of the primary ribs measured from the umbilical seam to the centre of the tubercle or point of furcation.

The graphs represent "mass curves", usually with more than one measurement taken from each specimen; points measured on phragmocone whorls are represented by open symbols, those from the body chamber are solid. Macroconchs (♀), microconchs (♂) and specimens from different localities are indicated by the use of symbols that are explained on each graph. Points joined by thin, continuous lines represent measurements made from various growth stages on the same specimen ("individual growth curves"); heavier continuous lines joining numbered points represent measurements from holotypes; other dashed lines joining numbered points represent measurements on "species" considered synonymous with the named species. Regression lines, based only on measurements from the phragmocone whorls, are represented by the heaviest continuous lines and, where relevant, the suggested sexes are indicated by the appropriate symbols. Approximate "growth lines" for each dimorph appear in small insets.

On graphs showing rib counts, individual points from different growth stages of each specimen are joined to produce a "growth curve" emphasising individual ontogenetic variation. The use of umbilical diameter allows data to be incorporated from the inner whorls of undissected specimens and figured specimens, at least for primary ribbing exposed on the lower flanks or umbilical walls.

The diameter at which individual sutures were drawn is indicated to the right of each figured suture. Shell diameters at which cross-sections were drawn are indicated on each diagram. Sections through body chamber whorls are shaded; black areas represent ribs and tubercles.

For the septal suture, the conventional symbols of the (mostly European) recent literature are used. These are, from venter to dorsum: E, external lobe; L, lateral lobe; U_1, U_2 etc., umbilical lobes in order of appearance; U_n, secondary internal lateral lobe; and I, internal lobe.

DIMORPHISM AND NOMENCLATURE

Each of the three Lower Bajocian faunas described from the Yakoun Formation on the Queen Charlotte Islands is of very low diversity, usually with only one stephanoceratid and one sphaeroceratid dimorphic species at each locality (Table 2). Thus difficulties due to possible overlap of morphological features between closely related species are minimized, especially among the less variable microconchs, allowing dimorphs to be paired at the species level. Corresponding dimorphs are regarded as the sexes of a single ammonite species and given the same name; the probable sex of each dimorph is indicated by the use of the appro-

priate biological symbols, though this is recognised as being a convention (Palframan, 1969, p. 148). In this usage we follow the procedures already proposed by Palframan (1969), Westermann (1969a), Makowski (1962) and Guex (1967), among others. Where one partner of a dimorphic pair has not been previously named, no new name is given and the two dimorphs are regarded as being conspecific. In those cases where both dimorphs have previously been described and named, rules of priority must be followed when choosing a single name for the species.

Full morphological descriptions and synonymies of both dimorphs are given separately, then comparison of their juvenile morphologies is made. The holotype of a dimorphic species must be that of the dimorph having priority (some will be macroconchs, others microconchs within a single genus); a specimen representing the other dimorph is designated an allotype.

There is a need for further study of data selection in attempts to demonstrate identical growth gradients for supposed dimorphs by means of formal statistical tests, *i.e.* whether one should use regression statistics based on (a) 'mass' curves in which numerous measurements are obtained from each individual at different growth stages; (b) 'mass' curves in which only one measurement is obtained from each specimen; or (c) individual growth curves using many measurements taken throughout the ontogeny of a single specimen representing each dimorph.

The biological meaning of the intercept ('b' in the allometric equation $Y = bX^a$) is uncertain (White and Gould, 1965). The inequality of 'b' found in most growth patterns for dimorphs tested here may represent some real difference in shape of the embryonic growth stages of males and females; on the other hand, dimensions and growth patterns on the first (nepionic) whorl of ammonites are known to diverge significantly from those found throughout the remainder of the phragmocone and so differences in 'b' may have no biological significance.

Order AMMONOIDEA Hyatt, 1889

Suborder AMMONITINA Hyatt, 1889

Superfamily **STEPHANOCERATACEAE** Neumayr, 1875

Family **STEPHANOCERATIDAE** Neumayr, 1875

Subfamily **STEPHANOCERATINAE** Neumayr, 1875

Genus **ZEMISTEPHANUS** McLearn, 1927

Type species. — *Ammonites richardsoni* Whiteaves, 1876 (by original designation).

Discussion. — The type-species, originally described by Whiteaves (1876, pp. 32, 33; pl. 5, figs. 1, 2), was based on a single specimen from the collection of fossils made by J. Richardson in 1872 from the shores of Skidegate Inlet, Queen Charlotte Islands. No further information on its precise locality was given. McLearn (1929, pp. 18-21) described an additional specimen as *Zemistephanus richardsoni* (Whiteaves) and two other specimens (designated *Z. vancouveri* McLearn, 1929 and *Z. funteri* McLearn, 1929) from the lower part of the Yakoun Formation on the shore of MacKenzie Bay, Maude Island, Skidegate Inlet (Text-fig. 1). More recent collections by Sutherland Brown (1968) and one of us [Hall] have produced a number of macroconch *Zemistephanus* only from this locality, so it seems certain that the holotype of *Z. richardsoni* came from here. *Zemistephanus* appears to be endemic to western North America (with the possible exception of "*Coeloceras*" *indicum* Kruizinga from the Sula Islands, Indonesia) and is of restricted stratigraphic range and diversity; yet the identity and affinities of the genus have been subject to widely varying treatment by other authors.

Another specimen from Skidegate Inlet recently recognized as belonging to this genus (Imlay, 1964) was described by Whiteaves (1876, pp. 38, 39) as *Ammonites carlottensis.* However, this species was not included by McLearn (1927, 1929) in his original discussion of the genus *Zemistephanus* and has been variously placed in *Perisphinctes, Stephanoceras* and *Pseudotoites.* Arkell (1954) believed there was a strong resemblance between *A. carlottensis* and the Western Australian *Pseudotoites leicharti* (Neumayr) and so transferred the poorly known Canadian species to that genus. Later work on material from southern Alaska, in which details of the suture were recognized for the first time, resulted in establishment of *A. carlottensis* Whiteaves as a *Zemistephanus* (Imlay, 1964). This was based on the strong similarities of the body chamber, ornamentation and suture with the type-species. Similar characters are used here in transferring the Alaskan material described by Imlay (1964, pp. B50, 51) as '*Teloceras itinsae* McLearn' to *Zemistephanus,* as *Z. alaskensis* n. sp. The aplanulate structure of the septum indicates that *Zemistephanus* should be affiliated with the family Stephanoceratidae rather than the Otoitidae which exhibit an abullate septum (Westermann, 1964a): E/L and I/U_n are much larger than the adjacent saddle elements in the suture.

The two additional species erected by McLearn, *Z. vancouveri* and *Z. funteri,* were based on single, incomplete and poorly preserved specimens. The holotype of *Z. funteri* is badly weathered and one side is missing; it is here regarded as a slightly smaller form of *Z. richardsoni.* As previously

suggested by Imlay (1964, p. B53), the incomplete holotype of *Z. vancouveri* appears to be identical with the Alaskan specimens of *Z. carlottensis* (Whiteaves).

Two alleged Australian species described by Arkell (1954), '*Z.*' *corona* and '*Z.*' *armatus*, have been shown to be coronate developments of *Pseudotoites* (Westermann, 1964a, p. 62). From mainland western Canada several specimens have previously been placed in *Zemistephanus*. '*Z. crickmayi*' Frebold (1957, pp. 52, 53) from the Rock Creek Member of the Fernie Group at Ribbon Creek in southern Alberta is a *Teloceras* [see discussion under *Teloceras crickmayi* (Frebold, 1957) ♀]. '*Zemistephanus* sp.' of Frebold and Tipper (1973, p. 1123), from Tenas Creek in north-central British Columbia, represented by a poorly preserved outer whorl with unknown coiling and cross-section, is of uncertain generic affinity.

A single impression of a stephanoceratid ammonite from the Lookout Section in Manning Park, southern British Columbia, was identified as *Z. richardsoni* (Whiteaves) by Frebold (in Frebold *et al.*, 1969, pp. 25, 26; pl. II, fig. 1; pl. IV, fig. 1). The specimen appears to be complete and not significantly larger than other described specimens from Alaska and the Queen Charlotte Islands. However, the occurrence of fine riblets or striae in place of relatively coarse secondary ribs on all visible parts of the flanks and venter and the loss of the large, conical nodes on the body chamber are not characteristic of *Z. richardsoni* (Whiteaves). Even the generic identity remains uncertain because of the unknown whorl shape.

'*Teloceras*' *warreni* McLearn is a large species of *Zemistephanus*, characterized by a narrow umbilicus, steep umbilical wall and large nodes situated low on the flanks.

Whiteaves (1876, p. 33) commented on the close relationship of *Ammonites richardsoni* to *A.* [*Erymnoceras*] *coronatus* Bruguière (1789) and *A.* [*Teloceras*] *Blagdeni* Sowerby (1818). Since then several authors (Warren, 1947, p. 72; Frebold, 1957, p. 53, Arkell, 1954, p. 579; Westermann, 1964a, pp. 62, 68) have questioned the genus-level distinction of *Zemistephanus* and *Teloceras*, though McLearn (1929) in his original description of the type species of *Zemistephanus* had already noted two important distinctions: the more dorsal position of the tubercles and the change on the body chamber to more serpenticone coiling. Imlay (1964, p. B52) distinguished the two genera by characters of the body chamber:

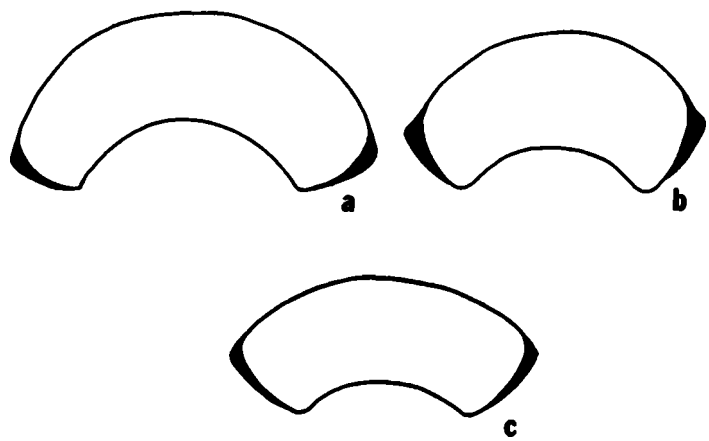
Zemistephanus is characterized by

> rather marked uncoiling of the body chamber, the low position of the tubercles on the flanks of the body chamber, a tendency for the tubercles to weaken near the aperture of the large, adult specimens.

Teloceras

> has a coronate adult body chamber, its tubercles occur higher on the flanks and remain strong on the body chamber, and the adult whorl contracts little or none at all from the preceding whorl.

Several specimens of *Teloceras* cf. *blagdeni* (Sowerby) from Goslar in northwest Germany have been compared with the inner whorls of *Z. richardsoni* from the Queen Charlotte Islands and southern Alaska. Clear differences are already apparent at 25 mm diameter. In *Zemistephanus*, the whorls are broader and more strongly arched than in *Teloceras*; the inner flanks are more strongly convex and grade into the almost vertical umbilical slope, forming a deep, crater-like umbilicus, and the nodes are more blunt and much lower on the flanks (Text-fig. 3). On the inner whorls,



Text-figure 3. — Cross-sections of phragmocone whorls of *Zemistephanus* and *Teloceras*. a-b, *Z. richardsoni* ♀ (McM J1797a at D = 70 mm and USNM 3000.1 at D = 62 mm), with whorl overlap of 40% and 33%, respectively; c, *T.* cf. *blagdeni* ♀ (McM J1875a from Goslar, Germany, at D = 63 mm), with whorl overlap of 25%. × 0.9.

the primary ribs in *T.* cf. *blagdeni* are more prominent than the low and more distant ribs on *Zemistephanus*. The nodes are not as large and conical as in *Zemistephanus* but instead are sharp terminations of the primaries. On later growth stages of typical *Teloceras* the primary ribs remain broader and prominent, the nodes become large and rounded and secondary ribs are thick and relatively sharp; the coronate cross-section becomes more pronounced, with a broad, flat venter. The end of the adult *Zemistephanus* phragmocone, in contrast, has weaker ornamentation, the primary ribs becoming broad undulations and the nodes blunt and rounded.

Species of *Stephanoceras (Stemmatoceras)*, that more closely resemble *Zemistephanus* in adult whorl section and coiling than does *Teloceras*, are all distinguished from *Zemistephanus* by the elliptical rather than ovate inner whorls

and the more prominent primary ribs (at least on the juvenile whorls).

Some species exhibit an exceptionally close homeomorphy to Australian and South American representatives of *Pseudotoites* Spath, in particular the relatively evolute and coarsely ribbed *Z. carlottensis* (Whiteaves) (*cf.* Arkell, 1954, pls. 35, 36, 39, 40; Imlay, 1964, pls. 25, 27). Besides its earlier age (*S. ovalis* vs. *S. humphriesianum* Zone), *Pseudotoites* differs in the denser, thinner ribbing of the juvenile whorls and, particularly, in the bullate, not planulate, septum (subequal first and second, internal and external lateral saddles in the suture). The two genera belong indeed to separate families (Westermann, 1964a).

Zemistephanus ♀ (macroconch)

Diagnosis. — Inner whorls cadiconic with broadly arched venter, well-defined lateral shoulder with large, conical nodes and steep, convex flanks forming a deep, narrow umbilicus. Primary ribs broad, faint on lower flanks, with 3-6 secondary ribs to each primary arching forward on the venter. Body chamber egresses strongly with marked decrease in whorl width and rounding of the whorl section, usually three-quarters to one whorl in length. Aperture simple with flared collar.

The following species are here included:

Zemistephanus richardsoni (Whiteaves, 1876) ♀ (Pl. 3, figs. 2a, b), lower part of the Yakoun Formation, Queen Charlotte Islands and Fitz Creek Siltstone, southern Alaska.

Zemistephanus carlottensis (Whiteaves, 1876) ♀ (Pl. 3, figs. 1a, b), lower part of the Yakoun Formation, Queen Charlotte Islands; Fitz Creek Siltstone and Cynthia Falls Sandstone, southern Alaska.

Zemistephanus alaskensis n. sp. ♀, lower part of the Yakoun Formation, Queen Charlotte Islands and Fitz Creek Siltstone, southern Alaska.

Zemistephanus warreni (McLearn, 1930) ♀, Fernie Group, Porcupine Creek, Kananaskis Valley, southern Alberta.

(??) Zemistephanus ursinum (Imlay, 1964), Fitz Creek Siltstone, southern Alaska.

Zemistephanus ♂ (microconch) ['**Kanastephanus**']

The microconch dimorph of the type species *Z. richardsoni* (from MacKenzie Bay) is described here for the first time. In all essential features it closely resembles those specimens from the same locality described by McLearn (1927) under *Kanastephanus*. Arkell (1957) placed *Kanastephanus* in synonymy with *Normannites* Munier-Chalmas, but *Kanastephanus* differs from that genus in having a deeper and narrower umbilicus, more coronate whorl section, large conical tubercles and a higher ratio of secondary to primary ribs on the phragmocone. Westermann (1964a, p. 68) placed *Kanastephanus* in synonymy with *Itinsaites* McLearn, 1927; however, '*I.*' *itinsae* is here shown to be the microconch of *Stephanoceras yakounense* McLearn (see discussion under *Stephanoceras (Stephanoceras) itinsae* (McLearn, 1927) ♀ & ♂). The "microconch genus" *Kanastephanus* McLearn, 1927 is here placed in synonymy with the corresponding "macroconch genus" *Zemistephanus* McLearn, 1927.

Diagnosis. — Phragmocone cadiconic, whorl section strongly depressed with broadly arched venter, lateral shoulder and steep, convex flanks forming a deep, narrow umbilicus. Primary ribs broad, faint on lower flanks, with large, conical nodes along the lateral shoulder and three to six secondary ribs to each primary. Body chamber egresses and contracts, aperture with ventro-lateral lappets. Ornamentation strong to aperture but nodes lost and primary ribs bifurcate.

The following species are included:

Zemistephanus richardsoni (Whiteaves, 1876) ♂, lower part of the Yakoun Formation, Queen Charlotte Islands and Fitz Creek Siltstone (possibly Cynthia Falls Sandstone also), south Alaska.

Zemistephanus crickmayi (McLearn, 1927) ♂, lower part of the Yakoun Formation, Queen Charlotte Islands; Fitz Creek Siltstone and Cynthia Falls Sandstone, south Alaska.

Zemistephanus richardsoni (Whiteaves, 1876) ♀ & ♂

Plate 1; Plate 2, figures 1-4; Plate 3, figure 2a b; Text-figures 4-7

Zemistephanus richardsoni ♀ (macroconch)

1876. *Ammonites richardsoni* Whiteaves, pp. 32, 33; pl. 5, figs. 1, 2.
1927. *Zemistephanus richardsoni* (Whiteaves); McLearn, p. 63.
1929. *Zemistephanus richardsoni* (Whiteaves); McLearn, p. 19; pl. 9, figs. 1, 2; pl. 10, fig. 2.
1929. *Zemistephanus funteri* McLearn, p. 20; pl. 10, fig. 1.
1964. *Zemistephanus richardsoni* (Whiteaves); Imlay, p. B51; pl. 25, figs. 6, 7; pl. 26, figs. 1-7.
?1969. *Zemistephanus richardsoni* (Whiteaves); Frebold *et al.*, pp. 25, 26; pl. 2, fig. 1; pl. 4, fig. 1.

Holotype. — GSC 5013, collected by J. Richardson in 1872, presumably from the lower part of the Yakoun Formation at MacKenzie Bay on the north shore of Maude Island, Skidegate Inlet, Queen Charlotte Islands (Text-fig. 1).

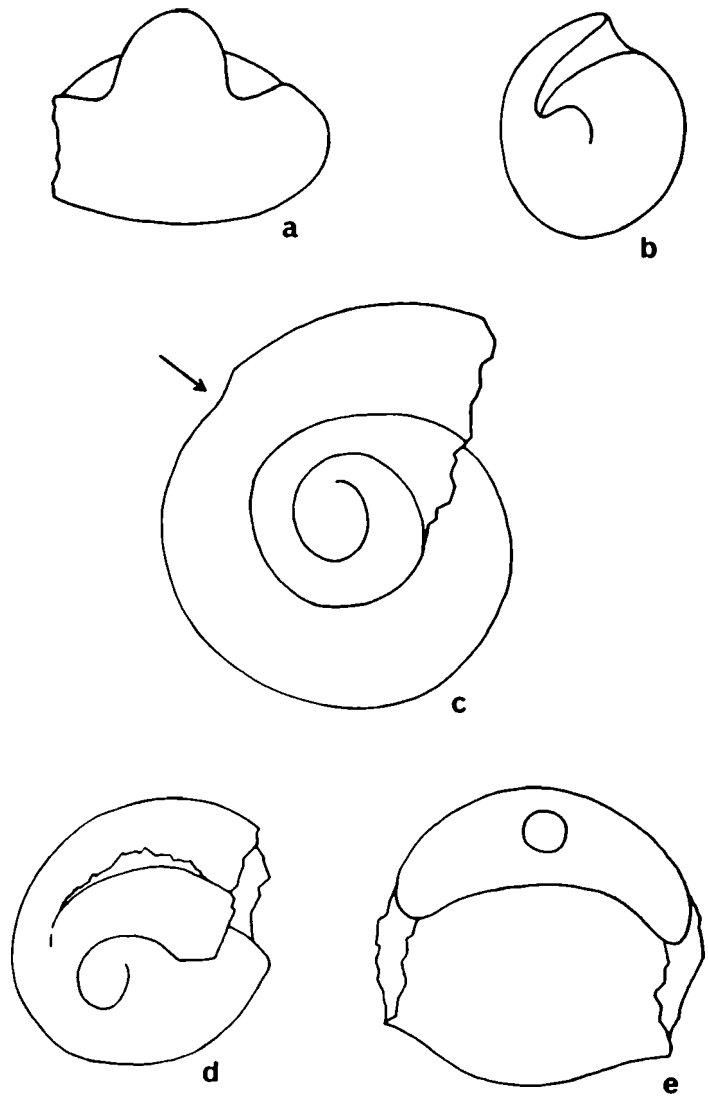
Material. — Two complete specimens (McM J1797a, b), five reasonably complete specimens (McM J1797c-g) from Yakoun Formation, lower 5 m of the section exposed at MacKenzie Bay; one specimen (McM J1797j) from shales 45 m above base of exposed section at MacKenzie Bay.

Seven specimens from the Fitz Creek Siltstone, southern Alaska (USGS Mesozoic locs. 2999, 3000, 10515, 26599). The holotype (GSC 5013), McLearn's "plesiotype" (GSC 9006) and another specimen from this locality (GSC 13639) were reexamined.

Description. — Protoconchs were obtained from one Alaskan specimen (USGS Mesozoic loc. 2999) and two juvenile specimens from MacKenzie Bay (McM J1797h, i). The protoconch is smooth and transversely elongate with a width of about 0.5 mm and a height of 0.35 mm. The first whorl is smooth with a broad, flattened venter that curves abruptly near the umbilical seam to form a short, convex flank. One whorl after the prosuture, at a diameter of 0.8 mm, there is a broad, faint constriction on the venter which fades approaching the umbilical seam. The siphuncle at this stage has a relatively large diameter and is central (Text-fig. 4).

Immediately following the constriction a distinct lateral shoulder forms with steep, convex flanks falling to the umbilical seam. A change in shell dimensions occurs at the end of the first whorl with positive allometry for whorl height; growth ratios remain constant from this point to the end of the phragmocone. At a diameter of 1.5 mm small, elongated tubercles appear along the lateral shoulder and extend onto the upper flanks; these are strongly prorsiradiate. At a diameter of 4 mm faint secondary ribs appear; they are broad and rounded, usually two to each primary. At this stage the tubercles are prominent and conical and the primary ribs are broad, curved undulations of the flanks that do not reach the umbilical seam. Whorl cross-section is depressed and coronate with H/W = 0.50 to 0.60. The venter is broad and only slightly arched while the flanks become very steep and almost vertical near the umbilical seam. Secondary ribs soon become sharper than the primaries and more densely spaced with increasing diameter, as many as five per primary. Primary ribs become rectiradiate but are still massive and rounded with large, conical nodes; their spacing during ontogeny at first decreases from 8-12 per half-whorl to a minimum of six to eight between umbilical diameters of 5 and 15 mm and then increases again to eight to ten on the final whorl (Text-fig. 5). Fine striae, which may entirely mask the secondary ribbing, appear on the outer shell surface of some specimens at diameters between 50 and 75 mm.

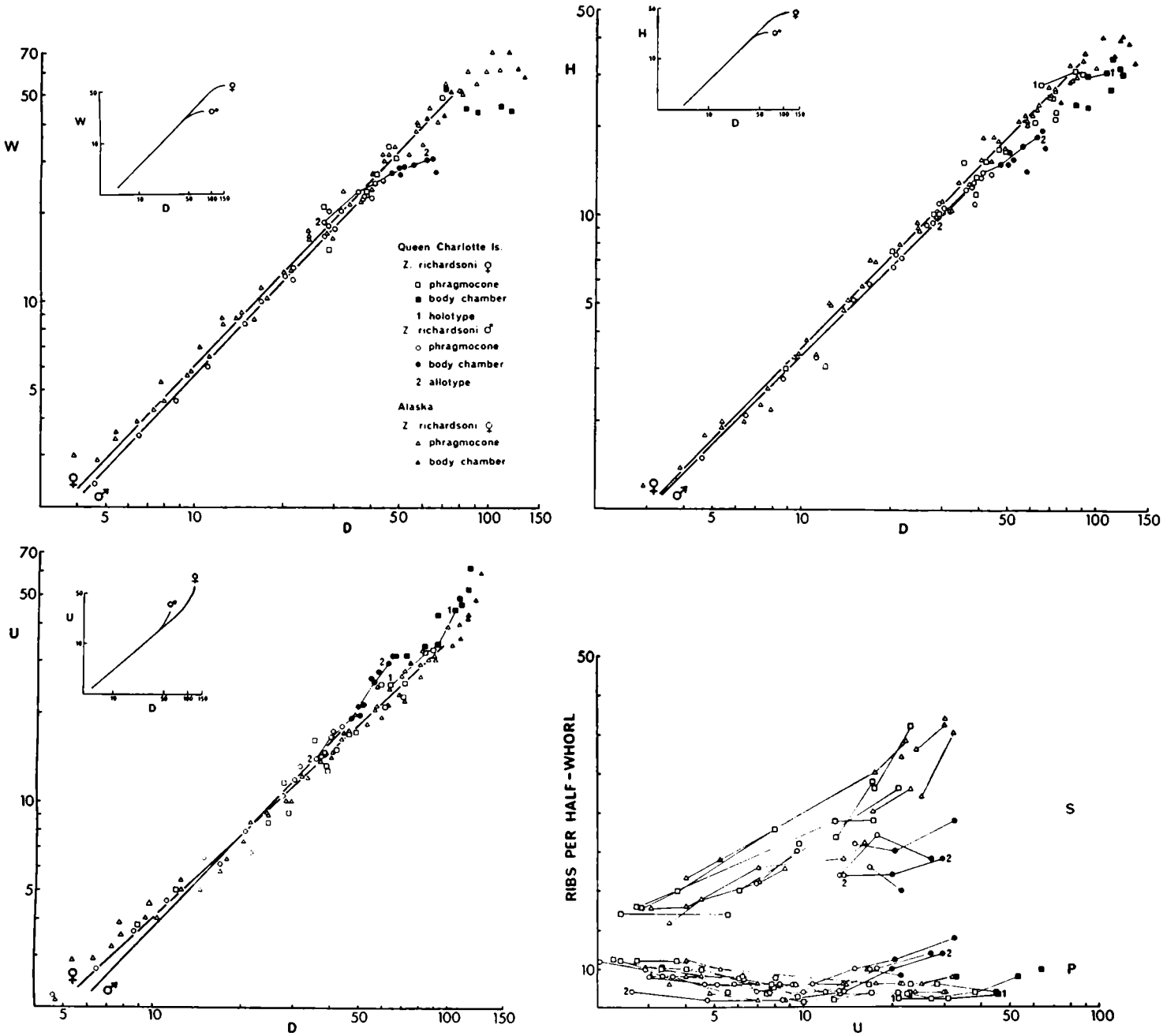
Marked changes in growth occur on the body chamber, which is about one whorl long. The umbilical seam egresses suddenly from the line of nodes on the previous whorl, the flanks become less steep and the relative (and sometimes absolute) whorl width decreases sharply, resulting in strong



Text-figure 4. — Protoconch and first whorl of *Zemistephanus richardsoni* ♀, × 62. a-b, ventral and lateral views of protoconch (McM J1797i); c, first whorl with nepionic constriction (arrow) at D = 0.8 mm (McM J1797h); d-e, lateral and ventral views of protoconch with part of first whorl, same specimen.

contraction of the whorl (Text-figs. 5, 7). H/W ratios increase to as much as 0.75 with rounding of the cross-section. Both primary and secondary ribs become faint and may disappear entirely. Nodes are blunt and rounded but persist to the aperture, situated low on the whorl. The aperture is marked by a slight constriction followed by an expanded collar and complete lip.

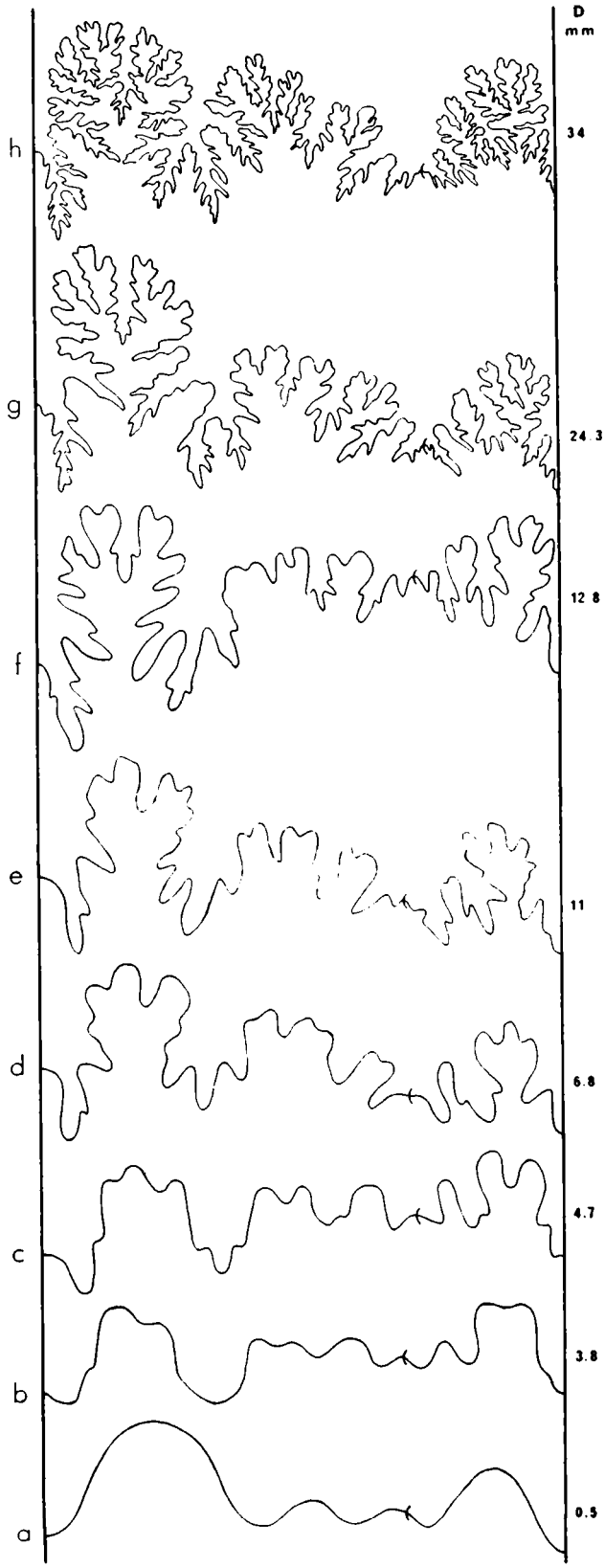
Remarks. — The specimens from southern Alaska and MacKenzie Bay agree closely in whorl shape, coiling and ornamentation (Text-figs. 5, 7). Most Alaskan representatives of the species attain larger sizes and have broader body



Text-figure 5.—Bivariate plots of whorl width (W), whorl height (H), umbilical diameter (U), shell diameter (D) and number of primary (P) and secondary (S) ribs per half-whorl, for *Zemistephanus richardsoni* ♀ and ♂ from Alaska and Queen Charlotte Islands. Based on 16 macroconchs and five microconchs, with some individual "growth lines" indicated. Insets are same graphs showing approximate positions of best-fit mass curves. See *Measurements* for additional explanations.

chambers and coarser secondary ribs. The character of the nodes on the body chamber also varies, some becoming blunt and rounded while others remain high and fairly sharp.

This species differs from *Z. carlottensis* (Whiteaves) in having more coronate whorls, losing strong secondary ribbing on the body chamber but retaining the large, conical nodes. *Z. alaskensis* n. sp. ♀ has similar phragmocone whorls but a higher body chamber whorl (H/W = 0.80 - 0.95) retaining strong secondary ribs with the nodes higher on the flanks.



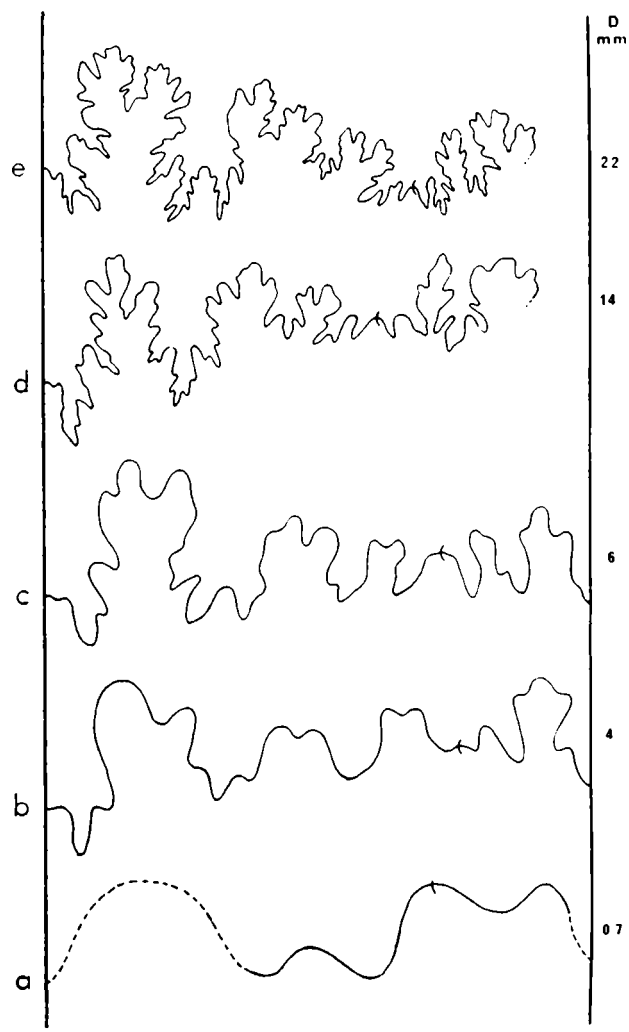
Text-figure 6. — Sutural ontogenies for *Zemistephanus richardsoni* ♀, left (a-c, USGS Mesozoic loc. 2999.3, d-h, USGS Mesozoic loc. 26599.1) and ♂, right (McM J1796d). The position of the umbilical seam is marked by (. See *Measurements* for additional explanations.

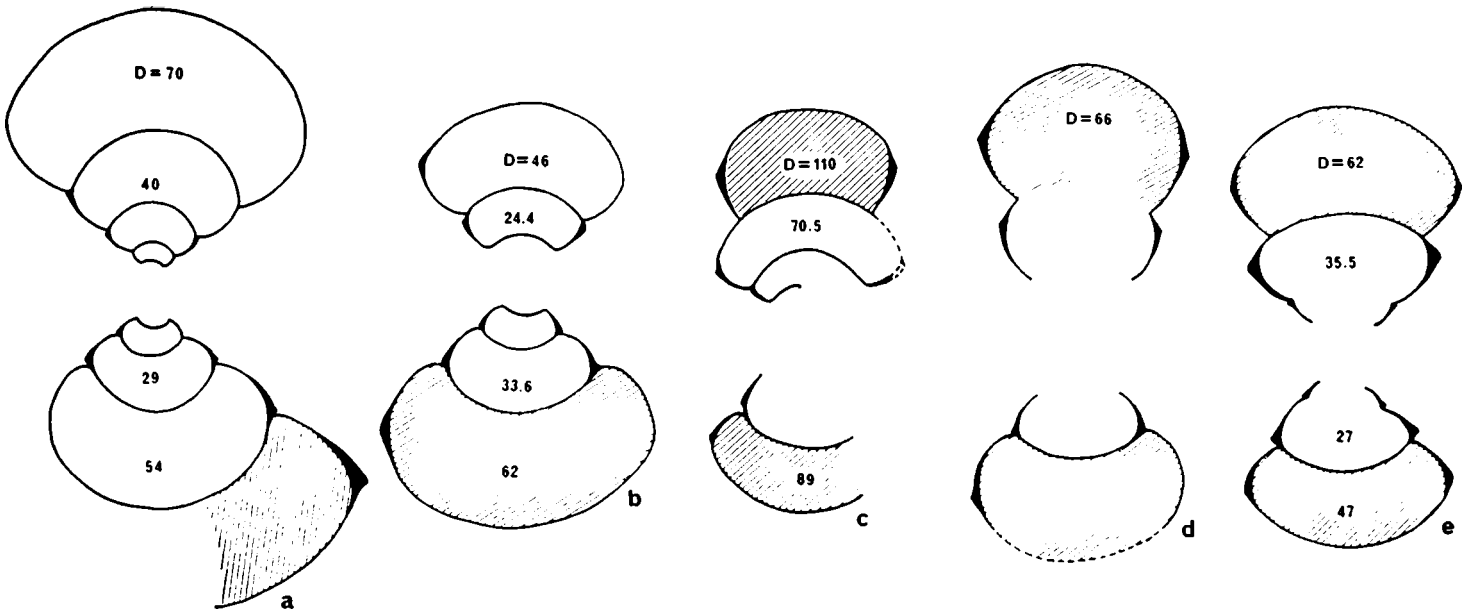
Zemistephanus richardsoni ♂ (microconch)

(This dimorph has not been described previously.)

Allotype. — McM J1796a, complete with lappets, from Yakoun Formation, lower 5 m exposed at MacKenzie Bay, Queen Charlotte Islands.

Material. — Four specimens (McM J1796a-d), three complete with aperture, from lower 5 m of Yakoun Formation exposed at MacKenzie Bay, Queen Charlotte Islands; GSC 48593 from the same locality; two crushed specimens from the Fitz Creek Siltstone in southern Alaska, USGS Mesozoic locs. 2999, 26599.





Text-figure 7. — Cross-sections of *Zemistephanus richardsoni* ♀ and ♂; × 0.9 except c = × 0.5; body chamber shaded. a-c, macroconchs (USGS Mesozoic locs. 3000.1, 26599.1, McM J1797a); d-e, microconchs (McM J1796b, McM J1796a). See *Measurements* for additional explanations.

Description. — The nature of the protoconch and earliest whorls is unknown. At a diameter of 3 - 4 mm small, sharp tubercles appear along the lateral shoulder, and extend onto the upper flanks as small ridges. The venter is smooth, broad and only slightly arched while the flanks are steep, falling straight to the umbilical seam. As diameter increases the tubercles become sharper and conical and the primary ribs extend further towards the umbilical seam; there are nine to ten per half-whorl. Broad, faint secondary ribs appear at a diameter of about 10 mm, two to each primary. At this stage the primary ribs take on characteristics that are retained throughout the phragmocone: they are broad, rounded undulations of the flanks, fading near the umbilical seam, rectiradiate and terminating in high and sharp conical nodes on the lateral shoulder. Their number decreases to six to eight per half-whorl between umbilical diameters of 5 and 15 mm, but later increases to 10 to 12 near the end of the phragmocone. Secondary ribs become strong, curving forward from the nodes, then crossing straight over the venter. Density of secondaries increases to three per primary (Text-fig. 5).

Whorl cross-section throughout the phragmocone is coronate with steep, convex flanks forming a deep, crater-like umbilicus. Nodes are situated at the point of maximum whorl width on the abrupt lateral shoulder which is at 40 to 50 percent of the total whorl height. H/W ratios throughout the phragmocone are 0.50 - 0.60.

The body chamber occupies about three-quarters of a whorl with egression from the line of nodes on the previous whorl commencing just after the last septum. While remaining broad and depressed in section, relative height increases on the body chamber, H/W ratios increasing to 0.62 - 0.64. Ribbing remains strong to the aperture with only two secondaries to each primary and the loss of the high, conical nodes. Most primary ribs bifurcate simply at the point of maximum whorl width. Flanks are less steep than on the phragmocone whorls. Lappets are ventro-lateral, short and spatulate with prominent growth lines (Pl. 2, figs. 1a, 2a); there is no strong constriction preceding the aperture and only weak flaring of the flanks.

Remarks. — *Z. richardsoni* ♂ is distinguished from the four "species" of '*Kanastephanus*' described from this locality (McLearn, 1929) by having a broader living chamber (H/W = 0.62 - 0.64 vs. 0.68 - 0.74) and less arched venter. It also has more prominent conical nodes, broad rectiradiate primary ribs and steep, convex flanks; it resembles '*Kanastephanus*' spp. in the loss of nodes on the body chamber and decline of secondary rib density to two per primary.

Dimorphism. — Comparison of whorl dimensions and ribbing pattern on the phragmocone whorls shows these specimens to be identical with the inner whorls of *Z. richardsoni* ♀ (Text-figs. 5, 7; compare Pl. 1, figs. 3c, 4 with Pl. 2, fig. 3; Appendix 1) which occurs in the same beds both at MacKenzie Bay and in southern Alaska. Both show

steep flanks with massive primary ribs and large, conical nodes, broad venter and similar changes in ribbing density at the same growth stages during ontogeny. Ornamentation of the body chamber, however, is quite different: while the microconch retains strong, bifurcating ribs but loses the sharp, conical nodes (Pl. 2, figs. 1a, 2a, 4) the macroconch has an almost smooth body chamber except for the large, rounded nodes low on the flanks (Pl. 1, figs. 1, 2). Adult macroconchs are about twice the size of microconchs and in the MacKenzie Bay strata outnumber them approximately four to one.

Zemistephanus crickmayi (McLearn, 1927) ♂

[? microconch to *Z. carlottensis* (Whiteaves, 1876) ♀]

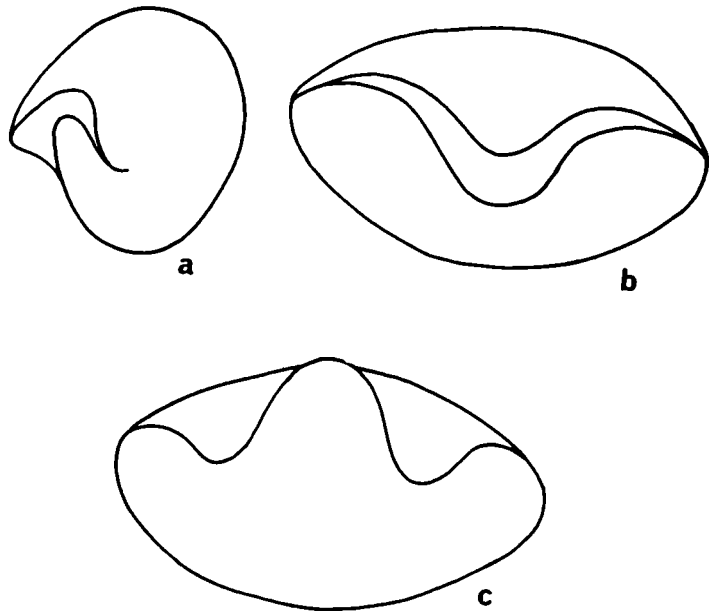
Plate 2, figures 5-8; Plate 5, figure 2; Text-figures 8-10, 11a-c

1927. *Kanastephanus crickmayi* McLearn, p. 73; pl. I, figs. 5, 6.
1929. *Kanastephanus crickmayi* McLearn; McLearn, pp. 23, 24; pl. XVI, figs. 7, 8.
1929. *Kanastephanus canadensis* McLearn, p. 25; pl. XV, figs. 4, 5.
1929. *Kanastephanus mackenzii* McLearn, p. 23; pl. XVI, figs. 1-3.
1929. *Kanastephanus altus* McLearn, p. 24; pl. XVI, figs. 4-6.
1949. *Normannites (Kanastephanus) crickmayi* McLearn; McLearn, pp. 13, 16.
1954. *Itinsaites crickmayi* (McLearn); Westermann, pp. 290-292; figs. 122, 123; pl. 27, fig. 3.
1964. *Normannites (Itinsaites) crickmayi* (McLearn); Imlay, pp. B43, 44; pl. 14, figs. 3-8, 13.
1964. *Normannites (Itinsaites) itinsae* (McLearn); Imlay, p. B44; pl. 14, figs. 1, 2.

Holotype. — GSC 9016, from the lower part of the Yakoun Formation on the north side of Maude Island, Queen Charlotte Islands, 10-22 ft. (3-7 m) above the base of the section exposed there (McLearn, 1929, p. 23).

Material. — The holotypes of McLearn's four "species" of *Kanastephanus* (GSC 9016, 9017, 9018, 9019) have been reexamined; one additional complete specimen (McM J1798a) and a number of fragments (McM J1798b-h) were collected from the type locality at MacKenzie Bay. Six specimens from the Fitz Creek Siltstone in southern Alaska (USGS Mesozoic locs. 2999, 3000, 19997, 21276).

Description. — The protoconch is smooth and elongated transverse to the plane of coiling; its width is about twice the diameter (Text-fig.8). Though the nepionic constriction was not observed, after approximately one whorl (at D = 0.85 mm) there is a sudden decrease in relative whorl width and an increase in umbilical diameter. After this point growth ratios remain constant throughout the phragmocone (Text-fig. 9). On the first whorl the shell is smooth and globose with broadly arched venter and strongly convex flanks falling steeply towards the umbilical seam. The siphuncle is central. At D = 3.0 mm small, elongated tubercles appear along the sharp lateral shoulder which is

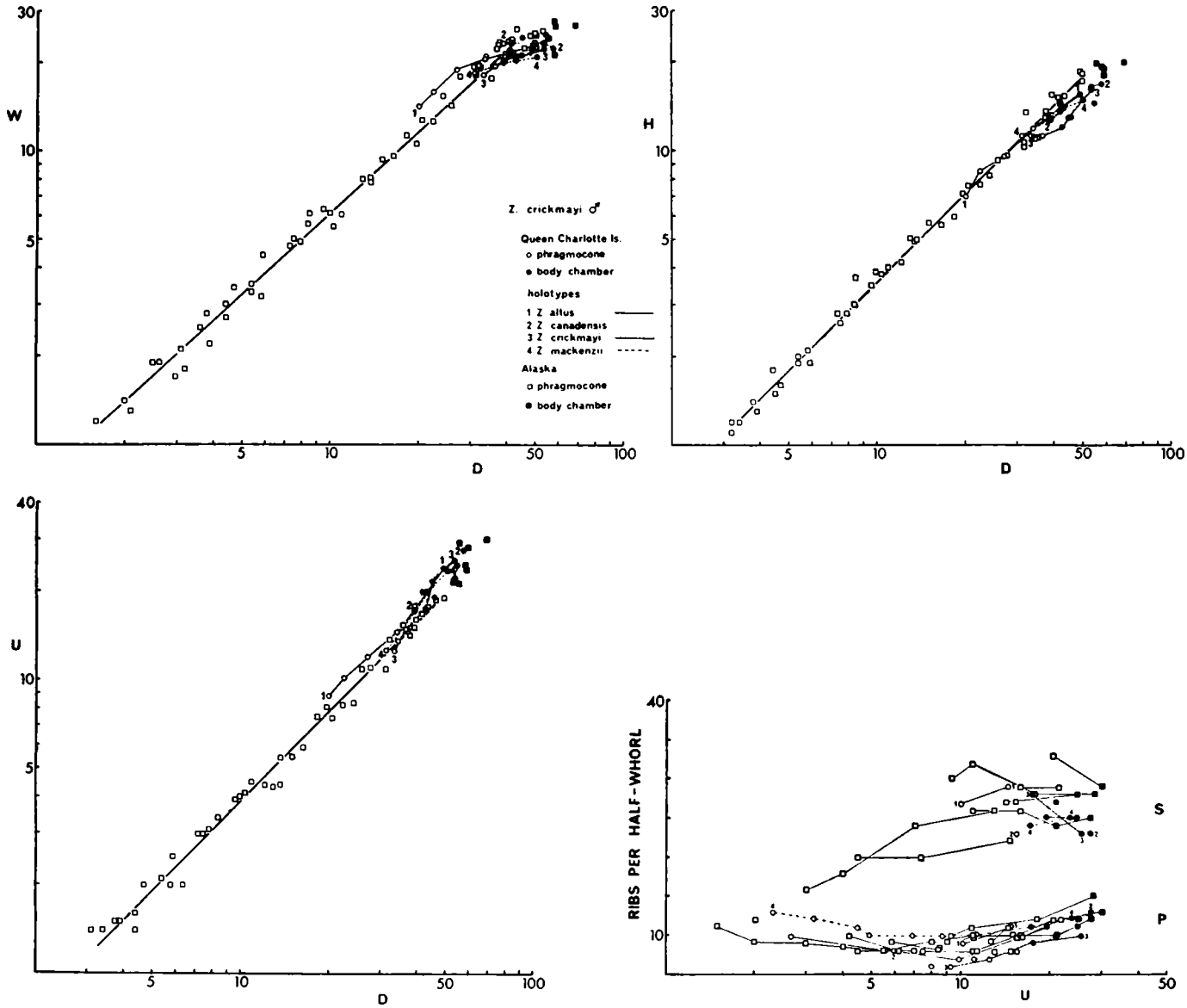


Text-figure 8. — Protoconch of *Zemistephanus crickmayi* ♂, three views of specimen from USGS Mesozoic loc. 2999.4. × 62.

situated at about 30 to 40 percent of the whorl height. These elongated tubercles are directed adapicad onto the upper flanks. They also extend slightly onto the venter as faint, broad undulations when the diameter reaches 4.5 mm. There are 10 per half-whorl. The whorl cross-section is broad and depressed (H/W = 0.50 - 0.60) with a gently arched venter, sharp lateral shoulder and steep, convex flanks. Secondary ribs extend across the venter becoming much stronger than the primary ribs and outnumbering them two to one. Throughout the phragmocone the whorls remain broad and depressed with H/W ratios of 0.55 - 0.65.

The number of primary ribs per half-whorl decreases from 10-13 on the earliest whorls to a minimum of six to eight at umbilical diameters between 5 and 15 mm, then increases again to eight to 12 at the end of the phragmocone (Text-fig. 9). At the same time the number of secondaries per half-whorl increases to a maximum, then declines on the body chamber. The ratio of secondary to primary ribs increases from two on the earliest whorls to three or four, and later decreases on the body chamber to two. Similar variations in ribbing density during ontogeny were also noted on *Z. richardsoni* ♂. Primary ribs are broad and massive, but a little sharper and more curved than on *Z. richardsoni* ♂; the nodes on *Z. richardsoni* ♂ are more massive and conical.

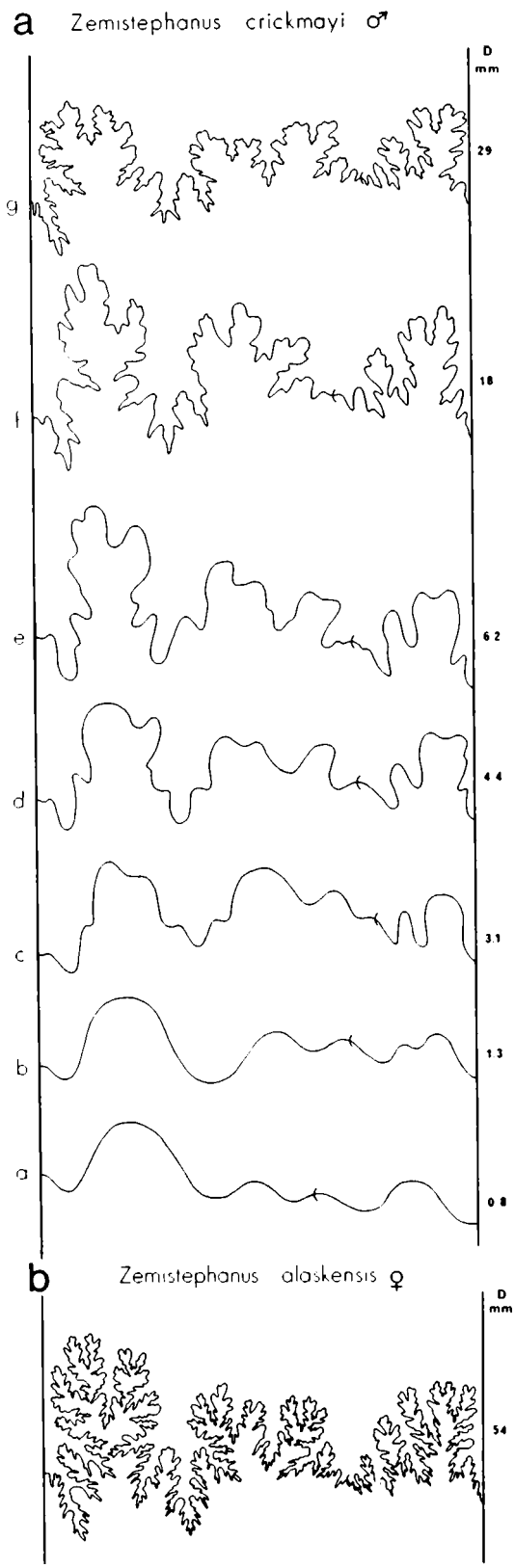
The body chamber is about one whorl in length and is marked by egression of the umbilical seam from the line of nodes on the previous whorl, changes in the shape of the



Text-figure 9. — Bivariate plots of whorl width (W), whorl height (H), umbilical diameter (U), shell diameter (D) and number of primary (P) and secondary (S) ribs per half-whorl, for *Zemistephanus crickmayi* ♂ (with synonyms) from Queen Charlotte Islands and southern Alaska. Based on 13 microconchs, with some individual "growth lines" indicated. See *Measurements* for additional explanations.

cross-section, and ornamentation. The flanks become less steep and the lateral shoulder more rounded with contraction of the whorl (Text-figs. 9, 11); H/W increases to 0.65 - 0.75. Primary and secondary ribbing remain strong to the aperture but the nodes disappear; primary ribs simply bifurcate. There is no constriction or flaring at the aperture, which terminates with lateral lappets.

Remarks. — The narrow body chamber clearly separates *Z. crickmayi* ♂ from *Z. richardsoni* ♂, which occurs in the same beds at MacKenzie Bay. There is only a slight increase in the H/W ratios on the body chamber of that species from 0.50 - 0.60 to 0.62 - 0.64 while the values for the same stages on *Z. crickmayi* ♂ are 0.55 - 0.65 and 0.65 - 0.75 respectively. The venter on the body chamber of *Z. crickmayi* ♂ is more highly arched, having a radius of curvature of 80-100 mm at a diameter of 30 mm compared with a radius of curvature of 140 mm for *Z. richardsoni* ♂ at the same size. In addition, the nodes on the phragmocone whorls of *Z. richardsoni* ♂ are more massive and the primary ribs broader and rectiradiate. In the material described by Imlay



Text-figure 10. — a, sutural ontogeny for *Zemistephanus crickmayi* ♂ (USGS Mesozoic loc. 2999.4); b, adult septal suture of *Z. alaskensis* ♀ (USGS Mesozoic loc. 2999.2), both from southern Alaska. The position of the umbilical seam is marked by (. See *Measurements* for additional explanations.

(1964) as *N. (I.) crickmayi* (McLearn) from southern Alaska he notes a considerable variation in whorl width; this suggests that specimens of both *Z. richardsoni* ♂ and *Z. crickmayi* ♂ are present, but because of crushing are not easily distinguished.

McLearn (1929) distinguished four "species" in his genus *Kanastephanus* based on minor differences in coiling and ornamentation. *K. altus* supposedly has a higher whorl section and wider umbilicus but growth curves (text-fig. 9) show these differences to be very small and intermediate between values from other specimens. *K. mackenzii* was separated from the type species because it had more primary ribs but the number per half-whorl is the same as on *K. altus* and some of the Alaskan material. *K. canadensis* has slightly broader whorls than the other "species" but is identical in this character with McM J1798a from MacKenzie Bay as well as some of the Alaskan material. Each of McLearn's four "species" was based on a single specimen, and the minor variations used to distinguish them are here shown to lie within the range of variation of all the available material representing *Z. crickmayi* (McLearn) ♂ (Text-fig. 9).

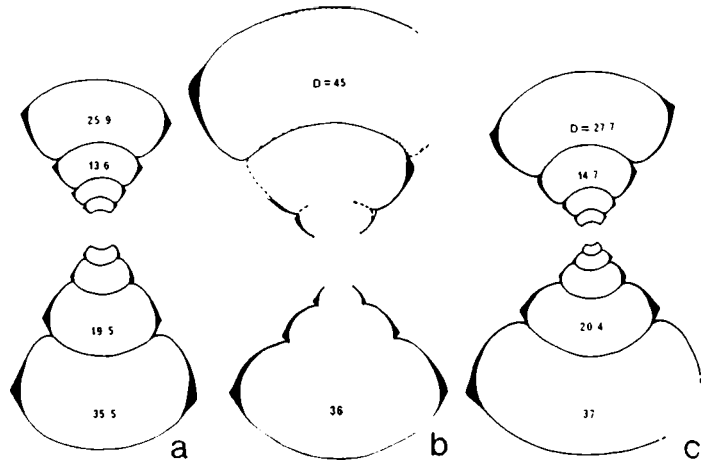
This species could be the microconch of *Z. carlottensis* (Whiteaves) ♀, the holotype of which (Pl. 3, fig. 1a, b) probably came from the MacKenzie Bay locality. Both forms are characterized by coronate early whorls that contract strongly on the body chamber, losing the steep flanks. At the same time relative whorl height increases with stronger arching of the venter. Broad ribbing persists to the aperture on both forms with a decline in the density of the secondaries to two per primary and loss of the prominent nodes. However, the only other macroconchs described are from southern Alaska (Imlay, 1964) and the few entirely septate specimens are too crushed to allow significant comparisons of whorl dimensions to be made.

Zemistephanus alaskensis n. sp. ♀

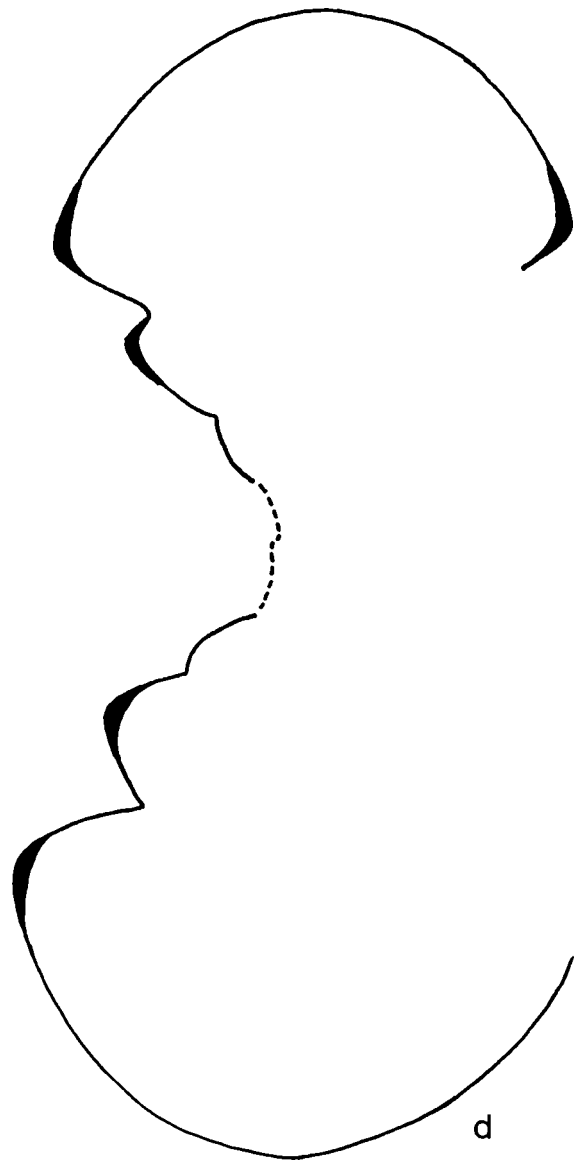
Plate 5, figure 1; Text-figure 12

1964. *Teloceras itinsae* McLearn; Imlay, p. B50; pl. 23, figs. 9, 10 (holotype); pl. 24, figs. 5, 7, (? 1, 2).
1964. *Zemistephanus richardsoni* (Whiteaves); Imlay, p. B51; pl. 25, fig. 6; *non* pl. 25, fig. 7 and pl. 26, figs. 1-7.

Holotype. — USNM 131434, described and figured by Imlay (1964; pl. 23, fig. 10) as a "plesiotype" of *Teloceras*



Text-figure 11. — a-c. Cross-sections of *Zemistephanus crickmayi* ♂, body chamber shaded, × 1.3. a,c, from southern Alaska (USGS Mesozoic loc. 21270.1); b, from Queen Charlotte Islands (McM J1798a); d, cross-section of *Z. warreni* (McLearn) ♀, fully septate holotype from southern Alberta, × 1. See *Measurements* for additional explanations.



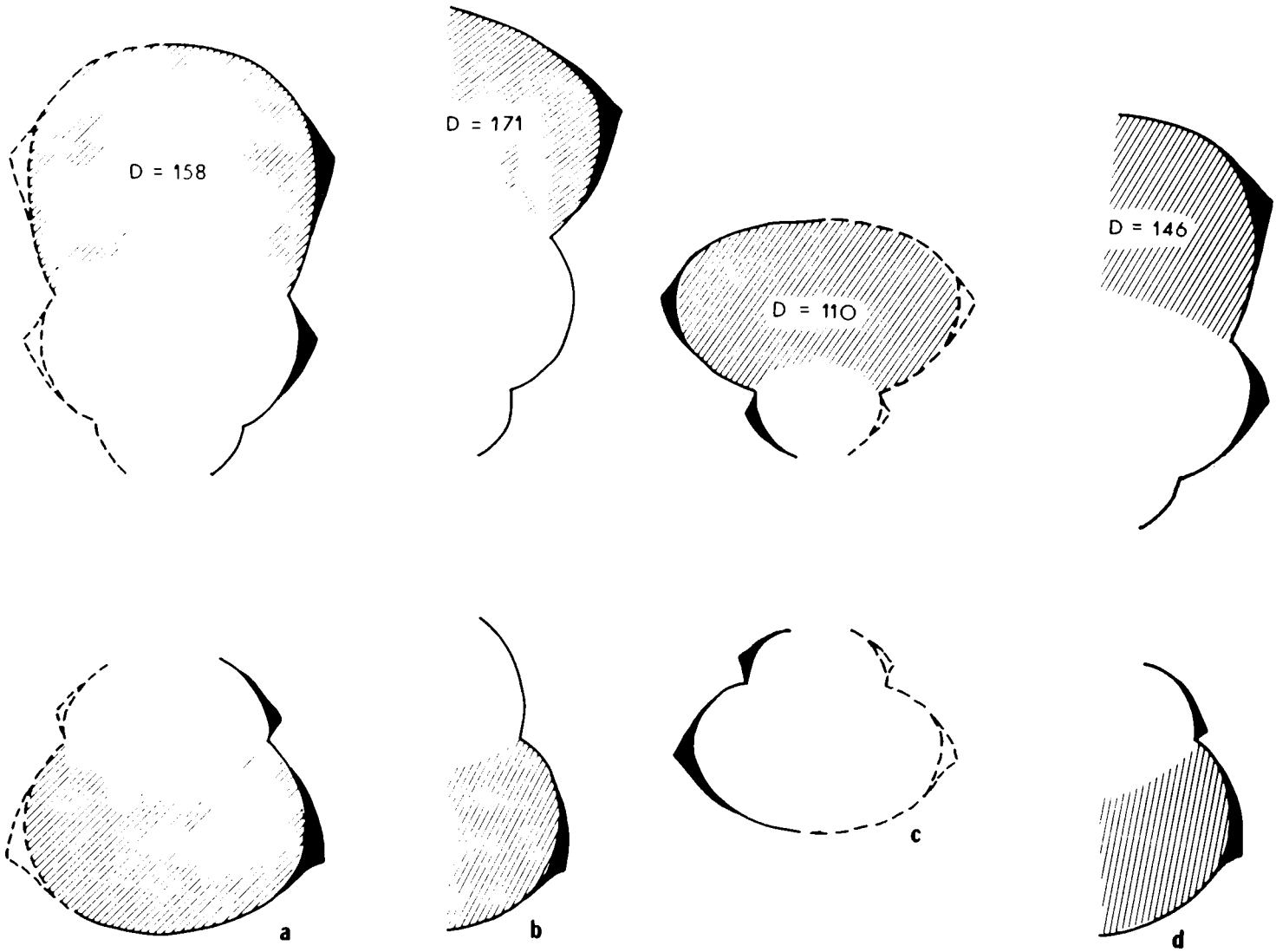
itinsae McLearn, from USGS Mesozoic loc. 21270 in the Fitz Creek Siltstone of the Tuxedni Group, Tuxedni Bay, southern Alaska.

Material. — Three other specimens from the Fitz Creek Siltstone in southern Alaska (USGS Mesozoic locs. 2999, 21270; McM J1245); two specimens (McM J1858a, J1858b) from shales 30 m above the base of the Yakoun Formation exposed at MacKenzie Bay and another (McM J1859) from 20 m above the base of the section.

Description. — The large macroconch reaches a maximum diameter of at least 155 mm and has a simple, slightly flared aperture. The phragmocone whorls are depressed (H/W = 0.55 - 0.65) with steep flanks, producing a deep umbilicus. The venter is broad and gently arched, curving sharply onto the flanks along the line of nodes to form an abrupt lateral shoulder (Text-fig. 12). The primary ribs are broad, rounded undulations, rectiradiate or slightly curved forward on the inner whorls, eight to 10 per half-whorl, and end in prominent, conical nodes. The secondaries are prominent and broad, curving forward from the nodes and then crossing straight over the venter, usually with 3 to 3.5 to each primary.

The body chamber is a single whorl long and is marked by changes in growth ratios similar to those in other species of the genus. Egression of the umbilical seam from the line of nodes on the previous whorl begins just after the last septum; it is accompanied by strong contraction of the whorl with a marked decrease in the absolute whorl width and increase of whorl height so that H/W ratios change from 0.60 - 0.65 at the end of the phragmocone to 0.80 - 0.95 near the aperture. The flanks become less steep and are nearly flat near the aperture while the undulations forming the primary ribs almost disappear. The conical nodes remain prominent to the aperture but occur successively higher on the flanks: at 36 percent of the whorl height on the phragmocone, 40 to 45 percent on the early body chamber and 55 to 60 percent near the aperture. Strong, coarse secondary ribs persist to the end, usually 3 to 3.5 to each node. On two specimens (USNM 131437 and another from USGS Mesozoic loc. 21270) secondary ribbing is obscured by fine striae (shell) on the early parts of the body chamber.

Remarks. — Coronate inner whorls with steep flanks, low and distant primary ribs on the juvenile whorls becoming



Text-figure 12. — Cross-sections of *Zemistephanus alaskensis* ♀, body chamber shaded, × 0.8. a, holotype from southern Alaska; b-c, from Queen Charlotte Islands (McM J1858a, McM J1858b); d, from southern Alaska (USNM 131433). See *Measurements* for additional explanations.

shortened on the intermediate and obsolete on the outer whorls, large conical nodes set low on the whorl, and a gently arched venter are all features of the other known species of *Zemistephanus*. The marked egression of the body chamber with decrease in whorl width and flattening of the flanks, and the persistence of the large, rounded nodes to the aperture are also diagnostic of the genus. This species differs from *Z. richardsoni* (Whiteaves) ♀ in the persistence of broad secondary ribs to the aperture, the higher position of the nodes on the flanks and the more rounded section of the body chamber. The density of secondary ribbing on the phragmocone is lower and the ribs are much coarser than on *Z. richardsoni*. *Z. carlottensis* (Whiteaves) ♀ differs in the steeper umbilical slope, the lower position of the lateral shoulder (at 30 to 35 percent of the whorl height), and in lacking the large, conical nodes on the last part of the body chamber. The similar *Stephanoceras (Stemmatoceras)* ex. gr. *acuticostatum* (Weisert) [*"Teloceras itinsae"* McLearn] differs in the prominent primary ribs on all but the ultimate one or two whorls.

The corresponding microconch is unknown.

Zemistephanus warreni (McLearn, 1930) ♀

Plate 4; Text-figure 11d

1930. *Teloceras warreni* McLearn, p. 3, pl. 1, fig. 4.
1932b. *Teloceras warreni* McLearn; McLearn, p. 113, pl. 3, fig. 4.

Description. — The holotype (re-examined by Hall) from Porcupine Creek in southern Alberta is almost entirely septate with a short and very steep to vertical umbilical wall, an acute shoulder with round, blunt nodes and a highly arched and rounded venter (see Text-fig. 11d). Gradual uncoiling commences one whorl before the end of the specimen. Primary ribs on the inner whorls are straight and broad; there are five secondary ribs per primary and they are strongly protracted except on the ultimate whorl.

Remarks. — McLearn (1932b, p. 113) noted the similarities of the inner whorls to *Zemistephanus 'vancouveri'* McLearn [= *Z. richardsoni* (Whiteaves) ♀] from the lower Yakoun Formation of the Queen Charlotte Islands. The nature of the umbilicus, whorl section and ribbing all indicate the close affinity to *Zemistephanus* that justifies the transfer of this species to that genus.

Genus **STEPHANOCERAS** Waagen, 1869, and allies

Type species. — *Ammonites Humphriesianus* J. de C. Sowerby, 1825 (by subsequent designation of Buckman, 1898). Type specimen refigured by Buckman (1908; pl. VII, fig. 1).

Discussion. — The taxonomy of the group of ammonites closely allied to *Stephanoceras* Waagen has long been the subject of disputed and varied treatments, especially in Europe; our treatment for North American forms is summarised in Table 5. Early nomenclatural difficulties arose because of doubt concerning the availability of the name *Stephanoceras;* this was settled by Spath (1944) who pointed out that the similar spelling of the older name *Stephanoceros* (Rotatoria) did not invalidate the younger name *Stephanoceras*. He also emphasized that the type species was *Stephanoceras humphriesianum* (J. de C. Sowerby, 1825) by a subsequent designation of Buckman (1898, p. 454).

Because he believed the name was pre-occupied, Buckman had emended *Stephanoceras* to *Stepheoceras*; while rejecting the need for this alteration Mascke (1907), however, retained both names, separating forms with a very wide umbilicus, weaker sculpture and a strongly enlarged aperture as *Stepheoceras* Buckman (group of *Amm. Humphriesi* Sow.). Those species with heavy sculpture, a deeper umbilicus and only slightly enlarged aperture (group of *Amm. Humphriesi mutabilis* Quenstedt) were described as *Stephanoceras* (Waagen) em. Mascke. In addition he proposed three new groups of closely related stephanoceratids: *Stemmatoceras, Skirroceras* and *Teloceras*.

Stemmatoceras (type species: *Amm. Humphriesianus coronatus* Quenstedt, 1886-1887 [= *S. frechi* Renz, 1913] included forms with a wide umbilicus, depressed whorls and medium-strong sculpture that declines on the body chamber, particularly in the density of secondary ribbing. *Skirroceras* (type species: *Amm. Humphriesianus macer* Quenstedt, 1886-1887) was separated from *Stemmatoceras* because of its more strongly incised sutures and wider, shallower umbilicus. The whorls expand only slowly and are rounded in section with less inflated flanks on the body chamber. *Teloceras* (type species: *Amm. Blagdeni* J. Sowerby, 1818) was characterized by great whorl thickness with large nodes on a sharp lateral edge that persisted onto the body chamber, with a decline in the strength of ornament (and rounding) only near the aperture.

In his extensive review of the group Weisert (1932) recognized only three subgenera: *Stephanoceras* Waagen, 1869; *Stemmatoceras* Mascke, 1907; and *Teloceras* Mascke, 1907. He included Mascke's *Skirroceras* and *Stepheoceras*, along with many of Buckman's vast array of stephanoceratid "genera" (*Kallistephanus, Skolekostephanus, Rhytostephanus, Oecostephanus, Stegeostephanus, Mollistephanus, Kumatostephanus*) in *Stephanoceras* Waagen. In addition he gave a detailed discussion of the characters distinguishing each genus throughout ontogeny. Included under *Stephanoceras* Waagen were forms with a rounded whorl section, narrow to broad umbilicus, ribbing of variable strength and low, fine nodes; the suture was said to be strongly differentiated, with a greatly subdivided "1st lateral" saddle (? L/U). *Stemmatoceras* on the other hand was defined as having a poorly-differentiated suture (of which the "1st lateral" saddle was only weakly subdivided), strong, high and pointed nodes with a decrease in the strength of the sculpture on the body chamber. The whorls are not as rounded as in *Stephanoceras* and the venter is not as highly arched.

While following Weisert's treatment of these closely related forms, Schmidtill and Krumbeck (1938) again made clear reference to the existence of two groups within *Stephanoceras* Waagen as defined by Weisert. They emphasized the contrast between the group [containing *S. umbilicum* (Quenstedt), *S. auerbachense* Schmidtill and Krumbeck and *S. mutabile* (Quenstedt)] having a deep umbilicus of medium width with broader whorls and the group [including *S. humphriesianum* (Sowerby), *S. zieteni* Quenstedt and *S. scalare* Mascke] that all have a fairly wide, shallow umbilicus, discoidal form and higher whorl section (p. 324). These two groups would correspond to "*Stephanoceras* Waagen emend. Mascke" and "*Stepheoceras* Buckman" respectively. Schmidtill and Krumbeck did not add further

to the discussion of the genus *Skirroceras* Mascke, but unlike Weisert (1932) accepted as reasonable its separation from *Stephanoceras*.

The lectotype of *S. humphriesianum* (BM 43908a) is an entirely septate specimen cut along the sagittal plane; thus nothing is known of its adult size or body chamber characteristics. It does, however, have a rounded whorl section lacking a lateral shoulder, and a broad, shallow umbilicus with uncoiling commencing at least one complete whorl (at U = 20-25 mm) before the end of the specimen. Clearly this specimen is closer to the more serpenticone forms of *Skirroceras* than to the *S. mutabile* group and its designation as type species is unfortunate. Also, those forms with broader whorl section, less arched venter, stronger and sharper ornamentation and later uncoiling (often at U = 40 mm and corresponding to the beginning of the body chamber) bear close resemblance to *Stemmatoceras*. Indeed, Schmidtill and Krumbeck (1938, p. 325) noted that forms of *S. mutabile* "show a remarkable approach" to *Stemmatoceras*. Other members of this group would include *S. brodiaei* (J. Sowerby), *S. umbilicum* (Quenstedt), *S. itinsae* (McLearn) and *S. skidegatense* (Whiteaves). However, it must be emphasized that whatever character or combination of characters is used, there are transitional forms that are difficult to place confidently in one subgenus. There seems to be a morphological trend within the group for prolongation of the juvenile morphology into successively mature stages of the conch (Mouterde *et al.*, 1971, p. 12). Thus in the early *Skirroceras* only the innermost whorls are comparatively involute cadicones; in *Stephanoceras* and *Stemmatoceras* these characters are present in successively later stages until in *Teloceras* most of them persist to the aperture. However, at any stratigraphic level a variety of morphologies is commonly present, and only the proportions change with time (Callomon, pers. comm.).

Although *Teloceras*, with its broad, coronate whorls persisting to the body chamber, is clearly distinguishable in the adult stage, a number of specimens of this genus from western Europe and South America examined by us (including *T.* cf. *blagdeni* from Goslar, Germany) have juvenile whorls that are difficult to distinguish from those of the relatively broadwhorled *Stephanoceras* spp. And, on the other hand, stephanoceratids with typically coronate juvenile and even intermediate stages may have rounded and planulate adult whorls (i.e., *Stemmatoceras*). This again emphasizes the necessity of basing systematic distinctions on mature specimens with known body chamber characteristics; many European species have been erected for small, incomplete phragmocones; *e.g.*, the small nucleus that is the type specimen of *S. umbilicum* Quenstedt makes proper interpretation of that species impossible. Similar situations have led to blurring of the distinctions between *Teloceras*, *Zemistephanus*, *Stephanoceras* s.s. and *S. (Stemmatoceras)* in North America.

In the *Treatise on Invertebrate Paleontology*, Arkell (1957, p. L289) treated *Skirroceras* as a subgenus of *Stephanoceras* s.s., but retained *Stemmatoceras* and *Teloceras* as separate genera. The great variety of *Stephanoceras s.l.* and the morphological resemblance of some species to *Stemmatoceras* require that *Stemmatoceras* also be classified as a subgenus of *Stephanoceras* Waagen. Indeed, many specimens from North America that have been identified with *Stemmatoceras* are here shown to be closely allied with the *Stephanoceras itinsae* group and thus also to the broader-whorled European forms such as *Stephanoceras mutabile* (Quenstedt). *Stemmatoceras* is here treated as a subgenus of *Stephanoceras* because of the gradation in characters previously used to distinguish the two "genera".

Teloceras is also linked to *S. (Stemmatoceras)* by a series of morphologically intermediate species that are characterized by relatively narrow but, nevertheless, coronate immature whorls bearing distant primaries with prominent nodes; the rounded adult ultimate whorl (including body chamber) clearly distinguishes complete specimens from typical *Teloceras*, but it is commonly crushed or missing. This is the species group of *'T.' acuticostatum* Weisert, including the poorly known European species *'T.' subblagdeni* Schmidtill and Krumbeck and *'Cadomites' blagdeniformis* Roché, which seems to be diagnostic of the lower *T. blagdeni* subzone in western Europe (Schmidtill and Krumbeck, 1938; Mouterde *et al.*, 1971) and continues to be classified in *Teloceras*. Following this precedent, Westermann (1964a, p. 68) originally classified *Stemmatoceras* tentatively as a subgenus of *Teloceras*, but later (Westermann and Getty, 1970, p. 248) because of the abundance of the intermediate forms (*'T.' acuticostatum* group) in the circum-Pacific area, united all in the single genus *Stephanoceras*. Most if not all of the North American species originally described under *Teloceras* belong to the *'T.' acuticostatum* group (see below), and this entire group is here transferred to *Stephanoceras (Stemmatoceras)*. Because of the great morphologic distance of the typical European *Teloceras* (the *T. blagdeni* group) from the typical *Stephanoceras* (*s.s.*), because of its well-established chronologic significance, and after consulting a number of colleagues, we decided to distinguish *Teloceras* at the genus level. *Teloceras* is thus redefined to include (in the macroconchs) only the extremely and completely coronate, large Stephanoceratinae.

Taxonomic treatment of microconch Stephanoceratidae has been varied. Westermann (1954) originally retained *Itinsaites* as a genus, making the Canadian species of *Kanastephanus* synonymous with it, but later (1964a) considered it tentatively as a subgenus of *Teloceras*. Arkell (1957, p. L289) placed *Normannites*, including *Itinsaites*, *Kanastephanus*, and others, in the family Otoitidae. McLearn (1949) and Imlay (1964) both regarded *Itinsaites* as a subgenus of *Normannites;* Imlay also transferred the four Canadian species of *Kanastephanus* to *Itinsaites*, rightly suggesting that they probably represented a single species. However, all *Kanastephanus* spp. differ from *Itinsaites* and are in fact the corresponding microconchs to *Zemistephanus* McLearn ♀.

The pairing of macroconch-microconch genera in the family Stephanoceratidae was attempted by Westermann (1964a); microconch equivalents for *Stephanoceras*, *Stemmatoceras*, *Teloceras*, *Zemistephanus*, *Kumatostephanus* and *Cadomites* were sought in the several subgenera or related genera of *Normannites* Munier-Chalmas. *Itinsaites* McLearn is here shown to be the microconch equivalent to *Stephanoceras* s.s. (part). Comparison of the inner whorls of *Itinsaites itinsae* McLearn and *Stephanoceras yakounense* McLearn from South Balch Island in the Queen Charlotte Islands indicates they are corresponding dimorphs (see discussion under *Stephanoceras itinsae*). *Itinsaites* McLearn thus becomes a junior subjective synonym of *Stephanoceras* Waagen.

Table 5. — Genus-group classification of North American *Stephanoceras* and *Teloceras* (excluding Mexico).

Stephanoceras (Stephanoceras)
- *Ammonites skidegatensis* Whiteaves (♀)
- *Stephanoceras skidegatensis* var. *laperousii* McLearn (♀)
- *S. yakounense* McLearn (♀)
- *S. obesum* Imlay (♀)
- *S. caamanoi* McLearn (♀)
- *Stemmatoceras mclearni* Warren (♀)
- *S. carri* Warren (♀)
- *S. arcicostum* Imlay (♀)
- *Itinsaites itinsae* McLearn (♂)
- *Dettermanites vigorosus* Imlay (♂)
- (?) *Normannites (Itinsaites?) variabilis* Imlay (♂)

Stephanoceras (Skirroceras)
- *S. (Sk.) juhlei* Imlay (♀)
- *S. (Sk.) nelchinanum* Imlay (♀)
- (?) *S. (Sk.) kirschneri* Imlay (♀) [?*Stephanoceras s.s.*]
- (?) *S. (Phaulostehanus?) oregonense* Imlay (♀)

Stephanoceras (Stemmatoceras)
- (?) *Stemmatoceras albertense* Warren (♀) [?*Stephanoceras s.s.*]
- (?) *Stemmatoceras palliseri* Warren (♀) [?*Stephanoceras s.s.*]
- *Teloceras dowlingi* McLearn (♀)
- (?) *Teloceras itinsae* McLearn (♀)
- (?) *Teloceras allani* Warren (♀)

Stephanoceras (Phaulostephanus)
- *S. mowichense* Imlay (♀)

Teloceras
- (?) *Teloceras stelcki* Warren (♀), nom. dub.
- *Zemistephanus crickmayi* Frebold (♀)

Subgenus **STEPHANOCERAS** Waagen, 1869

Stephanoceras (Stephanoceras) itinsae (McLearn, 1927) ♀ + ♂

Plates 6-8; Plate 11, figure 1; Plate 13, figure 2; Text-figures 13-17

S. itinsae ♀ (macroconch) ['**S. yakounense**']

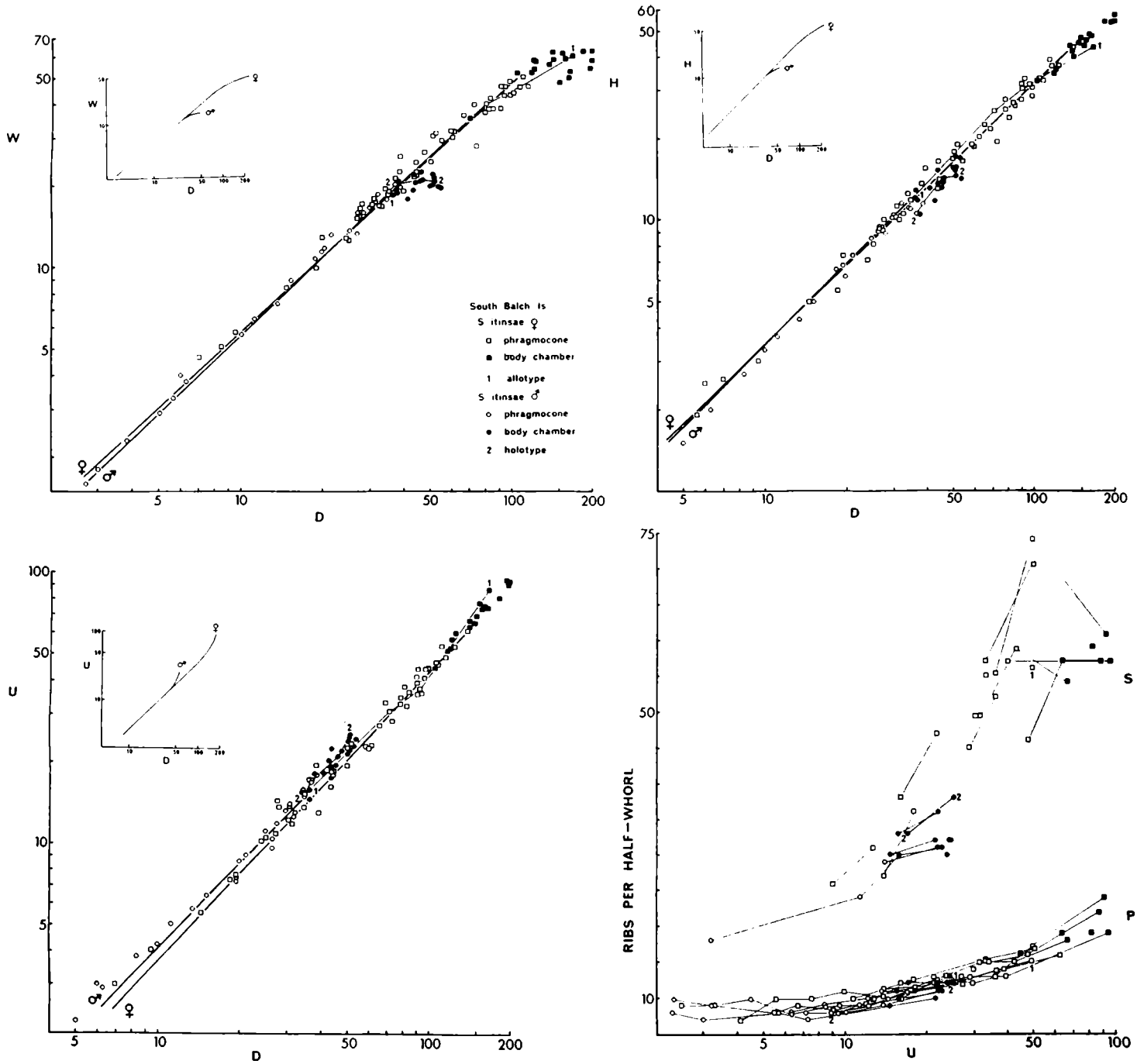
1900. *Perisphinctes skidegatensis* Whiteaves, p. 278 (part).
1930. *Stephanoceras yakounense* McLearn, p. 5; pl. 1, fig. 1 (holotype).
1932a. *Stephanoceras yakounense;* McLearn, p. 56; pl. 2, fig. 1 (holotype); pl. 3, figs. 2 and 6 (holotype), 4; pl. 5, fig. 10.
1947. *Stemmatoceras mclearni* Warren, p. 68; pl. I, fig. 1; pl. IV, fig. 1.
1947. *Stemmatoceras carri* Warren, p. 69; pl. IV, fig. 2; pl. VI, fig. 3.
1957. *Stephanoceras* ex. gr. *skidegatense* (Whiteaves); Frebold, p. 49; pl. XXI, fig. 1; pl. XXII, fig. 2; pl. XXV, fig. 2.
1957. *Stemmatoceras albertense* McLearn; Frebold, p. 50; pl. XXI, figs. 2a, b; pl. XXIII, figs. 1a-c.
(?)1964. *Stephanoceras obesum* Imlay, pp. B45, 46; pl. 18, figs. 5-11.
1967. *Stephanoceras* cf. *skidegatense* (Whiteaves); Imlay, p. 89; pl. 6, fig. 10.

Allotype. — GSC 9057, almost complete internal mold from lower part of the Yakoun Formation, South Balch Island, Skidegate Inlet, Queen Charlotte Islands [originally the holotype of *Stephanoceras yakounense*].

Material. — Ten relatively complete specimens (McM J1807-J1811) and many other phragmocone and body chamber fragments were collected from the type locality; also from this locality four specimens, three almost complete, were collected by Sutherland Brown (GSC locs. 48601, 44711) and two were collected by McLearn (GSC loc. 13634). One complete specimen and another with part of the body chamber were collected by Sutherland Brown from Reef Island, Queen Charlotte Islands (GSC loc. 40985). Two specimens, one almost complete, were obtained on loan from University of British Columbia Museum, labelled "Skidegate Channel". The holotype and McLearn's "plesiotype" were also reexamined.

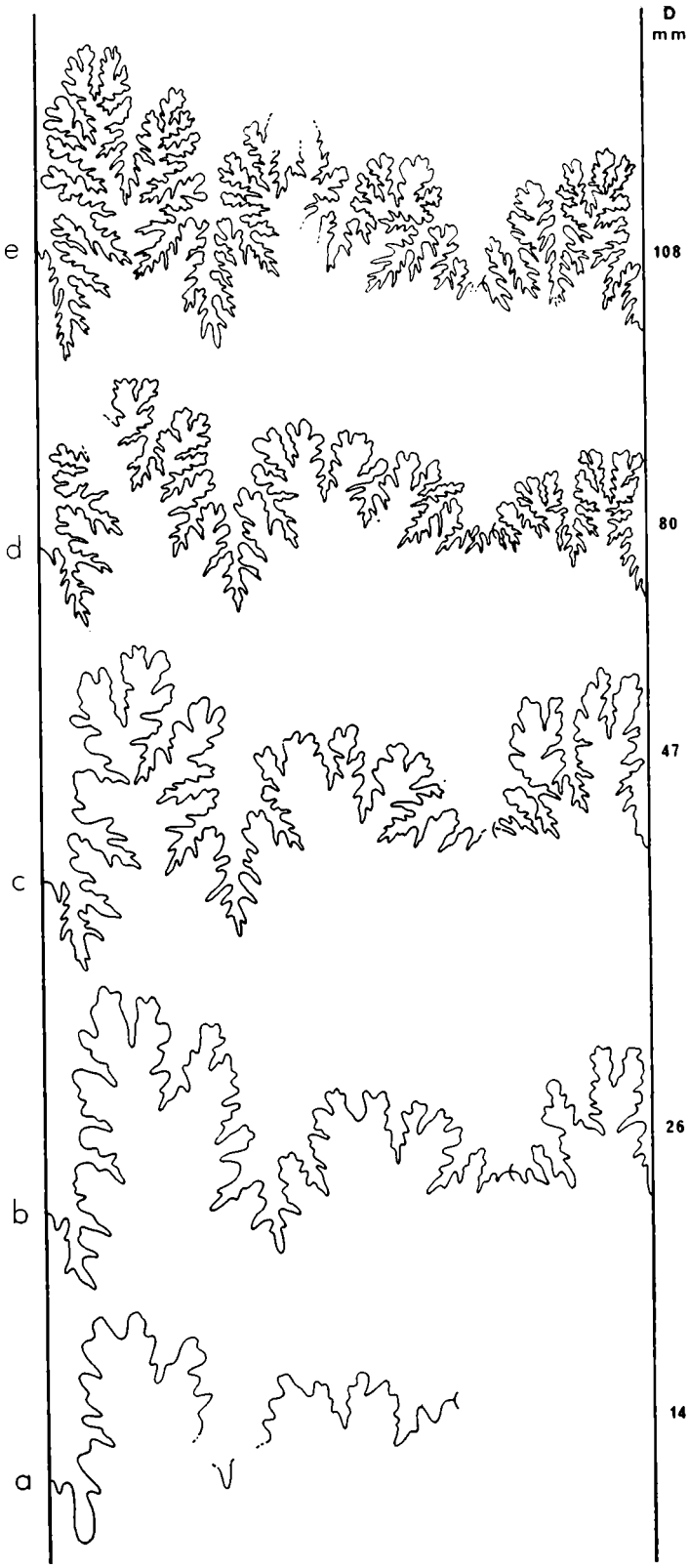
Description. — Maximum diameters of adult macroconchs range from 148-196 mm. One body chamber fragment is 66 mm wide at the position of the last septum and must have attained a greater diameter than any of the complete specimens examined; most adults reach a maximum width of only 60-65 mm at the aperture.

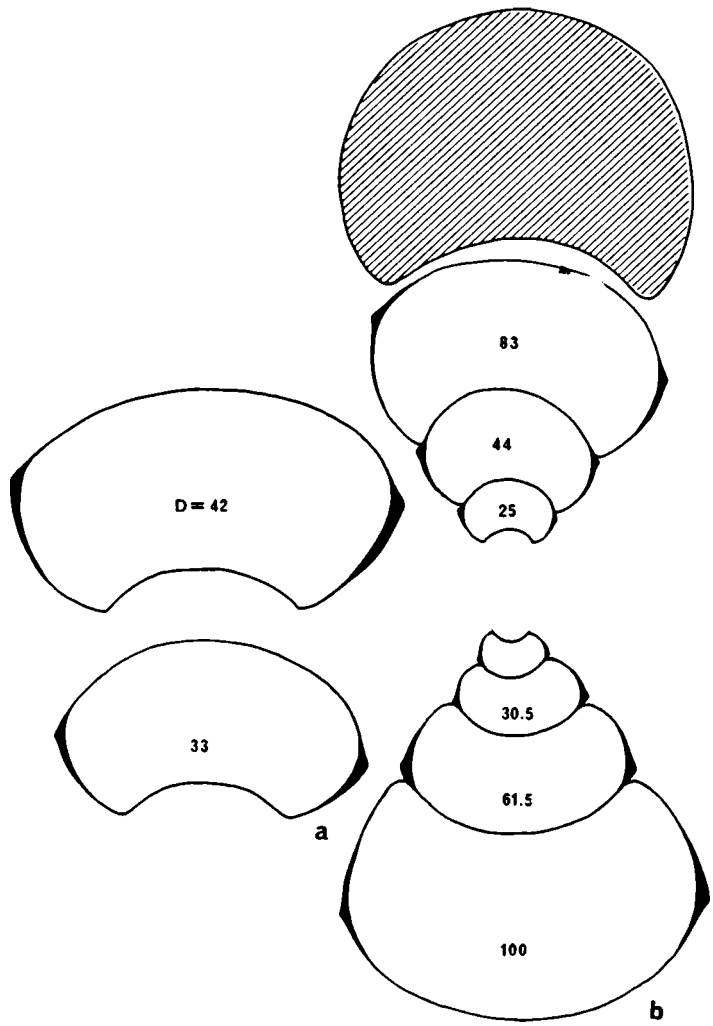
On the phragmocone the umbilical seam lies along the outer edge of the line of tubercles on the previous whorl. The whorls are much wider than high, with H/W = 0.55 - 0.60 at diameters up to 30 mm; throughout ontogeny the whorls gradually become higher with H/W = 0.65 - 0.70 between diameters of 30 and 100 mm (Text-fig. 15). The flanks are strongly convex with strong primary ribs that are prorsiradiate and curved forwards, ending in prominent conical tubercles situated at about 50-65 percent of the whorl height and just above the point of maximum whorl width.



Text-figure 13. — Bivariate plots of whorl width (W), whorl height (H), umbilical diameter (U), shell diameter (D) and number of primary (P) and secondary (S) ribs per half-whorl, for *Stephanoceras itinsae* ♀ and ♂ from Queen Charlotte Islands. Based on 13 macroconchs and 11 microconchs, with some individual "growth lines". Insets are same graphs showing approximate positions of best-fit mass curves. See *Measurements* for additional explanations.

The flanks curve strongly onto the broadly arched venter forming a lateral shoulder, particularly on the phragmocone whorls. The number of primary ribs per half-whorl increases slowly throughout ontogeny from nine at diameters up to 30 mm to 11 at 30-60 mm and 14 at 60-100 mm. The number of secondary ribs also increases from three per primary on the inner whorls to a maximum of four to five at diameters between 40 and 50 mm and declines again to three on the





Text-figure 15. — Cross-sections of *Stephanoceras itinsae* ♀, body chamber shaded; a, (McM J1808g) × 2; b, (UBC collections) × 1. See *Measurements* for additional explanations.

body chamber. Secondary ribs curve slightly forward from the tubercles, then cross straight over the venter.

At diameters between 95 and 130 mm uncoiling begins (this corresponds to U = 35 - 55 mm), the umbilical seam moving gradually away from the line of tubercles, coiling thus becoming more serpenticone. The point at which uncoiling begins also corresponds closely to the beginning of the adult body chamber, which is usually three-quarters of a whorl in length. On the body chamber there is a decrease in relative whorl height and width (Text-fig. 13) while the

Text-figure 14. — Sutural ontogeny of *Stephanoceras itinsae* ♀, incomplete for early stage (a-b, McM J1810b; c-d, J1808g; e, J1809a). The position of the umbilical seam is marked by (. See *Measurements* for additonal explanations.

H/W ratio increases to about 0.75 and may be as high as 0.90. In cross-section the body chamber is oval with the venter more strongly arched than on the phragmocone, is almost as high as wide, and has lost the strong lateral shoulder seen on the phragmocone. There are 16-26 (average 19) primary ribs on the final half-whorl and commonly three secondaries to each primary rib. However, the secondary ribs are much fainter than the primary ribs while the tubercles become blunt and rounded and sometimes obsolescent near the aperture. The primary ribs are shorter near the aperture, extending to less than 50 percent of the whorl height. The aperture is marked by a broad, shallow constriction and a slightly expanded, complete lip.

The suture is moderately deeply incised (Text-fig. 14). E is deep and narrow, L fairly broad, straight and trifid, almost as deep as E. Umbilical lobes are strongly retracted; U_n is deep, narrow, trifid and oblique; I is much deeper and narrow. E/L is broad and high; L/U_2 is very broad and not nearly as high as E/L. The tubercle is situated on L/U_2.

Remarks. — *S. itinsae* ♀ strongly resembles the European species group of *S. umbilicum* (Quenstedt), including *S. mutabile* (Quenstedt) and *S. brodiaei* (Sowerby), all of which are characterized by a relatively deep umbilicus with late uncoiling and broad whorls. *S. umbilicum* at similar diameters has a considerably smaller umbilicus (33 percent of diameter vs. 40-50 percent on *S. itinsae*), and relatively wider whorls (H/W of 0.58 vs. 0.68). *S. mutabile* has a less depressed whorl section than *S. itinsae* (H/W of 0.75 vs. 0.65 - 0.70 on *S. itinsae* at similar diameters). *S. itinsae* most closely resembles *S. brodiaei* differing only slightly in ornamentation: the holotype of *S. brodiaei* has fewer primary ribs (13 per half-whorl vs. 15-24 on *S. itinsae* at D = 105 mm). However, other specimens in the British Museum (Natural History) regarded as conspecific with *S. brodiaei* have ribbing densities closer to those on *S. itinsae* ♀. The tubercles on *S. brodiaei* are larger and more rounded than the sharp, pointed tubercles characteristic of *S. itinsae* ♀.

The single specimen from the Rock Creek Member in Ribbon Creek, southern Alberta, described by Frebold (1957, pp. 50, 51) as *Stemmatoceras albertense* McLearn, is very similar to *S. itinsae* ♀ in whorl dimensions, coiling and the style and density of ribbing but differs from the poorly preserved holotype in the narrower and probably more rounded whorls. The inner whorls are subcoronate, wider than high (H/W = 0.60 - 0.70 at D = 60 mm) with a broadly arched venter, distinct lateral shoulder and steep, convex flanks curving continuously to the umbilical seam that runs along the line of tubercles on the previous whorl. Primary ribs are strong and fairly sharp on the inner whorls, curving forward slightly with small, rounded tubercles at about 50 percent of the whorl height. The secondary ribs are also strong and sharp, passing almost straight over the venter from the tubercles. On the outermost preserved whorl (of which three-quarters is body chamber according to Frebold's description) the primary ribs become broader and less sharp but remain curved, the tubercles decline in strength and the secondary ribs are less sharp but still spaced 3 to each primary. The increase in the density of secondary ribbing from three to four per primary at diameters between 40 and 50 mm with the subsequent decline to three on the body chamber matches similar ontogenetic variation in *S. itinsae* ♀.

Two other specimens described and figured by Frebold (1957, pp. 49, 50; pl. 21, fig. 1; pl. 22, fig. 2; pl. 25, fig. 2) as *Stephanoceras* ex. gr. *skidegatense* (Whiteaves) are more probably identical with *S. itinsae* ♀, having three secondary ribs to each primary on the adult whorls; these secondary ribs are broad, rounded and faint, not sharp as on *S. skidegatense* ♀.

'*Stemmatoceras*' *carri* Warren (Pl. 13, fig. 2) was based on three syntypes characterized by depressed elliptical whorls with rounded lateral shoulders and strong rounding of the outer whorl(s). The long, curved primary ribs were said to be less robust and more closely spaced than those of '*Stemmatoceras*' *mclearni;* the inner flanks of *S. carri* are steeper and more inflated, at least on the early whorls (Warren, 1947, p. 69). The syntype from the headwaters of Sheep Creek near Burns Mine (UA Jr 485) is an incomplete phragmocone (D = 80 mm) with whorl dimensions similar to those of the inner whorls of *Stephanoceras itinsae* (McLearn) ♀ and with similar ornamentation: long, curved primary ribs (16 per half-whorl at D = 80 mm) terminating in small, sharp tubercles and with three to four fine secondary ribs per primary. '*Stemmatoceras*' *carri* is therefore a junior subjective synonym of *S. itinsae.*

The holotype of '*Stemmatoceras*' *mclearni* Warren (UA Jr 192) from Cadomin, Alberta, is very similar to *S. itinsae* ♀ in coiling and ribbing. The almost entirely septate holotype egresses near the end. The umbilicus is broad and shallow with slightly inflated flanks rounding gently onto the venter. On the internal mold the ribs are only of moderate relief, long and curved with small, round tubercles. On the greater part of the ultimate whorl the venter has been badly crushed and corroded but the secondary ribs are of moderate relief, and curve slightly forward from the tubercles, with about three to each primary at the beginning of the body chamber. Umbilical lobes of the suture are strongly retracted. This species bears a close resemblance to *Stephano-*

ceras (*ss.*) *itinsae* (McLearn) ♀, differing only in the shallower umbilicus and slightly coarser primary ribbing, particularly on the nucleus.

Stephanoceras obesum Imlay, from the lower part of the Tuxedni Formation in the Talkeetna Mountains of southern Alaska (Imlay, 1964, p. B45), has almost the same whorl dimensions as *S. itinsae* ♀ on the phragmocone whorls (H/W = 0.70), though the body chamber of the Alaskan material is unknown. Both have closely similar rib densities and style of ornamentation, although the secondaries on *S. obesum* are slightly finer and sharper.

Stephanoceras itinsae ♂ (microconch) ['Itinsaites']

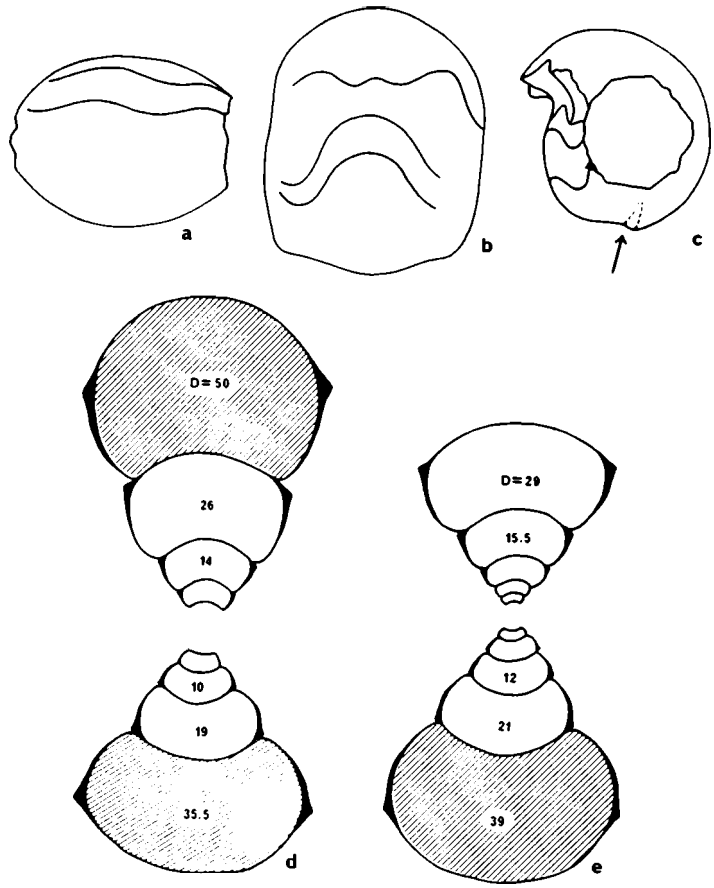
1927. *Itinsaites itinsae* McLearn, p. 73; pl. I, fig. 7 (holotype).
1929. *Itinsaites itinsae;* McLearn, pp. 26, 27; pl. XV, figs. 2, 3 (holotype).
?1933. *Otoites reesidei* Crickmay, p. 912; pl. 27, figs. 9-11.
1949. *Normannites (Itinsaites) itinsae* (McLearn); McLearn, pp. 15, 16.
1954. *Itinsaites itinsae* McLearn; Westermann, pp. 251-254; pl. 26, figs. 5a, b (holotype refigured); pl. 27, figs. 1a, b; text-figs. 101-107.

Holotype. — GSC 9020, from the lower part of the Yakoun Formation, South Balch Island, Skidegate Inlet, Queen Charlotte Islands.

Material. — The holotype was reexamined and compared with three other complete specimens in the GSC collections, Ottawa, that came from the type locality (GSC locs. 13634, 48601). Four complete specimens (McM J1799a-c, J1800) and numerous fragments and incomplete phragmocones (McM J1799d, J1801a, b and J1810b) were collected by Hall from the type locality.

Description. — The protoconch is cigar-shaped, elongated transverse to the plane of coiling, and is more than twice as wide as high (D/W = 0.43) with a diameter in the plane of coiling of 0.43 mm (Text-fig. 16). At a diameter of 0.8 mm the nepionic constriction, seen clearly on the venter, marks the end of the first whorl.

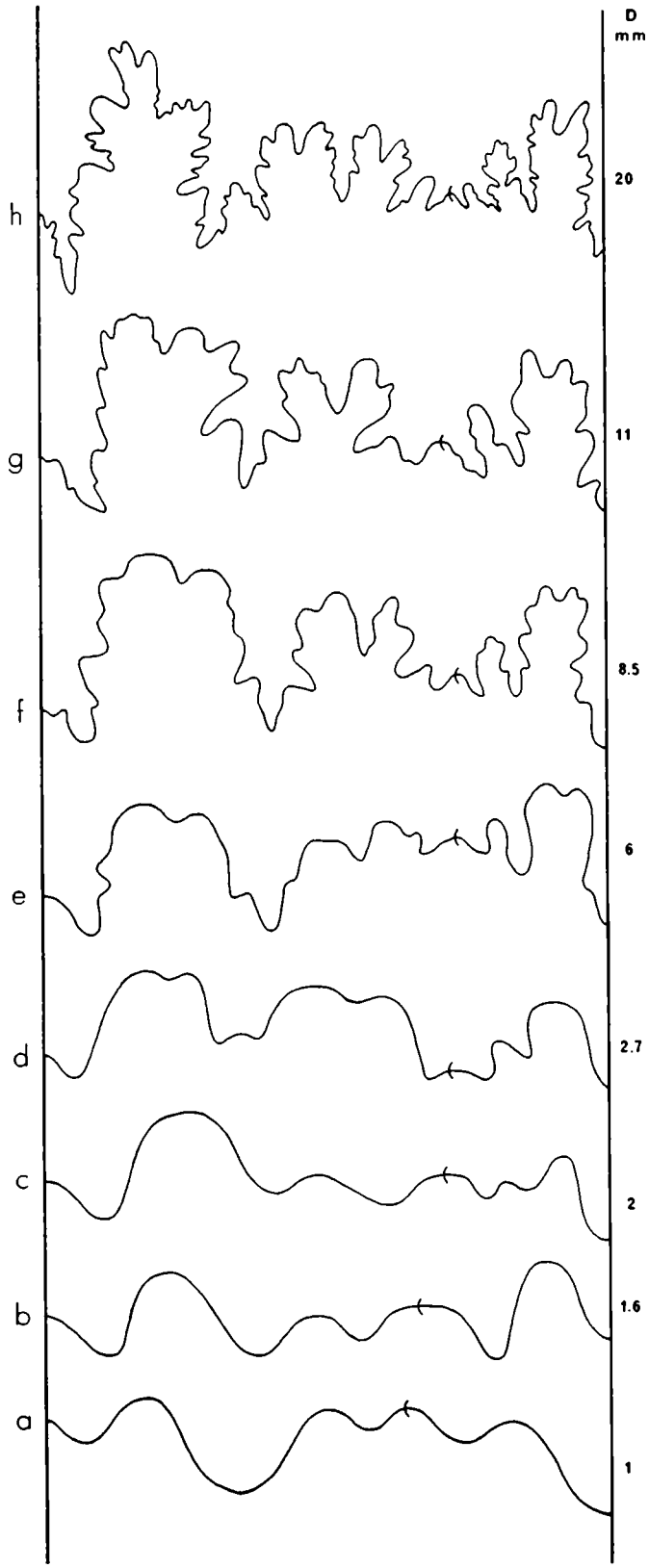
The shell is smooth up to a diameter of 2.5 mm. The first ornamentation consists of broad secondary ribs and small tubercles along the lateral shoulder; primary ribs are faint. By D = 6 mm strong, curved primary ribs extend from the umbilical seam to the lateral shoulder where they terminate in large, conical tubercles; two secondary ribs arise from each tubercle and at this stage in growth are stronger and sharper than the primary ribs. Whorls are much wider than high (H/W = 0.50-0.60) with a broad, slightly arched venter, pronounced lateral shoulder and gently convex flanks curving gradually to the umbilical seam that runs just on the ventral side of the row of tubercles on the previous whorl.



Text-figure 16. — *Stephanoceras itinsae* ♂ from Queen Charlotte Islands. a-b, protoconch (McM J1800) showing two closely spaced protosepta and the primary suture, × 50; c, first whorls with arrow indicating nepionic constriction at D = 0.8 mm (McM J1800), × 20; d-e, cross-sections with body chamber shaded (McM J1799a, J1800) × 1.4. See *Measurements* for additional explanations.

Throughout the phragmocone the whorls are coronate in cross-section, wider than high with a broadly arched venter (Text-fig. 16); the line of tubercles and the position of the pronounced lateral shoulder are a little higher on the flanks than the position of maximum whorl width (as measured between the primary ribs). Secondary ribs remain sharp, stronger than the primary ribs, and increase in density from two per primary to three and even four. Primary ribs are strongly curved forward reaching their maximum strength in the centre of the flanks and becoming weaker towards the umbilical seam and also towards the tubercles; the latter are large, conical and sharp. The number of primary ribs per half-whorl increases gradually from eight at diameters below 30 mm to 10 at diameters between 30 and 40 mm.

The body chamber is half a whorl or a little more in length, but conspicuous uncoiling occurs only in the last



Text-figure 17. — Sutural ontogeny for *Stephanoceras itinsae* ♂ from Queen Charlotte Islands (McM J1800). The position of the umbilical seam is marked by (. See *Measurements* for additional explanations.

quarter-whorl before the aperture where the umbilical seam moves away from the line of tubercles on the previous whorl. This is accompanied by a decrease in relative whorl height and width, with a marked change in the cross-section which becomes rounded with loss of the marked lateral shoulder. There is an increase in the H/W ratio from 0.60 at the end of the phragmocone to 0.70-0.80 just behind the aperture. Ornamentation on the body chamber remains strong, the sharp, curved primary ribs now stronger than the secondaries. There are 10 to 12 primary ribs on the last half-whorl with 2.5 to 3.0 secondary ribs to each primary; tubercles become a little less sharp but are still prominent. The aperture is marked by a narrow constriction with a flared lip extending into long, lateral lappets. Maximum diameters of about 50 mm are attained.

On the single protoconch obtained there are two closely spaced protosepta showing the large, rounded ventral saddle with adjacent narrow lobes on the external suture (Text-fig. 16). U_n appears high on the outer flank of I/U_1 at a diameter of 2 mm. The mature suture is not deeply dissected; E is deep and narrow, L short, broad and trifid and the umbilical lobes strongly oblique. E/L is high and broad and not deeply dissected (Text-fig. 17). The tubercle lies on the ventral edge of L/U_2.

Remarks. — McLearn (1929, pp. 26, 27) separated this species from the other lappet-bearing stephanoceratids on the Queen Charlotte Islands (i.e., '*Kanastephanus*' spp.) on the basis of the greater density of secondary ribbing. '*I.*' *itinsae* maintains a 3:1 ratio of secondary to primary ribs on the body chamber but on '*Kanastephanus*' this ratio declines to 2:1. All other material known from South Balch Island has a density of 3:1. The number of primary ribs per half-whorl increases gradually during ontogeny from eight to 12 whereas on '*Kanastephanus*' this number shows an initial decrease from 10 to 13 to six to eight with an increase to eight to 12 again on later parts of the phragmocone. In addition, the primary ribs on the phragmocone of '*I.*' *itinsae* are sharper and more strongly curved than those of '*Kanastephanus*' and the tubercles are smaller and sharper, persisting onto the body chamber while on '*Kanastephanus*' the tubercles are lost on the body chamber where the ribs bifurcate simply.

Two partially preserved specimens from the Rock Creek Member of the Fernie Group (UA Jr 491, 494) were tentatively placed in '*Itinsaites*' by Warren (1947, p. 73). In-

spection of the figures (pl. VI, fig. 2; pl. VII, fig. 2) indicates ribbing densities similar to those of '*Itinsaites*' but the growth stage is unknown. One complete specimen (Pl. 8, fig. 7) from the Rock Creek Member at Ribbon Creek, southern Alberta, shows the body chamber with three secondary ribs to each primary. The body chamber is a little wider than on the specimens from the Queen Charlotte Islands. Other incomplete small specimens from the Ribbon Creek locality probably also belong to this species.

The single specimen figured by Imlay (1964; pl. 14, figs. 1, 2) as *N. (I.) itinsae* (McLearn) from the Fitz Creek Siltstone, southern Alaska is almost fully septate with only a small part of the body chamber. On the last half-whorl preceding the body chamber the density of the secondary ribs has already declined to 2.4 per primary, which is characteristic of *Zemistephanus* ♂.

'*Otoites*' *reesidei* Crickmay from the Mormon Formation, Mt. Jura, California is very close to *S. itinsae* ♂ in most features, except that in the former the density of secondary ribs on the preserved part of the body chamber declines to 2.5 per primary whereas on specimens of *S. itinsae* ♂ the density usually remains at 3.0.

Dimorphism. — Dimensions and growth patterns throughout the phragmocone whorls of the macroconch and microconch agree closely (Text-fig. 13; Appendices 1, 2; compare Pl. 6, figs. 1c-d with Pl. 8, figs. 3-4); shape of the whorl cross-section is also similar (compare Text-figs. 15, 16d-e). The pattern and density of primary and secondary ribbing correspond closely (Text-fig. 13), both dimorphs showing a gradual increase in the density of secondary ribs from two to four per primary with a decrease to three in later growth stages. This 3:1 ratio is maintained on the body chamber of the microconch (Pl. 8, figs. 2-7) though on the macroconch, ornamentation declines in sharpness and the tubercles almost completely disappear near the aperture (Pl. 7, figs. 2-4). In both dimorphs the body chamber uncoils slowly, becoming relatively higher and rounder in cross-section with loss of the lateral shoulder.

These dimorphs occur together on South Balch Island (Table 2; Text-fig. 1), the macroconch being approximately three to four times the size of the microconch and almost four times as abundant. Specimens of both dimorphs also occur together in the Rock Creek Member of the Fernie Group at Ribbon Creek in southern Alberta. The specific epithet '*itinsae*' (1927) has precedence over '*yakounense*' (1930); the species is renamed *Stephanoceras itinsae* (McLearn, 1927).

Stephanoceras (Stephanoceras) skidegatense
(Whiteaves, 1876) ♀ & ♂

Pl. 9; Plate 10, figure 1; Text-figures 18-21

Stephanoceras skidegatense ♀ (macroconch)

1876. *Ammonites Skidegatensis* Whiteaves, p. 34; fig. 4; pl. 7.
1900. *Perisphinctes skidegatensis;* Whiteaves, p. 278.
1932a. *Stephanoceras skidegatense* (Whiteaves); McLearn, p. 54; pl. I, fig. 2; pl. 2, fig. 3; pl. 3, figs. 8, 9.
1932a. *Stephanoceras skidegatense* var. *laperousii* McLearn; pp. 54, 55; pl. 1, fig. 1; pl. 3, fig. 3.

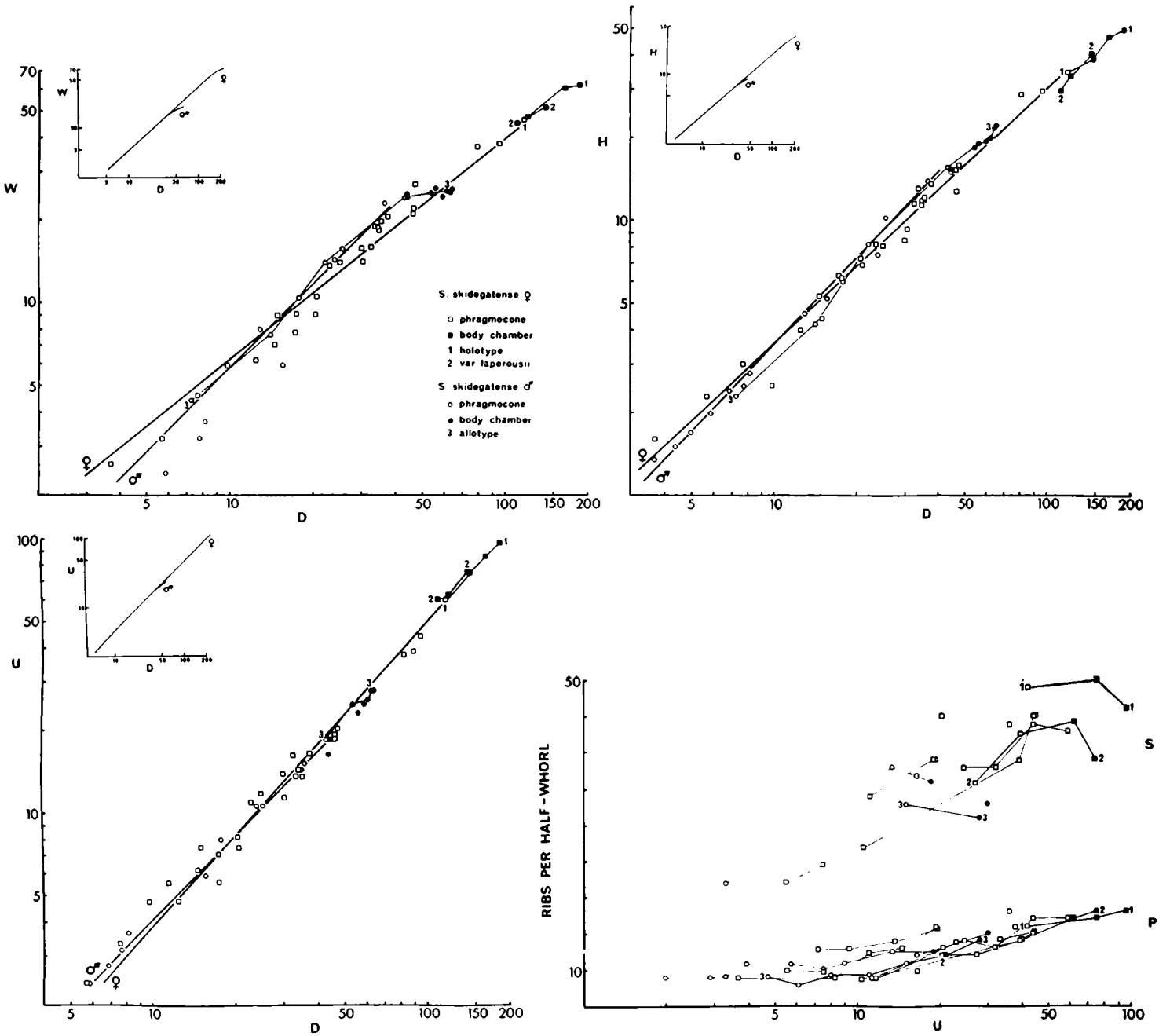
Holotype. — GSC 5011, collected by J. Richardson at Skidegate Inlet in 1872. The precise locality is unknown. However, the only other specimens from the Queen Charlotte Islands referable to this species have all been collected from the lower parts of the Yakoun Formation at Richardson Bay on the south shore on Maude Island (Text-fig. 1).

Material. — Two large body chamber fragments and a number of incomplete phragmocones have been collected by Hall and by Sutherland Brown from the Richardson Bay locality.

Description. — The inner whorls are coronate in cross-section, being much wider than high (H/W ratios average 0.65 up to diameters of 60 mm) with a broadly arched venter curving sharply onto flanks that are at first flat and steep. During ontogeny the flanks become more inflated with gradual rounding onto the more highly arched venter (Text-fig. 19).

Ribbing is very sharp and strongly developed throughout. At diameters less than 30 mm there are nine to 12 primary ribs per half-whorl. These are strongly prorsiradiate with prominent, conical tubercles. There are 2.5 secondaries to each primary. During growth the number of primary ribs increases slowly to 15 per half-whorl and the density of secondaries increases to 2.5-3.5 per primary.

The length of the body chamber is unknown, because complete specimens with aperture have not been found; but is in excess of three-quarters of a whorl. Uncoiling to the serpenticone condition begins at about 100 mm diameter but the umbilical seam moves only slowly away from the line of tubercles on the previous whorl. The body chamber is relatively higher and rounder than the phragmocone whorls and becomes ovate in cross-section (H/W = 0.75). Ribbing remains strong and sharp to the aperture but the tubercles are reduced to low, laterally elongate swellings on the flanks. Secondary ribs remain as strong as the primaries but decrease in density until there are only 2-2.5 per primary. There is also a loss of bifurcation, most secondaries arising

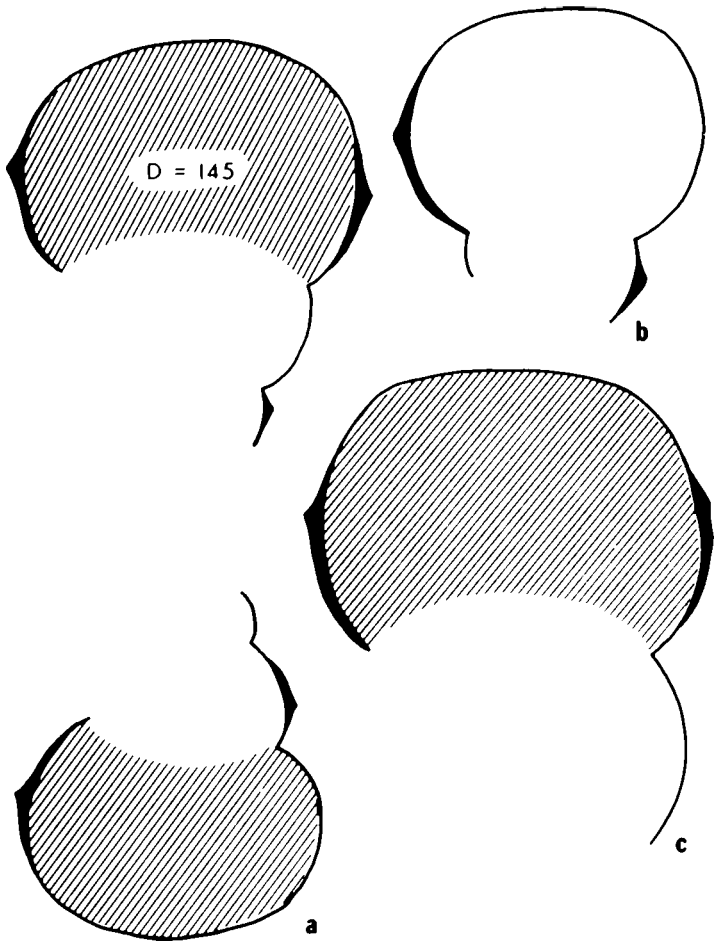


Text-figure 18. — Bivariate plots of whorl width (W), whorl height (H), umbilical diameter (U), shell diameter (D) and number of primary (P) and secondary (S) ribs per half-whorl, for *Stephanoceras skidegatense* ♀ and ♂ from Queen Charlotte Islands. Based on 12 macroconchs and three microconchs, with several individual "growth lines." Insets are same graphs showing approximate positions of best-fit mass curves. See *Measurements* for additional explanations.

by intercalation and lacking any connection with the tubercles; it is not uncommon for simple ribs to continue over the venter.

One body chamber fragment collected by Hall shows a small part of the apertural margin, which consists of a simple, flared lip preceded on the flank only by a broad and very shallow depression.

Remarks. — In whorl dimensions and ribbing this macroconch is closely allied with *S. itinsae* ♀ from the South Balch Island locality. However, the decline in the density of secondary ribbing and tubercle strength and the persistence of strong, sharp secondary ribs on the body



Text-figure 19. — Cross-sections of *Stephanoceras skidegatense* ♀, with body chamber shaded, from Queen Charlotte Islands. × 0.7, a, "var." *laperousii* (holotype, GSC 6482); b, holotype (GSC 5011); c, McM J1878. See *Measurements* for additional explanations.

chamber of *S. skidegatense* ♀ allow the two species to be separated at maturity. They do not co-occur at any locality.

Stephanoceras skidegatense ♂ (microconch)

Plate 9, figures 2-3; Text-figures 18, 20, 21

1884. *Ammonites Skidegatense;* Whiteaves, p. 210 [GSC 56694 here figured].

(This dimorph has not been previously described.)

Allotype. — McM J1802b (Pl. 9, fig. 3) from Richardson Bay on the south shore of Maude Island, Skidegate Inlet, Queen Charlotte Islands (Text-fig. 1); approximately 20 m above the base of the Yakoun Formation there exposed.

Material. — In addition to the allotype, eight other microconchs (McM J1802a,c-g, 1805, 1803) were collected from the same locality, four of them complete with aperture; the well-preserved specimen with 3/4 whorl body chamber (GSC 56694) was bought by Dawson from the Indians of Queen Charlotte Islands (Whiteaves, 1884, p. 210).

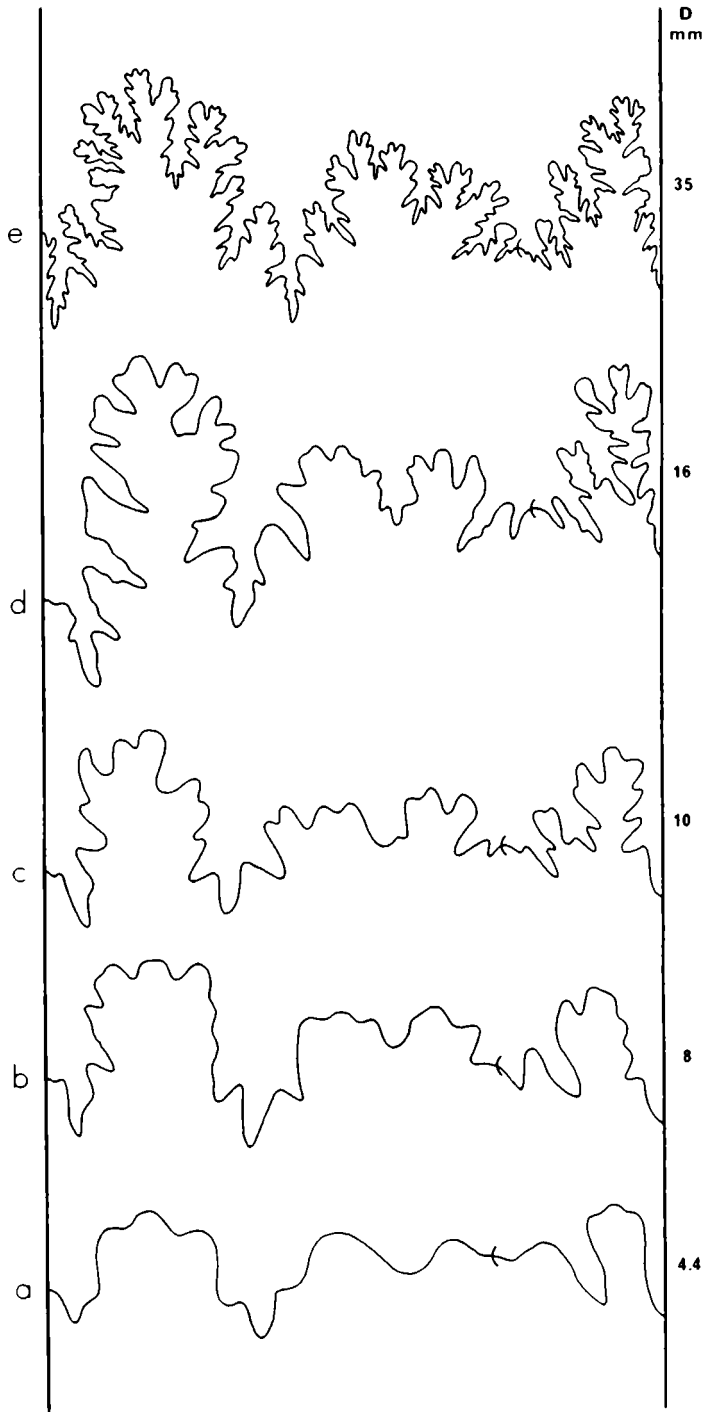
Description. — This is the largest microconch stephanoceratid from any of the Queen Charlotte Islands localities, attaining maximum diameters of 60-65 mm with robust ornamentation and only minor uncoiling on the last quarter-whorl.

At a diameter of 4 mm the whorl cross-section is coronate with a smooth shell, broadly arched venter, sharp lateral shoulder and slightly convex, moderately steep flanks. The only ornamentation at this stage consists of small, rounded tubercles on the lateral shoulder and extending adapicad as faint undulations about halfway down the flanks (incipient primary ribs). By a diameter of 6 mm faint secondary ribs appear, two to each primary, curving forward over the broad venter. Primary ribs are still seen only as faint extensions of the tubercles onto the upper parts of the flanks. Earliest whorls are much wider than high with H/W ratios of 0.55-0.65.

Primary ribs become strong by D = 10 mm, extending to the umbilical seam and curving forward with prominent, sharp tubercles on the lateral shoulder. Secondary ribs are broad and strong, curving forward from the tubercles and then crossing straight over the venter. Throughout the remainder of the phragmocone the primary ribs are stronger than the secondaries, conspicuously curved, reaching maximum height in the middle of the flanks and terminating in prominent tubercles at 43 to 53 percent of the whorl height. The number of primary ribs increases gradually from nine to 11 per half-whorl while the ratio of secondary to primary ribs increases from 2.5 to 3 by the end of the phragmocone. Whorl cross-section remains coronate with H/W ratios of 0.60-0.65, a broad venter and convex flanks (Text-figs. 18, 21). There is no distinct umbilical wall.

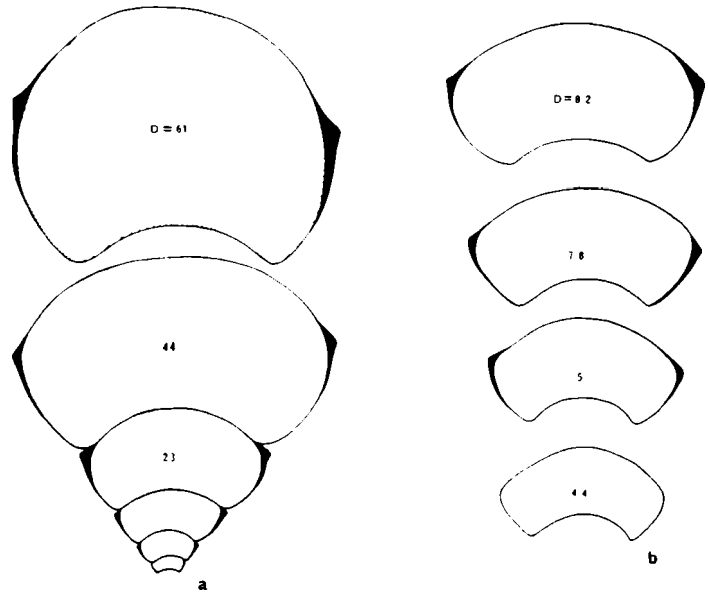
The body chamber is just over one-half whorl in length, marked only by slight uncoiling of the last quarter-whorl. The whorl cross-section is oval and relatively higher than on the phragmocone whorls (H/W = 0.75-0.85). Both primary and secondary ribbing remain strong to the aperture with 14 to 15 primaries on the last half-whorl but only 2-2.5 secondaries to each primary. The decline of secondary rib density is particularly noticeable on the last quarter-whorl where each primary rib bifurcates and intercalated ribs are absent. Tubercles become smaller, sharper and laterally elongate.

The aperture, not preceded by any constriction, is marked by a slightly flared lip; the beginning of the lateral lappets is visible on the four nearly complete specimens.



Text-figure 20. — Incomplete sutural ontogeny for *Stephanoceras skidegatense* ♂, from Queen Charlotte Islands (McM J1802a). See *Measurements* for additional explanations.

The mature suture is only moderately incised (Text-fig. 20). L is trifid and not as deep as E while U_2 is very short. U_3 is long, narrow and strongly oblique. E/L is higher than L/U_2; the tubercle is situated on the ventral side of the L/U_2 saddle.



Text-figure 21. — Cross-sections of two specimens of *Stephanoceras skidegatense* ♂, with body chamber shaded, from Queen Charlotte Islands, × 1.6 (a, McM J1802b; b, McM J1802a). See *Measurements* for additional explanations.

Remarks. — The decline to strong, bifurcating ribbing on the body chamber is similar to the body chamber ornamentation of *Zemistephanus* ♂. However, *S. skidegatense* ♂ reaches larger diameters, has less steep flanks, sharper ribbing, tubercles on the body chamber, more primary ribs on the last half-whorl (14-15 compared with 9-12 for *Zemistephanus* ♂) and a relatively higher whorl cross-section (H/W = 0.75-0.85 vs. 0.60-0.70).

S. itinsae (McLearn) ♂ differs strongly in having a higher density of secondary ribs on the body chamber (3 per primary) but fewer primary ribs per half-whorl (8-10 vs. 14-15).

S. skidegatense (Whiteaves) ♂ resembles '*Normannites*' *orbignyi* Buckman in coiling, whorl proportions and ornamentation. The holotype of '*N.*' *orbignyi* differs in rib density, having only 12 primaries on the last half-whorl; however, other specimens from Dorset [in the collections of the Geological Survey, London and the British Museum (Natural History)] show denser ribbing on the body chamber. The strong, curved primaries, sharp tubercles that decline on the body chamber and the strong, bifurcating secondary ribs all agree closely with the material here described as *S. skidegatense* (Whiteaves) ♂.

Dimorphism. — The specimens described here for the first time as *S. skidegatense* ♂ are morphologically identical throughout the phragmocone whorls with the previously known macroconch specimens from Richardson Bay (Text-fig. 18). Whorl shape and dimensions, density and form of ribbing and changes in ornamentation during ontogeny are closely similar in both dimorphs. Body chamber modifications are also similar in both (excluding apertural modifications), which is unusual for the stephanoceratids described here.

Slight egression of the last part of the body chamber is accompanied by rounding of the whorl section, decrease in the density of secondary ribs (2-2.5 per primary) and changes in the nature of the tubercles. The persistence of strong secondary ribbing right to the aperture enables separation of this species from the closely-related *S. itinsae.*

Adult macroconch specimens are about 3 times the size of the largest microconchs (60-65 mm); in all collections so far made from the Richardson Bay locality they occur in approximately equal numbers.

Stephanoceras sp. ♂ aff. **S. skidegatense** (Whiteaves, 1876) ♂

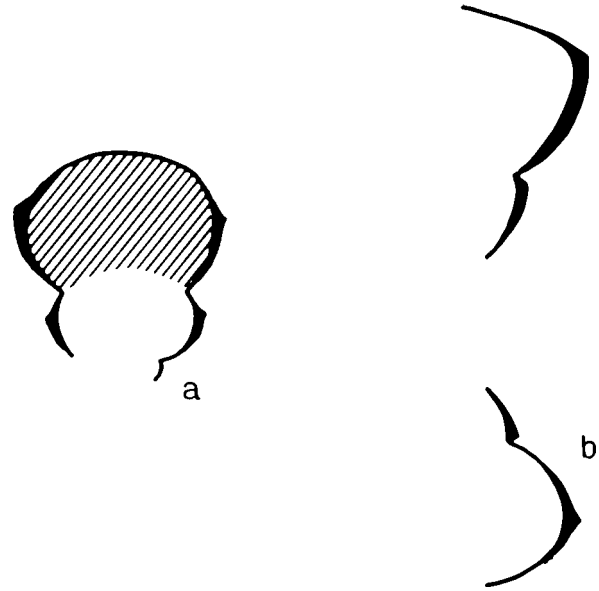
Pl. 10, figures 2-3; Text-figure 22

Material. — Two partially preserved specimens from the lower part of the Yakoun Formation at Richardson Bay, Maude Island, Queen Charlotte Islands: McM J1804, half of one body chamber whorl with one half-whorl of the preceding phragmocone attached, from 9 m above the base of the exposed section; and McM J1806, a smaller phragmocone found 20 m above the base of the section.

Description. — The phragmocone whorls are depressed (H/W = 0.63-0.72) with a strongly arched venter that rounds evenly onto the inflated flanks, there being no lateral shoulder (Text-fig. 22). The flanks slope gradually to the umbilical seam. Primary ribs are relatively short, reaching only 36 to 42 percent of the whorl height. At their ends a rounded, elongate swelling is developed and bifurcation occurs. The primary ribs are strong, rounded and curved forward; there are nine to 12 per half-whorl. Secondary ribs are also strong, curving slightly forward from the point of bifurcation and then crossing straight over the venter; there are 26 per half-whorl (*i.e.,* 2.4 to each primary).

The half-whorl of body chamber preserved shows gradual egression but the whorl section and ornamentation show little change. There are 14 primary ribs on the last half-whorl with 28 secondaries, all of which remain strong and rounded.

Remarks. — This species is distinguished from *S. skidegatense* (Whiteaves) ♂ by its stronger, more widely spaced secondary ribbing on the phragmocone whorls (compare Pl. 10, figs. 2-3 with Pl. 10, figs. 1b,c), more rounded whorl cross-section with highly arched venter, but lacks a lateral shoulder and has shorter, thicker primary ribs without sharp tubercles.



Text-figure 22. — Cross-sections with body chamber shaded, × 1. a, *Stephanoceras* aff. *skidegatense* ♂, (McM J1804); b, *Stephanoceras? (Stemmatoceras?)* ex gr. *S. acuticostatum* ["*Teloceras itinsae* McLearn"] (McM J1861). Both from Queen Charlotte Islands.

S. itinsae (McLearn) ♂ differs in the higher density of secondary ribbing that persists almost to the aperture, and the more coronate whorl section with a clearly defined lateral shoulder and broad venter on the phragmocone whorls.

Stephanoceras (Stephanoceras) pyritosum (Quenstedt, 1886-1887) ♀

1886-1887. *Ammonites Humphriesianus pyritosus* Quenstedt, p. 536; pl. 66, fig. 4.

Stephanoceras (Stephanoceras) pyritosum caamanoi McLearn, 1930 ♀

Pl. 10, figures 4a, b

1930. *Stephanoceras caamanoi* McLearn, p. 55; pl. 3, fig. 7; pl. 4, fig. 8.

?1969. *Stephanoceras* cf. *S. caamanoi;* Frebold *et al.,* p. 26, pl. 4, fig. 3.

The holotype came from South Balch Island in Skidegate Inlet, Queen Charlotte Islands, *i.e.,* the type locality of *Stephanoceras 'yakounense'* McLearn [= *S. itinsae* ♀]. The type is a large phragmocone (D = 110 mm) with typically

planulate, subcircular and evolute whorls, bearing sharp and dense, not prominent ribbing; small tubercles occur slightly below the middle of the whorls, at the end of the somewhat curved and projected primaries.

This specimen closely resembles the holotype of *S. pyritosum* (of which we possess a plaster cast) and is considered conspecific with the widely distributed European and Andean species (cf. Morton, 1971; Westermann and Riccardi, 1979). The only difference seems to be the longer primaries of *S. p. caamanoi*. Both resemble *S. humphriesianum* (Sowerby), the type-species, which differs in the coarser and stiffer ornament. This species clearly indicates the lower to middle *S. humphriesianum* Chronozone.

Subgenus **STEMMATOCERAS** Mascke, 1907

Type species. — *Am. Humphriesianus coronatus* Quenstedt, 1886-1887 [= *S. frechi* Renz, 1913; = "*Cadomites quenstedti*" Roché, 1938, obj. syn.]

Discussion. — Most North American species placed originally in *Stemmatoceras* are here transferred to the nominate subgenus *Stephanoceras*, i.e. '*Stemmatoceras*' *mclearni* Warren, '*S.*' *carri* Warren and (?) '*S.*' *arcicostum* Imlay, while '*S.*' *albertense* McLearn and '*S.*' *palliseri* McLearn are tentatively retained in *S. (Stemmatoceras)*; and '*S.*' *ursinum* Imlay is more tentatively transferred to *Zemistephanus*. Most of these species belong to the species group of *Stephanoceras umbilicum* (Quenstedt), characterized by broad subelliptical whorls with rounded lateral shoulders, and bear the coarse primary ribs of *S. brodiaei* (Sowerby).

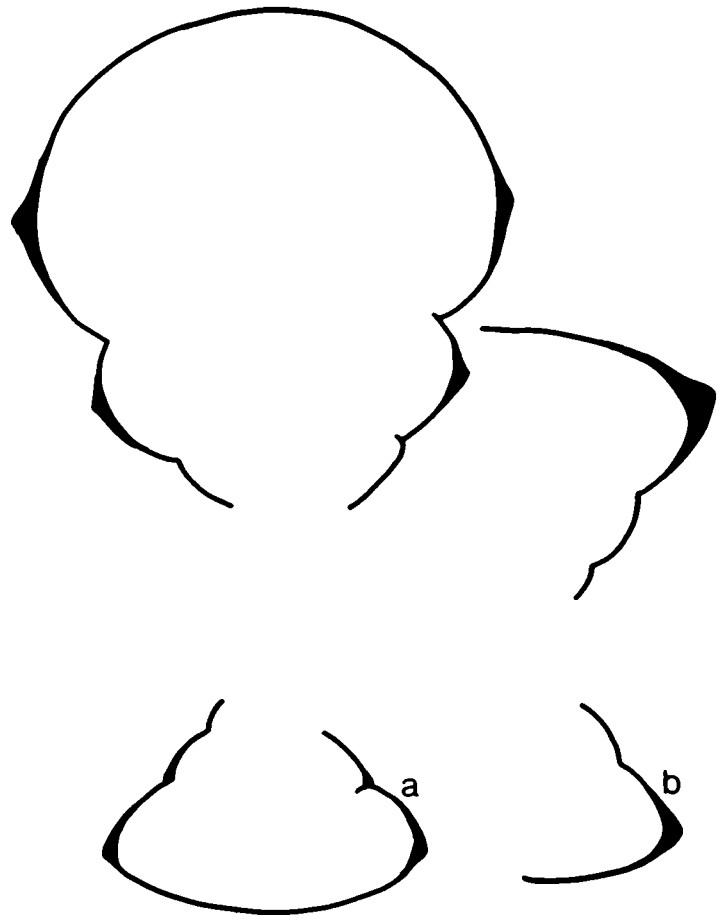
Typical *S. (Stemmatoceras)* from North America were previously placed in *Teloceras*, i.e. '*T.*' *dowlingi* McLearn, '*T*' *allani* Warren and probably "*T. itinsae*" McLearn [not '*Itinsaites*' *itinsae* McLearn].

Stephanoceras (Stemmatoceras) dowlingi (McLearn, 1930)

Plate 11, figures 2a, b; Text-figure 23a

1930. *Teloceras dowlingi* McLearn, p. 2.
1932b. *Teloceras dowlingi;* McLearn, p. 112, pl. 1; pl. 5, figs. 2, 3.

Remarks. — This species was based on an incomplete specimen, probably from Ribbon Creek in Alberta, that is wholly septate except for about one quarter-whorl of body chamber. In the first full description, McLearn (1932b, pp. 111-113) noted the essential deviations from *Teloceras* Mascke: the marked egression of the umbilical seam from the line of nodes, the high arching of the venter and rounding of the previously steep umbilical walls, and the marked decrease in the strength of the nodes. On the body chamber H/W increases to 0.71, far greater than any known values for undoubted species of *Teloceras* from western Europe



Text-figure 23. — Cross-sections through the phragmocones, × 1. a, '*Teloceras*' *dowlingi* McLearn, holotype (GSC 9050); b, "*Teloceras itinsae* McLearn" (GSC 6481), holotype.

(values of 0.40-0.50 are common at similar growth stages). In fact, whorl section and coiling closely resemble those of *S. frechi* Renz, type species of the subgenus *Stemmatoceras;* the only difference is in the more widely spaced, blunter primaries ending in more prominent tubercles.

Stephanoceras (Stemmatoceras?) allani (Warren, 1947)

Plate 12, figure 2

1947. *Teloceras allani* Warren, p. 70, pl. II, fig. 1.

Remarks. — The holotype from the Highwood-Sheep River area in Alberta is by far the largest of this group (D = 277 mm) and has almost one whorl of body chamber preserved. The inner whorls are tightly coiled with the umbilical seam following the line of nodes; these are small, round and situated on the pronounced lateral shoulder; the umbilical wall is steep producing a deep and narrow umbilicus. Uncoiling and rounding commence about half a whorl

before the end of the phragmocone. The umbilical wall becomes rounded merging more gradually with the arched venter (H/W increases from 0.64 to 0.84). The strong primary ribs that on the inner whorls are conspicuously directed forward as they approach the umbilical seam, become blunt on the body chamber and obsolete near the end; nodes and secondary ribs also decline on the body chamber. This species is also transferred to *Stephanoceras* and somewhat tentatively placed in the subgenus *Stemmatoceras;* confirmation awaits data from the section of the inner whorls.

Stephanoceras (Stemmatoceras?) albertense (McLearn, 1928) ♀

Plate 13, figure 1

1928. *Stemmatoceras albertense* McLearn, p. 20, pls. V-VII.

Remarks. — The holotype is a fully septate, distorted specimen from Sheep Creek, Alberta, on which only one side of the phragmocone is preserved; most of the ventral parts of the ultimate whorl have been crushed. In dimensions of the outer whorls and the long, curved primary ribs (18-20 per half-whorl) this specimen resembles *Stephanoceras* (*s.s.*) *itinsae* (McLearn) ♀, but *S. albertense* appears to have more coronate inner whorls. It may differ from '*Stemmatoceras*' *palliseri* McLearn in whorl section and coiling as suggested by Warren (1947, p. 68), and also has more primary ribs (18 vs. 13 per half-whorl at similar diameters). The relatively undistorted specimen from Ribbon Creek described and figured by Frebold (1957, p. 50; pl. 11, figs. 2a, b; pl. 23, figs. 1a-c) as "*Stemmatoceras albertense* McLearn" is identical with *Stephanoceras itinsae* (McLearn) ♀ from the Queen Charlotte Islands. The inner whorls as measured by Frebold (p. 51) correspond in all dimensions with *S. itinsae* and the increase in whorl height on the last preserved whorl and fine, dense secondary ribbing strongly confirm identification with that species.

The inner whorls of the holotype, however, are poorly known so that its true affinity remains uncertain.

Stephanoceras (Stemmatoceras?) palliseri (McLearn, 1930) ♀

Plate 12, figure 1

1930. *Stemmatoceras palliseri* McLearn, p. 3.
1932b. *Stemmatoceras palliseri;* McLearn, p. 114, pl. 2; pl. 5, fig. 1.

Remarks. — The holotype from Mountain Park, Alberta, is a wholly septate specimen, somewhat distorted and corroded. The inner whorls seem to be broad and coronate with moderately rounded lateral shoulders, curved lower flanks grading into the steep umbilical slope and a moderately deep umbilicus. The last phragmocone whorl, however, is depressed elliptical with well-rounded lateral shoulder and arched venter. The primaries are prominent. The "*Stemmatoceras* cf. *S. palliseri* McLearn" described by Imlay from southern Alaska (1964, pp. B48, 49; pl. 20, figs. 5, 6; pl. 21, figs. 2, 4) also has similar dimensions and coiling with a corresponding decline in ornament on the body chamber. However, the Alaskan specimen has fewer primary ribs (12 vs. 16-19 per half-whorl at D = 130 mm) with larger nodes. Secondary ribs are more dense on the inner whorls of Imlay's specimen but become broader and less dense (three secondaries per primary) on the body chamber. Since the section of the inner whorls is unknown, the subgeneric assignment of *S. palliseri* is uncertain. There is some resemblance to '*S.*' *albertense* (see above).

Stephanoceras? (Stemmatoceras?) ex gr. **S. acuticostatum** (Weisert, 1932) ♀

Plate 5, figures 3-4; Text-figure 23b

1932a. *Teloceras itinsae* McLearn, p. 51, pl. 10, figs. 1, 2.
1964. *Teloceras itinsae*; Imlay, p. B50, pl. 24, figs. 3, 4 (lectotype refigured), *non* figs. 1, 2, 5, 7 (= *Zemistephanus alaskensis* sp. nov.)

Material. — Three poorly preserved specimens from the Yakoun Formation at McKenzie Bay, Queen Charlotte Islands (type locality of "*Teloceras itinsae*"): a crushed phragmocone of 80 mm diameter (McM J1861) and a small fragment (McM J1860) from 10-15 m above the base of exposed formation and one fragmentary specimen with parts of the last two phragmocone whorls and of the body chamber (GSC MB858), collected by Sutherland Brown at the same general locality.

Description. — The holotype from McKenzie Bay is a well-preserved, possibly incomplete phragmocone of 70 mm diameter (Table 6). The phragmocone whorls are moderately depressed and not very wide (W/D $\cong$ 0.5) but subcoronate, with a narrowly curved lateral edge bearing prominent thick tubercles, and only gently curved inner flanks (with umbilical slope) and venter. The ornament consists of widely spaced, prominent primaries ending in the tubercles at about 3/5 whorl height, and coarse secondaries, three to four per primary; both primaries and secondaries are slightly projected (much of the curvature probably a result of the slight distortion). On the last preserved phragmocone whorl, the primaries withdraw somewhat from the umbilical seam and start to become blunt. The septal suture resembles that of typical stephanoceratids, with a much smaller L/U_2 than E/L saddle and strongly retracted umbilical elements.

While the smaller topotype (J1861) closely resembles the holotype except for the poorer preservation, the larger

one (GSC MB 858) has a phragmocone diameter of about 100 mm and parts of the body chamber which reached 130-150 mm diameter. The ultimate phragmocone whorl differs from that of the smaller holotype only in the blunter primaries that are further withdrawn from the seam. Unfortunately, the specimen is preserved with only one side so that the exact width of the whorl is unknown (Text-fig. 23b). The body chamber becomes much more rounded in section, with strong negative allometry of height and (?)width growth, and probably becomes more evolute; the ornament becomes increasingly blunt. The incomplete preservation and partial distortion of the body chamber whorl, however, do not permit the complete reconstruction of the specimen, in particular with regard to the change in adult whorl section and coiling.

Remarks. — It appears that "*Teloceras itinsae* McLearn" closely resembles the European "subcoronates" of the '*Teloceras*' *acuticostatum* group. The species group differs from true *Teloceras* as redefined herein in the narrow phragmocone whorls and the rounded outer whorl(s), and is placed in *Stephanoceras (Stemmatoceras)*. This species group includes the similar, poorly known '*T.*' *subblagdeni* Schmidtill and Krumbeck and '*T.*' *blagdeniformis* Roché, and appears to mark the lower part of the *T. blagdeni* Subzone (Mouterde *et al.*, 1971). This species group is also represented in New Guinea (cf. Boehm, 1913, pl. 3, fig. 2). The incomplete preservation of the holotype and topotypes of '*T. itinsae*', however, makes definite comparison impossible.

If transferred to *Stephanoceras*, however, '*Teloceras itinsae*' McLearn, 1932a, becomes a junior homonym of '*Itinsaites*' *itinsae* McLearn, 1927, which is a true *Stephanoceras* (*i.e.*, the definite microconch of *S.* '*yakounense*' McLearn, 1930 ♀). '*Teloceras itinsae*' would therefore need a new name (epithet), unless it can be demonstrated to be conspecific (and synonymous) with another described species. Since '*Teloceras itinsae*' and the European '*T.*' *acuticostatum*, '*T.*' *subblagdeni* and '*T.*' *blagdeniformis* are known from poorly preserved material only, no clear specific distinctions are evident to us. We therefore abstain from replacing the homonymous epithet. There is a remarkable resemblance to *Zemistephanus alaskensis* sp. nov. with which the rare "*Teloceras itinsae*" may be associated. *Z. alaskensis* is distinguished by the low, wide-spaced primaries on the juvenile whorls and their obsolescence on the outer whorls.

"*Teloceras itinsae*" is of appreciable biostratigraphic and chronologic significance. The name has been used erroneously by Imlay (1964) for a faunule or assemblage zone in South Alaska, based on misidentified *Zemistephanus alaskensis* sp. nov. "*T. itinsae*" has now been located stratigraphically at its type locality (McKenzie Bay) where it occurs within the range zone of *Zemistephanus richardsoni*. This provides good evidence for correlation with the (lower to middle) *S. humphriesianum* Standard Zone (*D. romani* Subzone).

Table 6. — Measurements (in mm) of "*Teloceras itinsae*".

Holotype (GSC 6481)	D	W	H	U
phragmocone	71	36.5	22.5	32.5
"	44	22	14.8	17.9

Genus **TELOCERAS** Mascke, 1907

Type species. — *Am. blagdeni* J. Sowerby, 1818, by original designation.

Discussion. — *Teloceras* Mascke was originally distinguished from other stephanoceratids by its strong nodes, sharp umbilical shoulder and the great thickness of the whorls (Mascke, 1907, p. 31). The distinctiveness of this genus was further emphasized by Weisert (1932) and Schmidtill and Krumbeck (1938): although the complete body chamber was rarely preserved, the mature shell exhibited a diminishing sculpture, slight rounding of the whorl section and, in some forms, minor egression. Dr. J. Wiedmann kindly examined for us 40 large Swabian *Teloceras*, 3 of them with partial to complete body chamber, in the Geologisch-Paläontologisches Institut and Museum in Tübingen and one of us [Hall] has studied a number of almost complete specimens from the Inferior Oölite in the British Museum (Natural History). At least the outer whorls of the phragmocone and the adapical parts of the adult body chamber are coronate with strongly depressed, broad whorls and flat venter with large, round nodes along the acute lateral shoulder. Only the second half of the body chamber rounds markedly. In many specimens the venter is almost smooth as a result of the loss of secondary ribbing. In consultation with Dr. J. Callomon, *Teloceras* is distinguished at the genus level and defined to include as macroconchs only species with an extremely coronate, large adult stage.

A detailed discussion of the taxonomy and nomenclature of this genus is given under the Genus *Stephanoceras*.

Most North American species described under *Teloceras*, however, are here transferred either to *Stephanoceras (Stemmatoceras)*, i.e. '*T.*' *dowlingi* McLearn, '*T.*' *allani* Warren and '*T.*' *itinsae* McLearn [becoming a junior homonym of *S. itinsae* (McLearn)], or to *Zemistephanus*, i.e. '*T.*' *warreni* McLearn.

Probably the only true *Teloceras* described from North

America are the poorly preserved holotypes of *'Teloceras stelcki'* Warren and *'Zemistephanus' crickmayi* Frebold, both from the Rock Creek Member of southern Alberta. These may be conspecific.

'T. stelcki' Warren (1947, p. 71, pl. VI, fig. 1) was based on a large (D = 232 mm) specimen with body chamber; but the right side and the inner whorls are missing and at least the body chamber is somewhat crushed. The preserved last whorl of the phragmocone (D ≅ 140 mm) is typically coronate with a sharp lateral edge bearing very prominent subconical tubercles. The primaries are strongly reduced to obsolete while the secondaries are comparatively dense, fine and projected (?due to distortion). The large body chamber (3/4 whorl; D ≅ 230 mm) displays strong egression and rounding above 170 mm diameter, but this could be in part a result of distortion. In the absence of the inner whorls, it cannot be decided whether this is an unusual *Teloceras* or an exceptionally large coronate *Zemistephanus*. The type material of this "species" is considered unsatisfactory and the name therefore is a *nomen dubium*.

Teloceras crickmayi (Frebold, 1957) ♀

1957. *Zemistephanus crickmayi* Frebold, p. 52, pl. XXV, fig. 1; pl. XXVI, fig. 1; pl. XXVII, fig. 1.

The fully septate and fragmentary holotype from Ribbon Creek, Alberta (reconstructed phragmocone D ≅ 150-160 mm) is characterized by a deep, conical umbilicus, extremely broad and depressed cross-section, with only slightly arched venter and high, conical tubercles. Primary ribs are obsolete; from the tubercles situated on the acute lateral shoulder the relatively fine secondary ribs pass slightly projecting over the venter. There are about 11 tubercles per half-whorl and four secondary ribs to each tubercle on the last two whorls. Though there is slight egression of the last preserved whorl from the line of tubercles, there is no decline in the steepness of the umbilical wall, nor rounding of the whorl section, as in *Zemistephanus* (H/W = 0.53 → 0.50). One of us [Westermann] has collected several topotypes, including one specimen with body chamber. Rounding of the whorl section occurs only with the second half of the body chamber, much as in *Teloceras* of the *T. blagdeni* group. This species is distinguished from *T. blagdeni* in the reduction of the primaries, similar to *T. banksii* (Sowerby), and in the denser secondaries. It could be synonymous with the poorly known *'T. stelcki'* Warren (*nomen dubium*) and is transferred to *Teloceras*. This species is of stratigraphic and chronologic significance, because it remains the only *Teloceras* in what has been called the "*Teloceras* fauna" (Warren, 1934) and indicates the *T. blagdeni* Subzone.

Family **SPHAEROCERATIDAE** Buckman, 1920

Genus **CHONDROCERAS** Mascke, 1907

Type species. — *Am. gervillii* J. Sowerby, 1817 (by original designation).

Discussion. — McLearn (1927) erected two new genera for Bajocian sphaerocone ammonites from western Canada: *Defonticeras*, occurring in the Queen Charlotte Islands and *Saxitoniceras*, from Alberta. He later (1949) came to regard *Defonticeras* as a subgenus of the European *Chondroceras* Mascke.

'Defonticeras' McLearn was separated from the then mainly European genus *Chondroceras* by the latter having a "3-ridged mouth, somewhat regular umbilicus, and fine ribs sloping well forward near the anterior end of the ultimate whorl" (McLearn, 1929, p. 13). Comparison of larger collections of *'Defonticeras'* from the Queen Charlotte Islands, Alberta and southern Alaska with plastotypes and figured specimens of European *Chondroceras* shows that in both groups sudden umbilical enlargement begins about two-thirds to one-half of a whorl before the aperture, corresponding to the beginning of the body chamber. Though a three-ridged mouth border is never seen in the eastern Pacific material, it is by no means universal in European species of *Chondroceras* either (see Westermann, 1956, pls. 1-3). Primary ribs on the body chamber of eastern Pacific species are coarser and more widely spaced than on European *Chondroceras* but it is believed this minor variation in ornamentation of the body chamber is insufficient to warrant separation, even at the subgeneric level, and *'Defonticeras'* McLearn is here treated as synonymous with *Chondroceras* Mascke. This follows the classification adopted in the 'Treatise' (Arkell, 1957, p. L292).

The two species of *'Saxitoniceras'* were separated from *'Defonticeras'* by their less dissected suture line and less abrupt umbilical enlargement (McLearn, 1927, 1928). Again, these minor variations are not considered sufficient for generic or subgeneric distinction. Indeed, *Chondroceras oblatum* (Whiteaves) ♀ from the Queen Charlotte Islands, tentatively placed in *'Defonticeras'* by McLearn (1929, p. 17) has a similarly simplified suture line (compare Text-fig. 26) and is regarded as conspecific with *Saxitoniceras marshalli* McLearn. *'Saxitoniceras'* McLearn is also placed in synonymy with *Chondroceras* Mascke. Umbilical enlargement in *'Saxitoniceras'* occurs suddenly over the last half- to quarter-whorl and is similar in character to that on *'Defonticeras'*.

Westermann (1956) treated *'Defonticeras'* and *'Saxitoniceras'* as subgenera of *Chondroceras;* he later (1964a,

pp. 55, 64, 65) tentatively placed *'Saxitoniceras'* in synonymy with *'Defonticeras'* as the corresponding microconch form. However, no *'Saxitoniceras'* has been found with the macroconch *'Defonticeras'* faunas on the Queen Charlotte Islands and indeed the corresponding microconchs are much smaller than *'Saxitoniceras'* (see Pl. 14, fig. 8; Pl. 15, fig. 6). The two known species of *Chondroceras* from the Queen Charlotte Islands are strongly dimorphic in that the macroconch is approximately three times the size of the corresponding microconch. However, the apertural modifications in both dimorphs are similar: a constriction followed by a flared collar.

Chondroceras oblatum (Whiteaves, 1876) ♀ & ♂

Pl. 14, figures 1-8; Text-figures 24-26

Chondroceras oblatum ♀ (macroconch)

1876. *Ammonites loganianus* Form A Whiteaves, p. 29; pl. 4, figs. 2, 2a.
1884. *Stephanoceras oblatum* Whiteaves, p. 209.
1927. *Saxitoniceras marshalli* McLearn, p. 68.
1928. *Saxitoniceras marshalli* McLearn, p. 22; pl. VIII, figs. 3, 4.
1929. *Defonticeras oblatum* (Whiteaves); McLearn, pp. 16, 17, pl. XV, fig. 1.
1956. *Chondroceras (Defonticeras) oblatum* (Whiteaves); Westermann, pp. 102-4; figs. 18, 61; pl. 11, figs. 4, 5.
1957. *Chondroceras marshalli* McLearn var.; Frebold, p. 54; pl. XXV, figs. 3a, b; pl. XXVI, figs. 2a, b.
1964a. *Chondroceras (Defonticeras) oblatum* (Whiteaves); Westermann, pp. 55, 64.
1964. *Chondroceras* cf. *C. marshalli* (McLearn); Imlay, p. B43, pl. 12, figs. 1-3.

Holotype. — GSC 4964, on Pl. 14, figs. 1a, b collected by J. Richardson in 1872 and labelled "Skidegate Channel". McLearn (1929, p. 17) assumed this specimen came from the lower part of the Yakoun Formation at Richardson Bay on the south shore of Maude Island. However, it does not resemble any of the specimens collected and described from that locality; it does agree closely with material from South Balch Island in Skidegate Inlet and the holotype may well have come from the lower Yakoun Formation exposed at this locality.

The holotype of *'Saxitoniceras' marshalli* was originally described as from near the base of the Fernie Formation on the head waters of Sheep Creek, Alberta (McLearn, 1928, p. 22).

Other Material. — Eight macroconch specimens (McM J1795, 1832a-d, 1833 and 1834a, b) were obtained from several horizons in the lower Yakoun Formation of South Balch Island. Other specimens described as *'Saxitoniceras' marshalli* came from the Rock Creek Member of the Fernie Group in Ribbon Creek, southern Alberta (Frebold, 1957), and the Tuxedni Formation of southern Alaska (Imlay, 1964). Further material from the Ribbon Creek locality in southern Alberta collected by Westermann (McM J1837a-f) and Hall (McM J1836a-f) was also studied.

Description. — The early phragmocone whorls are sphaeroconic with a rounded venter curving continuously onto the convex flanks and to the umbilical seam. The umbilicus is very narrow, usually representing less than ten percent of the shell diameter. Whorls depressed, with H/W = 0.65-0.75. Up to a diameter of 13 mm the shell is smooth. At greater diameters on the phragmocone the primary ribs are faint on the umbilical wall, somewhat stronger but never sharp on the lower flanks, almost rectiradiate, eight to 10 per half-whorl. There are 3.5 secondary ribs to each primary rib and they pass straight over the venter; tubercles are absent at all stages of growth.

The body chamber is two-thirds of a whorl in length, egression beginning at the last septum with sudden enlargement of the umbilicus to as much as 30 percent of the shell diameter. Near the aperture the whorl section remains depressed (H/W = 0.70-0.75), the venter remains broadly rounded, but the flanks are slightly flattened. There are eight or nine primary ribs on the last half-whorl, that are slightly prorsiradiate on the flanks and give rise to three secondary ribs per primary. Maximum diameters attained are 55-60 mm, and the aperture is marked by a narrow constriction followed by a smooth lip.

Remarks. — This species is clearly differentiated from *C. defontii* ♀ and *Chondroceras* n. sp. indet. ♀ by ribbing density and sutural complexity. *C. defontii* ♀ is characterised by complex sutures that are deeply incised whereas the suture of *C. oblatum* ♀ has broader saddles with small incisions. While also having a simple suture, *Chondroceras* n. sp. indet. ♀ has much denser and finer ribbing (12-14 primaries per half-whorl) and a broader whorl section than *C. oblatum* ♀ (H/W = 0.47-0.55 vs. 0.65-0.75 at similar diameters).

Comparison of the whorl dimensions of *'S' marshalli* ♀ with *C. oblatum* ♀ shows that the two forms are similar (Text-fig. 24). They attain similar maximum diameters (58.7 mm for the holotype of *C. oblatum* ♀ and 51-66 mm for *'S.' marshalli* ♀), have simplified sutures (Text-fig. 26), depressed whorl sections (H/W = 0.60-0.80) and seven to ten primary ribs on the last half-whorl of the body chamber with three secondaries to each primary. Thus *C. oblatum* (Whiteaves) and *'S.' marshalli* McLearn are placed in synonymy, the former name taking precedence.

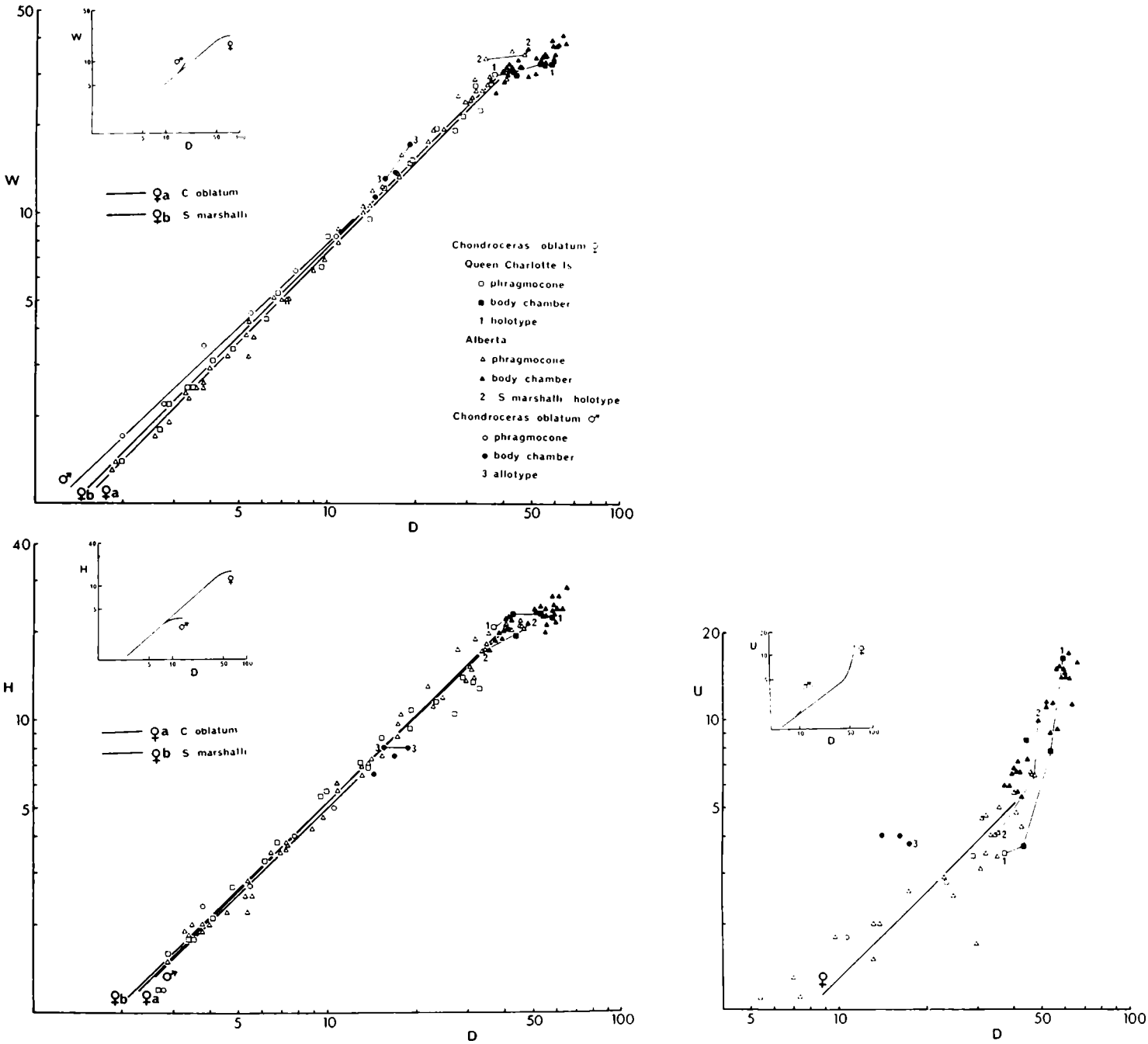
Chondroceras oblatum ♂ (microconch)

(This dimorph has not been previously described.)

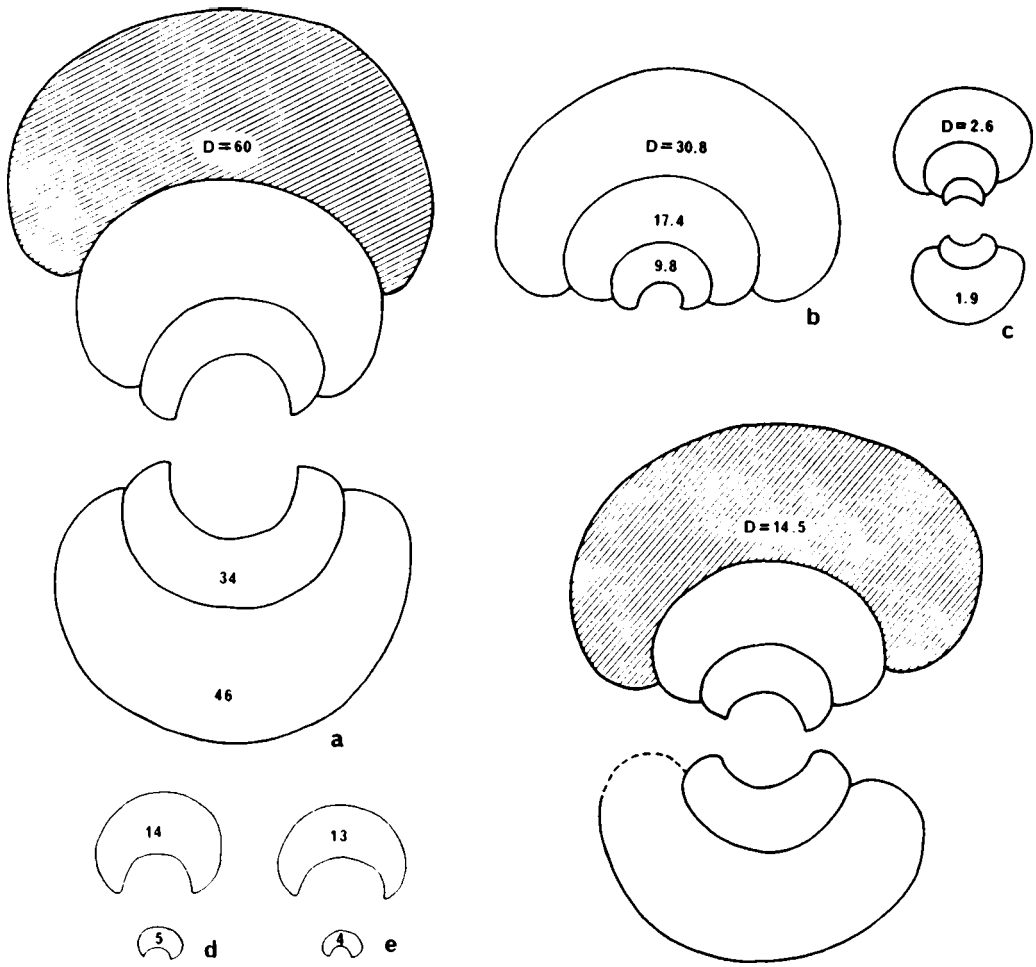
Allotype. — McM J1794a, on Pl. 14, fig. 8, from the lower part of the Yakoun Formation, eastern shore of South Balch Island, Queen Charlotte Islands.

Other Material. — One other complete specimen, McM J1835, from 44 m above the base of the exposed Yakoun Formation on South Balch Island.

Description. — Microconchs are characterised by egression of the last quarter-whorl, which terminates with a nar-



Text-figure 24. — Bivariate plots of whorl width (W), whorl height (H), and umbilical diameter (U) against shell diameter (D) for *Chondroceras oblatum* ♀, including the synonymous "*Saxitoniceras marshalli*", respectively from Queen Charlotte Islands and southern Alberta. Based on 22 macroconchs and three microconchs, with several individual "growth lines" indicated. Insets are same graphs showing positions of best-fit mass curves. See *Measurements* for additional explanations.



Text-figure 25. — Cross-sections of *Chondroceras oblatum* ♀ and ♂, body chamber shaded. a-c, *"Saxitoniceras marshalli"* ♀ from southern Alberta, a-b, × 1.6 (McM J1836e, McM J1837g); c, × 12 (McM J1836f); d-e, juvenile whorls of macroconch at diameters indicated, × 1.5 (McM J1795b, McM J1795a); f, microconch × 4.5 (McM J1794b). See *Measurements* for additional explanations.

row constriction followed by a smooth lip. Microconchs reach a maximum size of 19 mm, only one-third the size of the adult macroconchs. The shell is oblate with a broadly rounded venter, narrow umbilicus and convex flanks. Primary ribs are slightly prorsiradiate with seven on the last half-whorl; there are three secondaries to each primary rib at this stage. Dimensions and suture are similar to those of the macroconch (Text-figs. 24-26).

Dimorphism. — Text-figure 24 and Appendix 1 show that the whorl dimensions of the few known microconchs are very similar to those of the macroconchs. Ribbing style and density on the body chambers of the two dimorphs are similar, while the apertural modifications are identical: a narrow constriction followed by a broad, smooth collar. The macroconch is approximately three times the size of the microconch. No microconch specimens are known from the Ribbon Creek locality of southern Alberta.

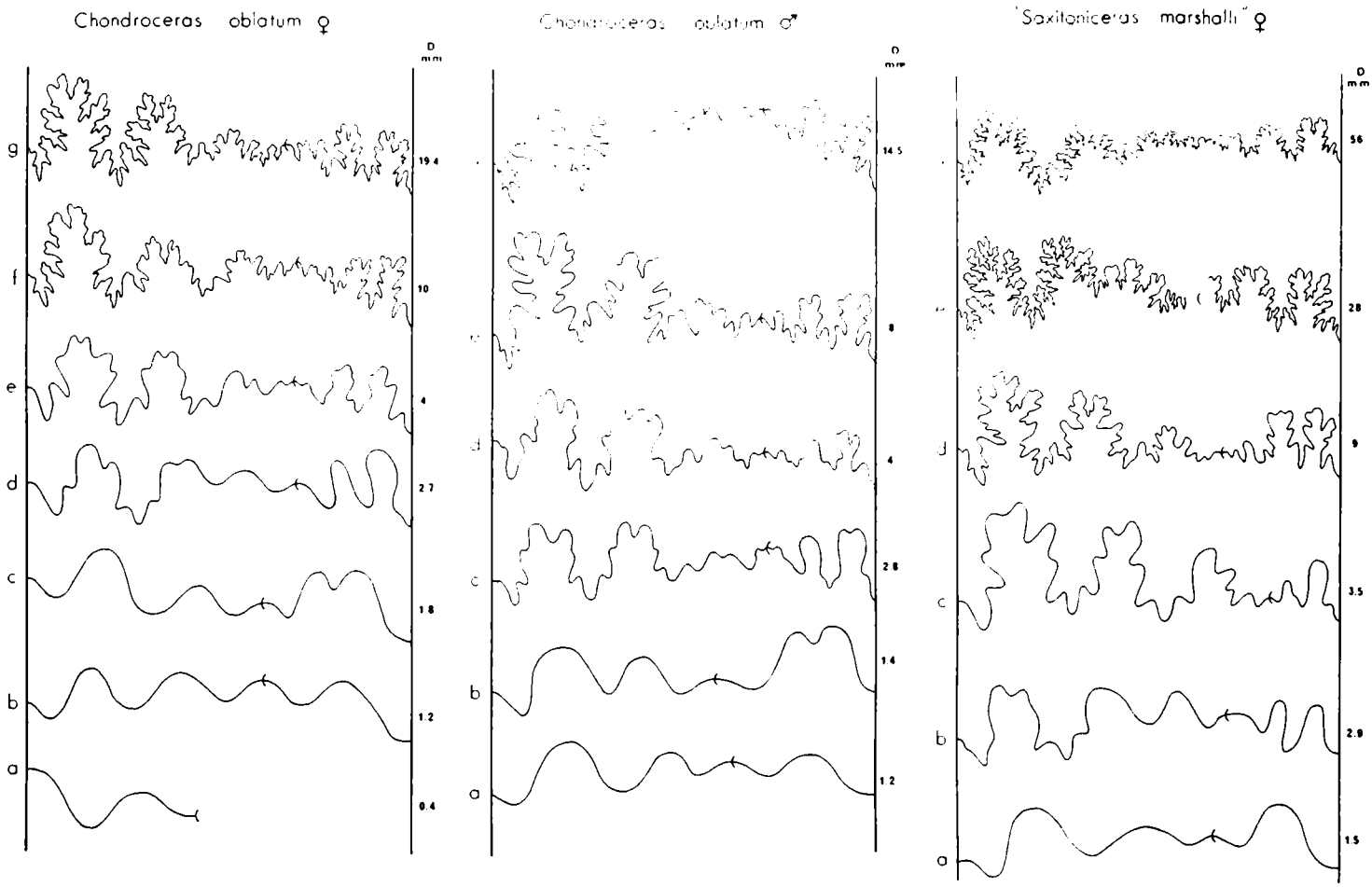
Westermann (1964a, p. 65) considered that *'Saxitoniceras' allani* was a possible microconch equivalent of *'Defonticeras' oblatum.* However, the microconchs here described and figured from the macroconch type locality on South Balch Island are very much smaller than either the macroconch *C. oblatum* or *'S.' allani;* no specimens similar to *'S.' allani* have been found on the Queen Charlotte Islands.

Chondroceras defontii (McLearn, 1927) ♀ & ♂

Plate 15, figures 1-6; Text-figures 27-29

Chondroceras defontii ♀ (macroconch)

1927. *Defonticeras defontii* McLearn, p. 72; pl. 1, fig. 3.
1929. *Defonticeras defontii* McLearn; McLearn, pp. 13, 14; pl. XII, figs. 1-3.
1929. *Defonticeras colnetti* McLearn, pp. 15, 16; pl. XIII, figs. 4, 5.



Text-figure 26. — Sutural ontogenies for *Chondroceras oblatum* ♀ (left) and ♂ (center), from Queen Charlotte Islands (McM J1795a, McM J1794b); right, "*Saxitoniceras marshalli*" from southern Alberta (McM J1837h, McM J1836d). The position of the umbilical seam is marked by (. See *Measurements* for additional explanations.

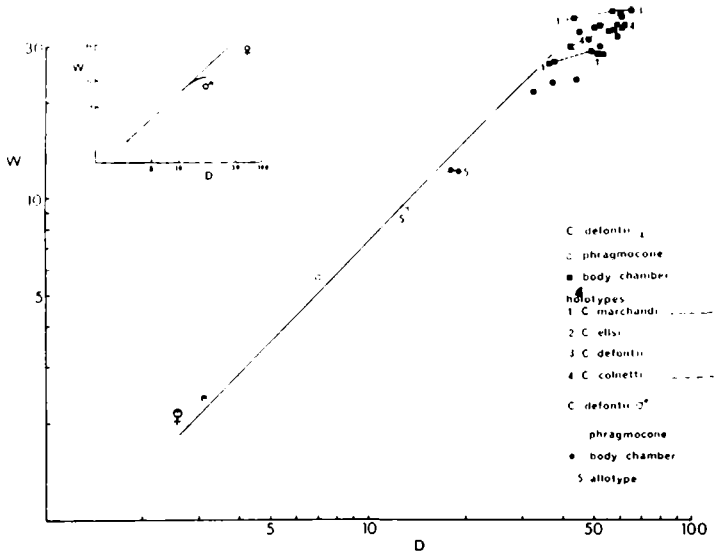
1929. *Defonticeras ellsi* McLearn, p. 16; pl. XIII, figs. 2, 3; pl. XIV, fig. 1.
1929. *Defonticeras marchandi* McLearn, pp. 14, 15; pl. XII, figs. 4, 5.
1949. *Chondroceras (Defonticeras) defontii* (McLearn); McLearn, pp. 10, 16.
1956. *Chondroceras (Defonticeras) defontii* (McLearn); Westermann, pp. 100-102; pl. 11, fig. 3; figs. 57, 59.
1964. *Chondroceras defontii* (McLearn); Imlay, p. B42; pl. 12, figs. 8, 11-14.

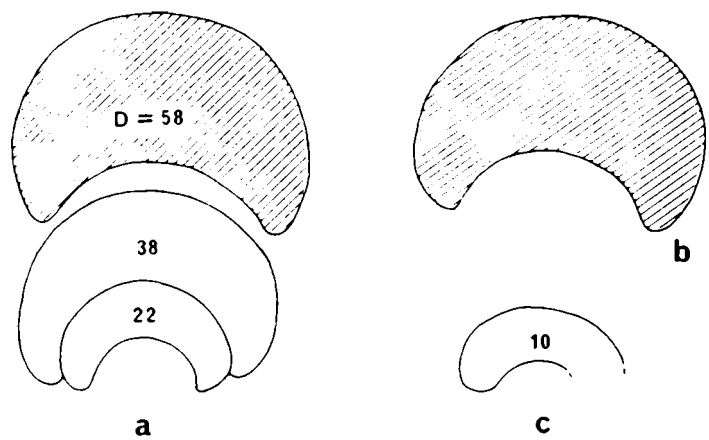
Holotype. — GSC 9009 on Plate 15, figures 1a,b, from talus on the ledges of the lower Yakoun Formation at Richardson Bay on the south shore of Maude Island, Queen Charlotte Islands.

Other Material. — Seven complete (some crushed) macroconchs (McM J1792a-f, J1829) and a number of body chamber fragments complete with aperture from the type locality.

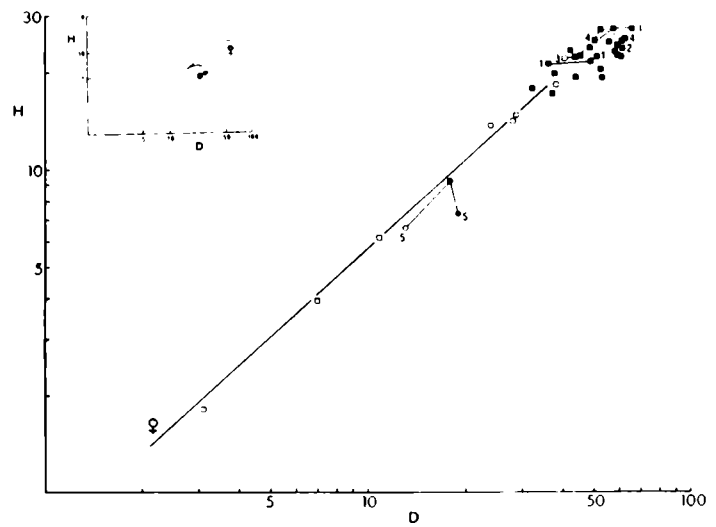
Description. — Phragmocone whorls globose with broadly rounded venter, inflated flanks and short, steep umbilical wall (Text-fig. 28). Umbilicus very narrow, usually less than 10% of the shell diameter, and deep. Cross-section depressed with H/W ratios of 0.60-0.75. Primary ribs fairly strong, rectiradiate on the steep umbilical wall, then strongly curved forward on the lower flanks; 11 to 15 per half-whorl. Secondary ribs fine and closely spaced, up to 3.5 per primary, curved forward from the point of furcation and then crossing straight over the venter. Nodes not developed.

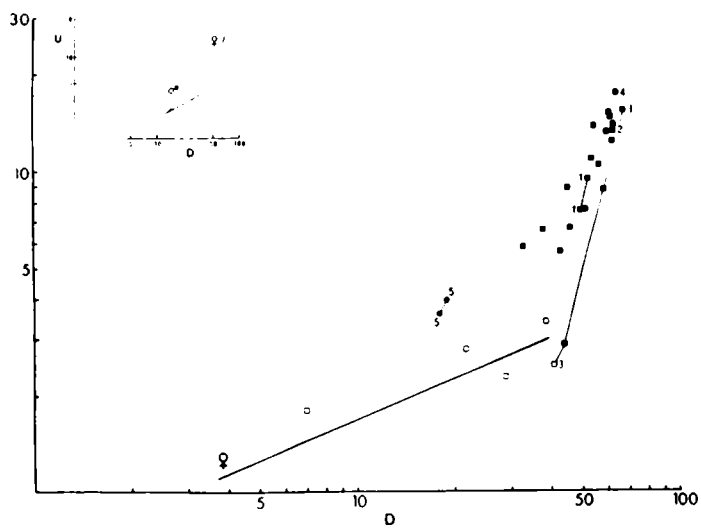
Egression of the body chamber abrupt, beginning close to the last septum, with a sudden increase in umbilical diameter from less than 10 percent to 20 to 25 percent of the shell diameter. The body chamber occupies one-half to two-thirds of a whorl and terminates with a strong constriction and a broad, smooth lip. There is a decrease in relative whorl width and height (Text-fig. 27) near the aperture but the whorl remains broadly rounded with H/W ratios of





Text-figure 28. — Cross-sections of *Chondroceras defontii* ♀ with body chamber shaded. a-b, × 1 (McM J1792c, McM J1792b); c, × 2 (McM J1792c). See *Measurements* for additional explanations.





Text-figure 27. — Bivariate plots of whorl width (W), whorl height (H) and umbilical diameter (U) against shell diameter (D) for *Chondroceras defontii* ♀ and ♂ (with synonyms) from Queen Charlotte Islands. Based on 13 macroconchs and one microconch, with several individual "growth lines". Insets are same graphs showing approximate positions of best-fit mass curves. See *Measurements* for additional explanations [In plot of U vs. D (lower right), U is measured on the ordinate axis.].

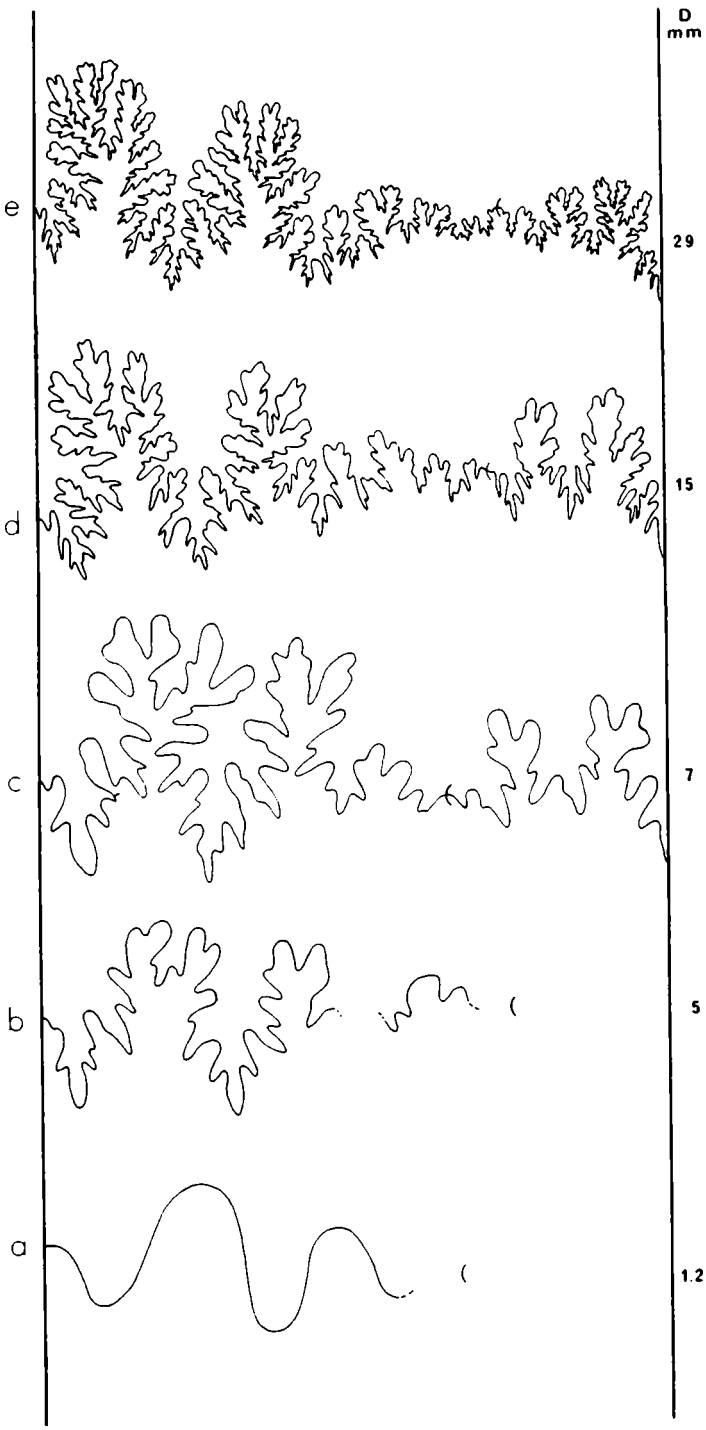
0.65-0.75; the flanks become flattened with loss of the steep umbilical wall seen on earlier whorls. Maximum diameters are between 51 and 65 mm. Ornamentation on the body chamber remains strong to the aperture. Primary ribs are straight and more widely spaced but directed forward on the flanks with 10-13 on the last half-whorl. The density of secondary ribs is reduced to 2.5 to each primary and they are noticeably coarser than those on the phragmocone.

The mature suture (Text-fig. 29) is complex with deeply incised saddles. E/L is only a little higher than L/U_2 but U_2/U_3 is very short and broad. L is about as deep as E and trifid, while U_2 is broader and trifid. The umbilical lobes are short and not retracted.

Remarks. — *Chondroceras defontii* ♀ commonly has 10-13 primary ribs per half-whorl on the body chamber and last parts of the phragmocone and up to 15 per half-whorl on earlier whorls. This density is similar to that seen on *Chondroceras* sp. from MacKenzie Bay, but *C. defontii* ♀ has a more complex suture with longer and narrower lobes and much more deeply incised saddles; its whorl section is also narrower and higher.

C. oblatum (Whiteaves) ♀ is distinguished by having a simpler suture line with broader and less deeply incised saddles (compare Text-figs. 26, 29), fewer primary ribs on

the body chamber whorl and coarser secondary ribbing on the phragmocone. *C. oblatum* ♀ also has a broader cross-section with H/W ratios of 0.60-0.65 compared with 0.65-0.75 on *C. defontii* ♀.



Text-figure 29. — Incomplete sutural ontogeny for *Chondroceras defontii* ♀ (McM J1792c). See *Measurements* for additional explanations.

No specimens intermediate in size between the holotypes of *C. defontii* ♀ and the larger *C. maudense* (McLearn) from the same locality have been found and so the latter is retained as a separate, though poorly defined, species. A specimen described as *C. defontii* by Imlay (1964; pl. 12, fig. 8) has a diameter similar to that of the holotype of *C. maudense* (75 mm just behind the aperture) but has much finer, denser ribbing with 18 primary ribs on the last half-whorl.

The four species of '*Defonticeras*' from Richardson Bay that were erected by McLearn (1929), were each based on a single specimen. *C. colnetti* was distinguished from *C. defontii* as being smaller with narrower whorls and a more strongly contracted body chamber. Table 7 shows that *C. defontii,* with a maximum shell diameter of 66 mm, is at the upper extreme of a range in shell sizes that includes *C. colnetti, C. ellsi* and *C. marchandi* and a number of other specimens of intermediate sizes. No whorl measurements from the phragmocone of *C. colnetti* are available as a result of poor preservation, so that relative contraction of the body chamber cannot be estimated; the body chamber is only slightly narrower than that on the holotype of *C. defontii.*

C. ellsi was characterized by having narrower and lower whorls, more rounded flanks, less arched venter and deeper saddles than the holotype of *C. defontii.* A slight difference in rounding of the venter and flanks is too subjective a basis for distinction. The published figures (McLearn, 1929; compare pl. XII, fig. 1 and pl. XIII, figs. 2, 3) show little difference in the suture; in addition, the suture illustrated for *C. ellsi* is probably almost a half-whorl before the last septum whereas that for *C. defontii* appears to be the last septum.

C. marchandi is at the lower extreme of the range for maximum shell diameter in this group and the relative width and height of the body chamber (phragmocone not preserved on the holotype) are similar to those of the other specimens from this locality (Table 7). Secondary ribs on the last half-whorl of the body chamber are not more numerous than on the other specimens of *C. defontii,* as stated by McLearn (1929, p. 15).

Chondroceras defontii ♂ (microconch)

(This dimorph has not been previously described.)

Allotype. — McM J1793a (Pl. 15, fig. 6) from Richardson Bay, Queen Charlotte Islands, 18-21 m above the exposed base of the Yakoun Formation.

Table 7. — Comparative measurements (in mm) for *Chondroceras defontii* (McLearn) ♀, with synonyms, from Richardson Bay, Queen Charlotte Islands.

Specimen	Maximum Diameter	U%	W%	H%	H/W	P	S
C. defontii							
GSC 9009 (holotype)	66	24	58	41	0.70	12	30
C. colnetti							
GSC 9012 (holotype)	63	28	55	40	0.73	10	27
C. ellsi							
GSC 9013 (holotype)	62	22	55	38	0.68	11	28
McM J1792a	61.5	23	59	40	0.68	13	31
McM J1792f	61	21	61	36	0.59	11	27
GSC 48594	60	25	58	40	0.70	11	30
McM J1792c	59.5	15	53	37	0.70	12	29
McM J1792b	59	23	57	39	0.68	11	32
GSC 48594	54	14	52	34	0.64	13	37
McM J1792e	51	21	57	38	0.67	12	30
C. marchandi							
GSC 9011 (holotype)	51	19	55	43	0.78	10	27

Other Material. — McM J1792g and GSC 48594, from Richardson Bay, Queen Charlotte Islands.

Description. — Maximum diameter attained is 19 mm, about one-third the size of the corresponding macroconchs. The aperture is marked by a prominent, narrow constriction followed by a narrow, smooth lip; the body chamber occupies three-quarters of a whorl. The whorl section is slightly depressed with H/W = 0.60-0.75; the venter is broad and rounded, the flanks slightly flattened. Primary ribs are strong and curved forward; there are nine on the last half-whorl with about three secondaries to each primary. Phragmocone whorls have not been preserved.

Dimorphism. — The few measurements available from microconch specimens indicate that they are similar to the macroconch in relative whorl dimensions (Text-fig. 27). Microconchs attain about one-third the size of corresponding macroconchs. Style and density of ribbing on the body chamber are similar in both dimorphs; apertural modifications are also similar, consisting of a constriction followed by a smooth lip.

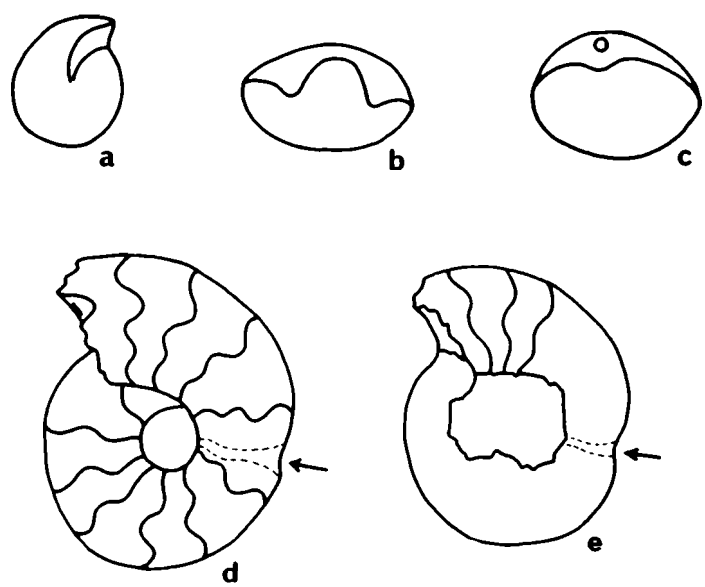
Chondroceras allani (McLearn, 1927) ♀

Pl. 15, figures 7, 8; Text-figures 30-33

1927. *Saxitoniceras allani* McLearn, p. 72; pl. 1, fig. 4.
1928. *Saxitoniceras allani* McLearn; McLearn, pp. 21, 22; pl. VIII, figs. 1, 2.
1956. *Chondroceras (Saxitoniceras) allani* (McLearn); Westermann, pp. 107, 108; pl. 12, fig. 3 (holotype re-figured).
1957. *Chondroceras allani* (McLearn); Frebold, p. 53; pl. XXVII, figs. 2a, b.
1964. *Chondroceras allani* (McLearn); Imlay, pp. B42, 43; pl. 12, figs. 4-7, 9, 10.
1964. *Chondroceras allani* (McLearn); Frebold, pp. 20, 21; pl. VIII, figs. 1-5.
1964a. *Chondroceras (Defonticeras) allani* (McLearn); Westermann, p. 55, (?microconch ♂).
1973. *Chondroceras allani* (McLearn); Imlay, p. 81; pl. 40, figs. 11, 12.

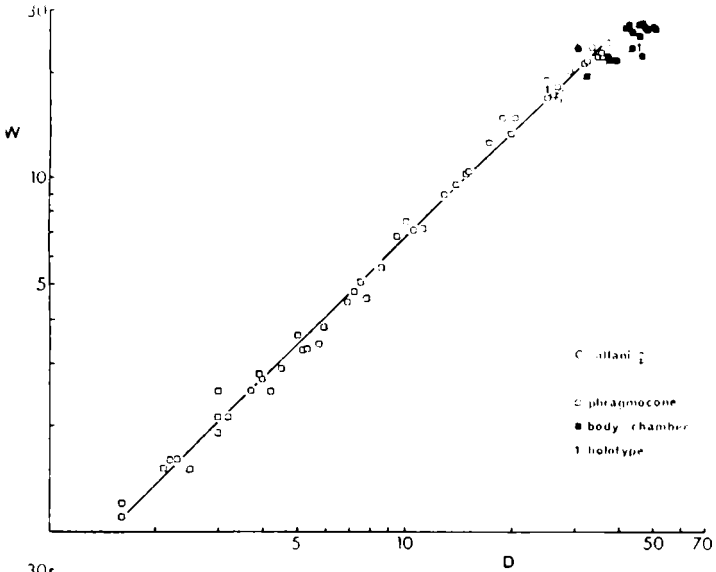
Holotype. — GSC 9021. According to the collector (quoted in McLearn, 1928, p. 22) the holotype came from the base of the Fernie Group on the headwaters of Sheep Creek, Alberta. Frebold (1957, p. 53) considers this stratigraphic position unlikely because the species is an index fossil in the Middle (= Lower, here) Bajocian Rock Creek Member which never forms the base of the Fernie.

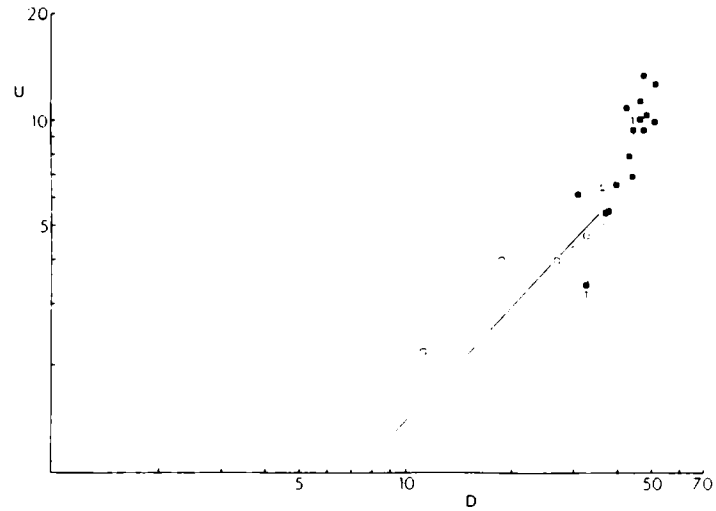
Other Material. — Fourteen relatively complete specimens (McM J1830a-i, 1831a-e) from the Rock Creek Member at Ribbon Creek, southern Alberta are now available for quantitative study of the species.

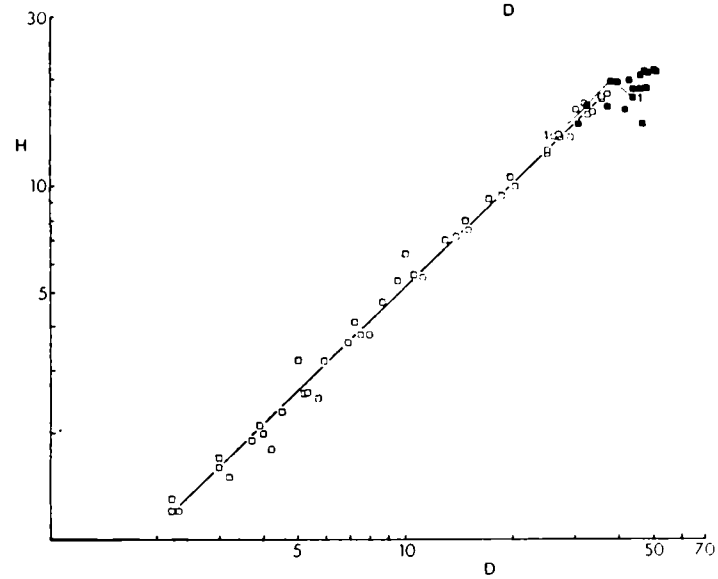


Text-figure 30. — Protoconch and nepionic whorls of *Chondroceras allani* ♀, arrow pointing to nepionic constriction, × 30 (a-d, McM J1830f; e, J1830c).

Text-figure 31. — Bivariate plots of whorl width (W), whorl height (H) and umbilical diameter (U) against shell diameter (D) for *Chondroceras allani* ♀ from southern Alberta. Based on 15 macroconchs, with several individual "growth lines". See *Measurements* for additional explanations.





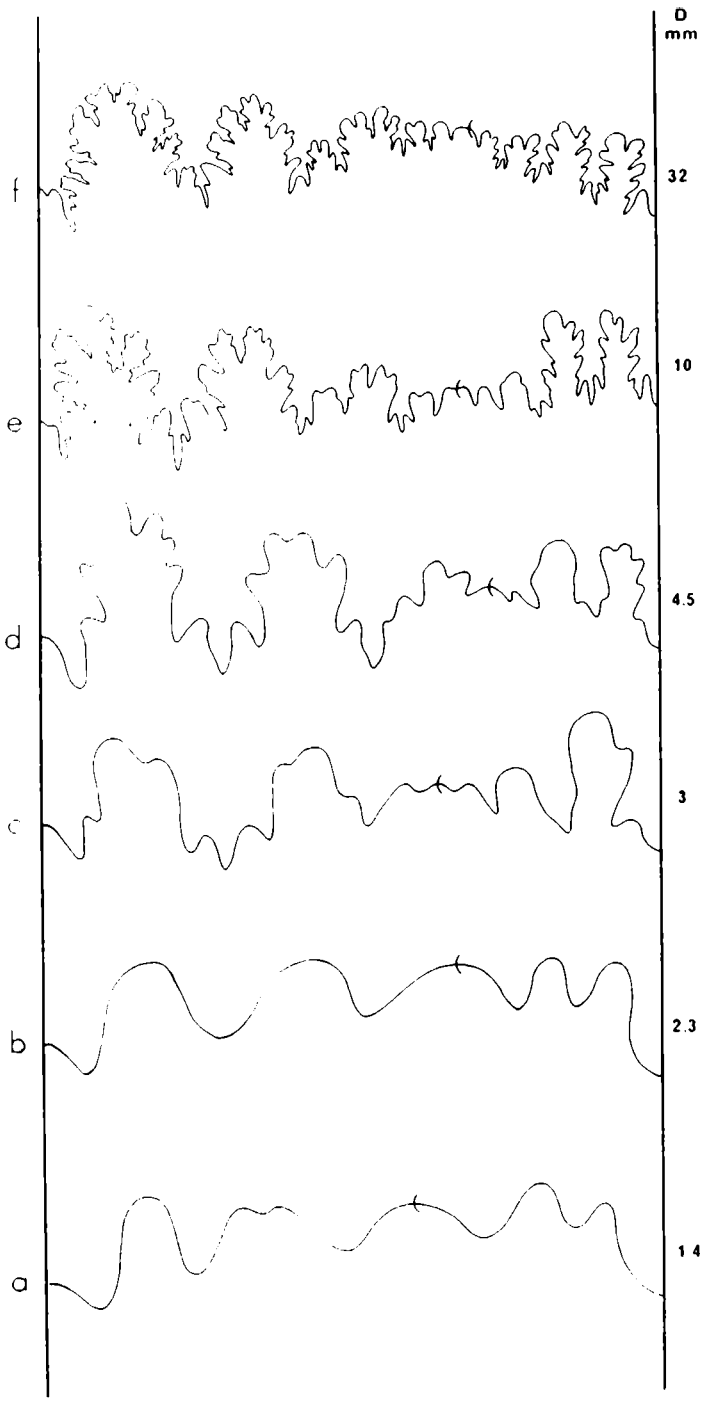


Description. — The protoconch is elongated transverse to the direction of coiling, having a width of 0.60 mm and a diameter of 0.50 mm at the position of the proseptum. The nepionic constriction occurs at a diameter of 0.80 mm, the end of the first whorl (Text-fig. 30). Ornamentation first appears at D = 5 mm in the form of broad ribbing on the ventral region only. Faint primary ribs do not appear on the flanks until a diameter of about 12 mm is reached. Shell sphaeroconic, the phragmocone whorls moderately depressed (H/W = 0.68-0.85, mostly 0.75-0.80), with flattened flanks and a strongly arched venter (Text-fig. 33). The umbilicus is deep and narrow, commonly less than five percent of the shell diameter with almost vertical walls rounding smoothly onto the flanks (Text-fig. 31). Ornamentation is not strong. There are seven to nine primary ribs per half-whorl, strongly curved forward on the flanks with three to four secondary ribs to each primary. Secondary ribs arise by bifurcation and intercalation at 50 percent of the whorl height, curving forward on the upper flanks then crossing straight over the venter. Nodes are absent.

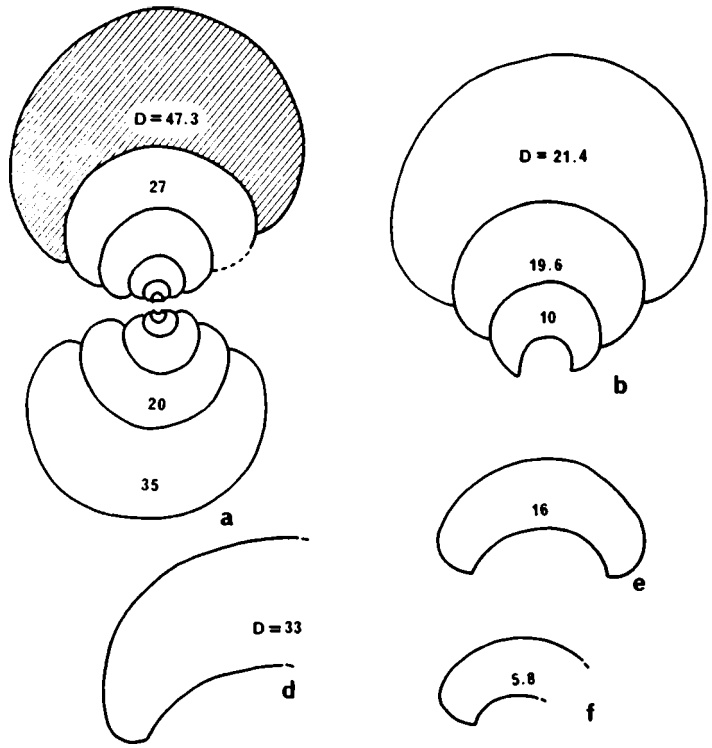
Strength and density of the ribbing remain unchanged on the body chamber except that the number of secondary ribs decreases to 2.0-2.5 per primary. The last half-whorl before the aperture bears seven to nine primary ribs curving forward on the flanks but not extending onto the smooth umbilical wall. Sudden umbilical enlargement, beginning about half a whorl before the aperture and corresponding approximately to the position of the last septum, is accompanied by a decrease in both the relative whorl width and height. The whorl section changes little (H/W = 0.75-0.80). The aperture is marked by a broad, shallow constriction, and a smooth lip.

Remarks. — *Chondroceras allani* ♀ and *C. oblatum* ♀ occur together in abundance in the Rock Creek Member of the Fernie Group at Ribbon Creek, southern Alberta. Comparison of six complete specimens of *C. allani* with fourteen complete specimens of *C. oblatum* shows that the former species has a consistently smaller adult diameter (41.5-50.0 mm vs. 51.2-65.7 mm). Phragmocone whorls of *C. allani* are

relatively narrower than those of *C. oblatum* from the Ribbon Creek locality (H/W = 0.75 vs. 0.68; W/D = 0.68 vs. 0.74 respectively).



Text-figure 32. — Sutural ontogeny for *Chondroceras allani* ♀ (McM J1830c, McM J1830h). The position of the umbilical seam is marked by (. See *Measurements* for additional explanations.



Text-figure 33. — Cross-sections of *Chondroceras* with body chamber shaded. a-b, *C. allani* ♀ from southern Alberta, respectively × 1.5 and × 3 (McM J1830g, McM J1830h); d-f, *C.* n. sp. indet. ♀ from Queen Charlotte Islands; d-e (McM J1857b, McM J1857a) × 2, f (McM J1857a) × 5. See *Measurements* for additional explanations.

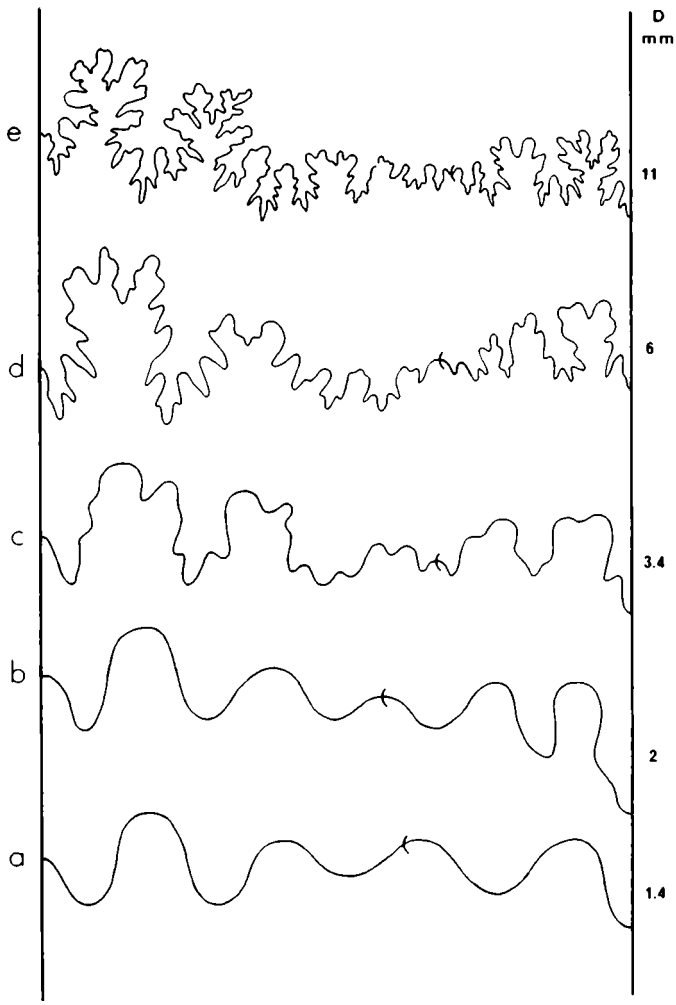
Chondroceras n. sp. indet. ♀

Plate 15, figure 9; Text-figures 33, 34

Material. — Two incomplete phragmocones (McM J1857a, b) from the lowest exposed bed of the Yakoun Formation at MacKenzie Bay; they are the first *Chondroceras* recorded at this locality.

Description. — Shell globose with broad and gently arched venter and a deep, narrow umbilicus. Whorl section strongly depressed with H/W ratios of 0.50-0.65 (Table 8). Umbilical wall short, almost vertical, rounding strongly and abruptly onto the inflated flanks to produce an umbilical shoulder.

Primary ribs long and fine, beginning at the umbilical seam and extending beyond the umbilical shoulder; they are closely spaced with 12-14 per half-whorl at diameters between 13 and 33 mm. There are usually three secondary ribs to each primary. On the umbilical wall the primary ribs are straight but crossing onto the flanks become prorsiradiate; the fine secondary ribs cross straight over the venter. Nodes are not developed.



Text-figure 34. — Sutural ontogeny of *Chondroceras* n. sp. indet. ♀ (McM J1857b). The position of the umbilical seam is marked by (. See *Measurements* for additional explanations.

The suture is only moderately complex at a diameter of 11 mm, lacking any deep incisions of the saddles (Text-fig. 34). E is shallow and broad, L slightly deeper, broad and trifid. U_2 is very broad, shallow and trifid. I and U_n are of the same length and narrow.

Remarks. — These two phragmocones differ from *C. defontii* ♀ and *C. oblatum* ♀ in details of the septal suture, ribbing density and whorl shape. *C. defontii* ♀ has much more complex sutures with deeply incised saddles and narrow lobes. *C. oblatum* ♀ has fewer primary ribs per half-whorl (7-10 vs. 12-14 at D = 15-30 mm) and higher whorl section with a narrower venter. *C. allani* ♀ has fewer primary ribs per half-whorl (8-10 vs. 12-15) and a more compressed whorl section with H/W = 0.68-0.85 on the phragmocone whorls.

Suborder PHYLLOCERATINA Arkell, 1950

Family **PHYLLOCERATIDAE** Zittel, 1884

Subfamily **CALLIPHYLLOCERATINAE** Spath, 1927

Genus **CALLIPHYLLOCERAS** Spath, 1927

Type species. — *Phylloceras disputabile* Zittel, 1869, (by original designation).

Calliphylloceras sp. indet.

Text-figures 35, 36

Material. — Three internal molds, all septate (McM J1826-8), the latter being a large fragment from a shell whose diameter must have exceeded 250 mm. Lower parts of the Yakoun Formation, South Balch Island, Queen Charlotte Islands.

Description. — Whorls strongly compressed and much higher than wide (H/W = 1.6-2.0). Venter rounded, flanks smooth and flattened. Umbilicus very narrow (U/D = 0.08 at D = 120 mm) with narrow umbilical walls that are almost vertical and curve strongly onto the flanks.

The internal molds show no sign of ribbing or other ornamentation. There are probably six to eight faint, narrow, gently sigmoidal constrictions per whorl. Features of the body chamber are unknown.

Remarks. — Phylloceratids comprise relatively small proportions of Bajocian ammonoid faunas at several localities in western North America. '*Macrophylloceras*' [= *Partschiceras*] and *Holcophylloceras* occur in southern Alaska (Imlay, 1964, pp. B32, 33); *Phylloceras, Holcophylloceras* and *Calliphylloceras* in eastern Oregon (Imlay, 1973, pp. 54, 55). Phylloceratids have not been recorded from the interior basins of western Alberta, Wyoming, Idaho and Utah.

Four specimens from the upper part of the Weberg Member of the Snowshoe Formation in eastern Oregon (correlated with the "Sowerbyi" Zone) were referred to *Calliphylloceras* by Imlay (1973, p. 54). Descriptions were brief and only one small specimen was illustrated, so that comparisons are not possible.

Geczy (1967) described and figured a large number of species of *Calliphylloceras* from Hungary, mostly from the Lower Jurassic and lower Middle Jurassic. The present species is probably much larger than most figured by Geczy and the constrictions are not as strong as those on the Hungarian species. Whorl proportions are closest to those of *C. connectens frechi* (Prinz) though the flanks of that species are more inflated. *C. liasicum* Geczy has similar whorl dimensions but a lower degree of whorl overlap and greater rate of whorl expansion.

Table 8. — Measurements (in mm) for *Chondroceras* n. sp. indet ♀ from MacKenzie Bay, Queen Charlotte Islands.

Specimen	D	U	W	H	P	S
McM J1857a	16.0	—	14.0	7.6	14	42
" "	11.3	2.1	10.0	6.7	11	32
" "	7.8	1.5	6.5	4.5	—	—
" "	5.8	1.1	4.4	2.6	—	—
" "	3.4	0.8	2.8	1.4	—	—
" "	2.0	0.6	1.5	0.8	—	—
McM J1857b	33.0	—	34.0	16.0	13	35
" "	16.3	—	11.2	6.9	12	28

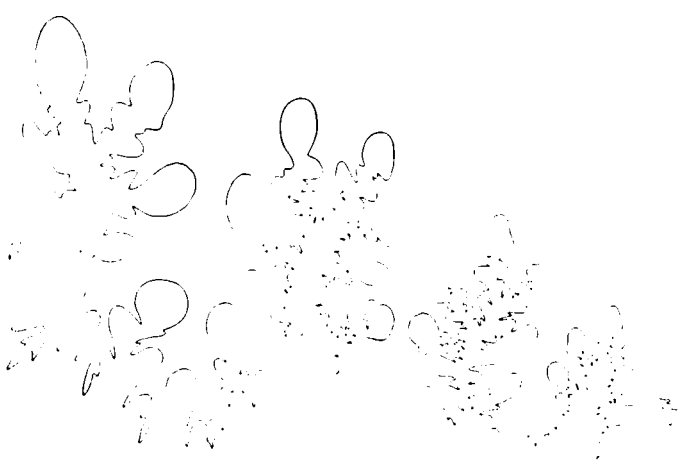
Order NAUTILIDA

Family **NAUTILIDAE** De Blainville, 1825

Genus **CENOCERAS** Hyatt, 1884

Type species. — *N. intermedius* J. Sowerby, 1816 (original designation).

Discussion. — *Cenoceras* is the most diverse of the post-Triassic nautiloids, represented by some 97 described species (Kummel, 1956, p. 361) with a cosmopolitan distribution throughout the Lower and Middle Jurassic. However, Jurassic nautiloids have only rarely been recorded from North America: four species representing two genera were described by Kummel (1954) and another two by Castillo and Aguilera (1895). *C. imlayi* (Kummel) is from the Kialagvik Formation of southern Alaska (the Aalenian *E. howelli* Zone of Westermann, 1964c) and *C. lupheri* (Kummel) is from the Weberg Formation of Oregon ("Sowerbyi" Zone).



Text-figure 35. — Part of external septal suture of *Calliphylloceras* sp. indet. at D = 110 mm (McM J1826).

Cenoceras sp. indet.

Text-figure 36

Material. — A single internal mold, wholly septate and about 50 mm in diameter (McM J1825b), from the lower parts of the Yakoun Formation on the north-western shore of South Balch Island, approximately 38 m above the lowest exposure of the Formation on this island. Another from the lowest exposed strata of the Yakoun Formation at MacKenzie Bay (McM J1881).

Description. — Shell globose with deep, narrow umbilicus (15 percent of shell diameter); umbilical wall vertical with a broadly rounded shoulder that is the widest part of the whorl. Flanks only slightly flattened, convergent toward the venter, which is broad and somewhat flattened. Whorls depressed, H/W = 0.66 at a diameter of 50 mm. Siphuncle subdorsal; septa concave toward aperture.

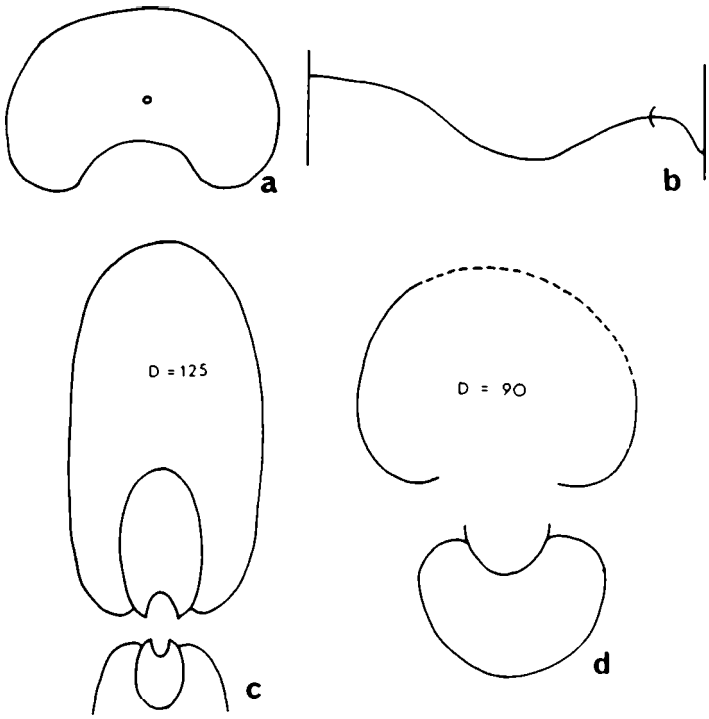
The inner whorls have faint ornamentation preserved in places, consisting of both growth lines and lirae. Suture smooth, straight across the venter with a broad, shallow lateral lobe and a short, narrow V-shaped dorsal lobe (Text-fig. 36b).

Remarks. — This specimen resembles *C. lupheri* (Kummel) from the Weberg Member of the Snowshoe Formation in eastern Oregon in whorl shape and relative umbilical diameter. Erroneously described as "compressed" by Kummel (1954, p. 323) the whorl section of *C. lupheri* has a H/W ratio of 0.72 at the adoral end of the phragmocone which is similar to that of the present specimen (0.66 at D = 50 mm; Text-fig. 36a). However, no shell ornamentation is preserved on the holotype of *C. lupheri*.

Genus **EUTREPHOCERAS** Hyatt, 1894

Type species. — *Nautilus dekayi* Morton, 1834 (by original designation).

Discussion.—Morphological distinction between the two large nautiloid stocks represented by *Cenoceras* and *Eutrephoceras* appears difficult to maintain. According to the detailed discussions of Kummel (1956), *Eutrephoceras* has an



Text-figure 36. — a-b, cross-section (× 0.6) of septate whorl and septal suture of *Cenoceras* sp. indet. at D = 50 mm (McM J1825b); c, cross-section of fully septate *Calliphylloceras* sp. indet. (McM J1826), × 0.6; d, cross-section of fully septate *Eutrephoceras?* sp. indet., × 0.6 (McM J1825a).

essentially smooth conch and globose whorl section while *Cenoceras* typically bears fine lirae and growth lines and has a more quadrate whorl section. The name *Cenoceras* is usually given to Upper Triassic - Middle Jurassic forms, while *Eutrephoceras* ranges from the Upper Jurassic to Miocene (Kummel, 1964, p. 449). However, the range of whorl shapes illustrated for *Cenoceras* (Pia, 1914, reproduced as figs. 8, 9, and 10 in Kummel, 1956) clearly includes forms that also appear within the range illustrated for *Eutrephoceras* (Kummel, 1956, fig. 13). There are specimens from the Cenomanian (Upper Cretaceous) of France in the McMaster University collections (McM K44) which bear fine lirae and growth lines. *Eutrephoceras* sp. indet. from the Queen Charlotte Islands, an apparently smooth-shelled form with globose whorl section, is of Early Bajocian (Middle Jurassic) age.

If ornamentation and whorl shape are used to separate *Cenoceras* and *Eutrephoceras*, then their supposed age difference must be disregarded. Indeed, Kummel (1954) has already listed several Middle Jurassic species that he considers belong to *Eutrephoceras* (p. 321).

?Eutrephoceras sp. indet.

Text-figure 36

Material. — Fully septate, internal mold about 90 mm in diameter, McM J1825a, from the lower part of the Yakoun Formation, about 38 m above the exposed base of the formation on South Balch Island.

Description. — Shell globose with slightly flattened sides converging towards the venter; umbilical shoulder broadly rounded and the widest part of the whorl (Text-fig. 36d). Umbilical wall vertical; umbilicus very narrow (15 percent of shell diameter). Whorls depressed with H/W = 0.68-0.74. No ornamentation preserved. Position of the siphuncle unknown. Suture with external ventral and lateral lobe, both broad and shallow; internal suture unknown.

Remarks. — *E. montanensis* Kummel from the Callovian Rierdon Formation in Montana has much narrower, compressed whorls (H/W = 1.3 vs. 0.68-0.74). Another specimen of *Eutrephoceras* has been recorded from the Bajocian Twin Creek Limestone in Wyoming (Kummel, 1954, p. 321), but it was not described or illustrated.

REFERENCES

Arkell, W. J.
1950. *A classification of the Jurassic ammonites.* J. Paleontol., vol. 24, pp. 354-364.
1956. *Jurassic geology of the world.* Oliver & Boyd (London & Edinburgh), xv + 804 pp., pls. 1-46.

Arkell, W. J., Kummel, B., and Wright, C. W.
1957. *Mesozoic Ammonoidea,* in *Treatise on Invertebrate Paleontology, Part L, Mollusca 4.* Geol. Soc. Amer. & Univ. Kansas Press, pp. L80-L437, figs. 124-555.

Arkell, W. J., and Playford, P. E.
1954. *The Bajocian ammonites of Western Australia.* R. Soc. London. Philos. Trans. Ser. B, vol. 237, pp. 547-601.

Billings, E.
1873. *On the Mesozoic fossils from British Columbia collected by Mr. James Richardson in 1872.* Can., Geol. Surv., Rep. Prog., 1872-1873, pp. 71-75.

Boehm, G.
1913. *Unteres Callovien und Coronatenschichten zwischen Mac Cluer-Golf und Geelvink-Bai.* In C. E. A. Wichmann, Nova Guinea, vol. 6, livr. 1 (Geologie), 20 pp.

Bruguière, J. C.
1789-92. *Encyclopedie Méthodique; histoire naturelle des. vers.* Vol. 1, pt. 1, pp. 1-344.

Buckman, S. S.
1898. *On the grouping of some divisions of so-called "Jurassic" time.* Geol. Soc. London, Q. J., vol. 50, pp. 442-462.
1908. *Illustrations of type specimens of Inferior Oolite ammonites in the Sowerby collection.* Palaeontogr. Soc., Monogr., vol. 62, pls. 1-7.
1919-21. *Type ammonites,* part III, pp. 1-64. Weldon & Wesley, London.

Castillo, A. del, and Aguilera, J. G.
1895. *Fauna fósil de la Sierra de Catorce, San Luis Potosí.* Com. Geol. Mexico, Bol. 1, IX + 55 pp., 24 pls.

Clapp, C. H.
1914. *A geological reconnaissance on Graham Island, Queen Charlotte Group, B. C.* Can., Geol. Surv., Summ. Rep. 1912, pp. 12-40.

Colloque Du Jurassique, Luxembourg
1971. *Colloque du Jurassique a Luxembourg, 1967.* Fr., Bur. Rech. Geol. Minieres, Mem. No. 75.

Crickmay, C. H.
1933. *Mount Jura investigation.* Geol. Soc. Am., Bull. 44, pp. 895-926.

Dawson, G. M.
1880. *Report on the Queen Charlotte Islands, 1878.* Can., Geol. Surv., Rep. Prog. 1878-1879, vii + 239 pp., 52 illus., maps.
1889. *On the earlier Cretaceous rocks of the northwestern portion of the Dominion of Canada.* Am. J. Sci., ser. 3, vol. 38, pp. 120-127.

Dean, W. T., Donovan, D. T., and Howarth, M. K.
1961. *The Liassic ammonite zones and subzones of the northwest European province.* Br. Mus. (Nat. Hist.), Bull., Geol., vol. 4, pp. 435-505.

Dickinson, W. R., and Vigrass, L. W.
1965. *Geology of the Suplee-Izee area, Crook, Grant and Harney Counties, Oregon.* Oregon, State Dep. Geol. Miner. Ind., Bull., 58, 109 pp., illus.

Ells, R. W.
1906. *Report on Graham Island, B. C.* Can., Geol. Surv., Annu. Rep., n.s., vol. XVI, pt. B, 1904, pp. 1-46.

Frebold, H.
1957. *The Jurassic Fernie Group in the Canadian Rocky Mountains and Foothills.* Can., Geol. Surv., Mem. 287, 197 pp., incl. 44 pls.
1961. *The Jurassic faunas of the Canadian Arctic. Middle and Upper Jurassic ammonites.* Can., Geol. Surv., Bull. 74, 43 pp., 21 pls.
1964. *Lower Jurassic and Bajocian ammonoid faunas of northwestern British Columbia and southern Yukon.* Can., Geol. Surv., Bull. 116, 31 pp., 8 pls.
1976. *The Toarcian and lower Middle Bajocian beds and Ammonites in the Fernie Group of southeastern British Columbia and Parts of Alberta.* Can., Geol. Surv., Pap. 75-39, 33 pp., illus.

Frebold, H., Mountjoy, E. W., and Tempelman-Kluit, D. J.
1967. *New occurrences of Jurassic rocks and fossils in central and northern Yukon Territory.* Can., Geol. Surv., Pap. 67-12, 35 pp., illus., tables.

Frebold, H., and Tipper, H. W.
1970. *Status of the Jurassic in the Canadian Cordillera of British Columbia, Alberta, and southern Yukon.* Can. J. Earth Sci., vol. 7, pp. 1-21.
1973. *Upper Bajocian-Lower Bathonian ammonite fauna and stratigraphy of Smithers area, British Columbia.* Can. J. Earth Sci., vol. 10, pp. 1109-1131.

Frebold, H., Tipper, H. W., and Coates, J. A.
1969. *Toarcian and Bajocian rocks and guide ammonites from southwestern British Columbia.* Can., Geol. Surv., Pap. 67-10, 55 pp., illus.

Geczy, B.
1967. *Ammonoides Jurassiques de Csernye, Montagne Bakony, Hongrie. Part II. (excl. Hammatoceratidae).* Geol. Hung. Ser. Palaeontol. No. 35, 282 pp., illus.

Gordon, W. A.
1976. *Ammonoid provincialism in space and time.* J. Paleontol., vol. 50, pp. 521-535.

Guex, J.
1967. *Dimorphisme sexuel d'un groupe d'*Hammatoceras *et position systématique du genre* Onychoceras. Lausanne Univ., Lab. Géol., Minéral., Géophys., Mus. Géol., Bull., No. 168, pp. 1-12.

Hallam, A.
1969. *Faunal realms and facies in the Jurassic.* Palaeontology, vol. 12, pp. 1-18.
1971. *Provinciality in Jurassic faunas in relation to facies and paleogeography.* In *Faunal provinces in space and time.* F. A. Middlemiss et al. (Eds.). Geol. J., Special Issue No. 4, pp. 129-152.

Howarth, M. K.
1973. *Lower Jurassic (Pliensbachian and Toarcian) ammonites.* In *Atlas of Palaeobiogeography.* A. Hallam (Ed.). Elsevier, pp. 275-282.

Hyatt, A.
1883-84. *Genera of fossil cephalopods.* Boston Soc. Nat. Hist., Proc., vol. 22, pp. 253-338.
1889. *Genesis of the Arietidae.* Smithson. Contrib. Knowl. Vol. 26, no. 673, xi + 238 pp., illus.
1894. *Phylogeny of an acquired characteristic.* Am. Philos. Soc., Proc., vol. 32, pp. 349-647, illus.

Imlay, R. W.
1961. *New genera and subgenera of Jurassic (Bajocian) ammonites from Alaska.* J. Paleontol., vol. 35, pp. 467-474, pls. 63-64.
1962. *Late Bajocian ammonites from the Cook Inlet region, Alaska.* U.S. Geol. Surv., Prof. Pap. 418-A, 15 pp., 5 pls.
1964. *Middle Bajocian ammonites from the Cook Inlet Region, Alaska.* U.S. Geol. Surv., Prof. Pap. 418-B, 61 pp., 29 pls.
1965. *Jurassic marine faunal differentiation in North America.* J. Paleontol., vol. 39, pp. 1023-1038, 6 text-figs.
1967. *Twin Creek Limestone (Jurassic) in the Western Interior of the United States.* U.S. Geol. Surv., Prof. Pap. 540, 105 pp., 16 pls.
1973. *Middle Jurassic (Bajocian) ammonites from Eastern Oregon.* U.S. Geol. Surv., Prof. Pap. 756, 100 pp., 47 pls., map.
1976. *Middle Jurassic (Bajocian and Bathonian) ammonites from northern Alaska.* U.S. Geol. Surv., Prof. Pap. 854, 22 pp., 4 pls.

Khudoley, K. M.
1974. *Circum-Pacific Mesozoic ammonoid distribution: Relation to hypotheses of Continental Drift, Polar Wandering, and Earth Expansion.* In *Plate Tectonics — Assessments and reassessments.* C. F. Kahle (Ed.). Am. Ass. Pet. Geol., Mem. 23, pp. 295-330.

Kummel, B.
1954. *Jurassic Nautiloids from western North America.* J. Paleontol., vol. 28, pp. 320-324, pls. 34-35.
1956. *Post-Triassic nautiloid genera.* Harv. Univ., Mus. Comp. Zool., Bull. 114, No. 7, pp. 324-494, illus.
1964. *Nautiloidea — Nautilida. Treatise on Invertebrate Paleontology, Part K, Mollusca 3.* R. C. Moore (Ed.). Geol. Soc. Amer. & Univ. Kansas Press, pp. K383-K457, figs. 280-337.

Lupher, R. L.
1941. *Jurassic stratigraphy of central Oregon.* Geol. Soc. Am., Bull. 52, pp. 219-270, 4 pls., 3 figs.

MacKenzie, J. D.
1916. *Geology of Graham Island, British Columbia.* Can., Geol. Surv., Mem. 88, x + 221 pp., maps, illus.

Makowski, H.
1962. *Problem of sexual dimorphism in ammonites.* Palaeontol., Pol., vol. 12, 92 pp., 26 text-figures, 20 pls.

Mascke, E.
1907. *Die* Stephanoceras-*Verwandten in den Coronatenschichten von Norddeutschland.* Göttingen Univ., thesis, 38 pp.

McLearn, F. H.
1927. *Some Canadian Jurassic faunas.* R. Soc. Can., Trans., 3rd ser., vol. 21, sec. 4, pp. 61-73, 1 pl.
1928. *New Jurassic Ammonoidea from the Fernie Formation, Alberta.* Geol. Sur. Canada, Museum Bull. 49, pp. 19-22.
1929. *Contribution to the stratigraphy and palaeontology of Skidegate Inlet, Queen Charlotte Islands, B.C.* Natl. Mus. Can., Bull. 54, 27 pp., 16 pls.
1930. *Notes on some Canadian Mesozoic faunas.* R. Soc. Can. Trans., 3rd ser., vol. 24, sec. 4, 8 pp., 2 pls.
1932a. *Contributions to the stratigraphy and palaeontology of Skidegate Inlet, Queen Charlotte Islands, B.C. (continued).* R. Soc. Can., Trans., 3rd ser., vol. 26, sec. 4, pp. 51-84, 10 pls.

1932b. *Three Fernie Jurassic ammonoids.* R. Soc. Can., Trans., 3rd ser., vol. 26, sec. 4, p. 111-115, 5 pls.
1949. *Jurassic formations of Maude Island and Alliford Bay, Skidegate Inlet, Queen Charlotte Islands, British Columbia.* Can., Geol. Surv., Bull. 12, 19 pp., illus., maps.

Morton, N.
1971. *Some Bajocian ammonites from western Scotland.* Palaeontology, vol. 14, pp. 266-293, pls. 40-51.
1972. *The Bajocian ammonite* Dorsetensia *in Skye, Scotland.* Palaeontology, vol. 15, p. 504-518, pls. 102-105.
1975. *Bajocian Sonniniidae and other ammonites from western Scotland.* Palaeontology, vol. 18, p. 41-91, pls. 6-17.
1976. *Bajocian (Jurassic) stratigraphy in Skye, western Scotland.* Scott. J. Geol., vol. 12, Pt. 1, pp. 23-33, illus.

Mouterde, R., Enay, R., Cariou, E., Contini, D., Elmi, S., Gabilly, J., Mangold, Ch., Mattei, J., Rioult, M., Thierry, J., and Tintant, H.
1971. *Les zones du Jurassique en France.* Soc. Géol. Fr., C. R. No. 2, pp. 76-102.

Neumayr, M.
1875. *Die Ammoniten der Kreide und die Systematik der Ammonitiden.* Zeitschr. deutsch. geol. Gesell., Band 27, pp. 854-943.
1883. *Über klimatische Zonen während der Jura- und Kreideperiode.* Denkschr. d. Kais. Akad. Wiss. Wien, vol. 47, pp. 276-310.

Palframan, D. F. B.
1969. *Taxonomy of sexual dimorphism in ammonites: Morphogenetic evidence in* Hecticoceras brightii *(Pratt).* In *Sexual dimorphism in fossil metazoa and taxonomic implications.* G. E. G. Westermann (Ed.). Intl. Union Geol. Sci., ser. A, No. 1, pp. 126-152, pls. 6-8.

Parsons, C. F.
1974. *The* sauzei *and 'so-called'* sowerbyi *Zones of the Lower Bajocian.* Newsl. Stratigr., vol. 3, pp. 153-179, illus.
1976. *A stratigraphic revision of the* humphriesianum/subfurcatum *Zone rocks (Bajocian Stage, Middle Jurassic) of Southern England.* Newsl. Stratigr., vol. 5, pp. 114-142, 5 figs., 3 tables.

Pavia, G.
1969. *Nouvelles données sur le bajocien de Digne (Basses-Alpes); Coéxistence de* Normannites, Itinsaites, Stemmatoceras, Teloceras *avec les premiers représentants de la superfamille Perisphinctaceae à la base de la zone à Subfurcatum.* Soc. Geol. Ital. Boll., vol. 88, p. 445-451, illus.

Pavia, G., and Sturani, C.
1968. *Etude biostratigraphique du bajocien des chaines subalpines aux environs de Digne (Basses Alpes).* Soc. Geol. Ital. Boll., vol. 87, pp. 305-316, illus.

Pia, J.
1914. *Untersuchungen über die liasischen Nautiloidea.* Beitr. z. Paläontol. u. Geol. Osterreichungarns Orients, vol. 27, pp. 19-86, 7 tables.

Quenstedt, F. A. v.
1886-87. *Die Ammoniten des Schwäbischen Jura, v. 2, Der Braune Jura.* Stuttgart. pp. 441-815.

Renz, C.
1913. *Zur Geologie des östlichen Kaukasus.* Neues Jahrb. Min. Geol. Palaeontol., Beil.-Band 36, pp. 651-703, Taf. XXV-XXIX, 3 text-figs.

Richardson, J.
1873. *Report on the Coal-Fields of Vancouver and Queen Charlotte Islands.* Can. Geol. Surv., Rep. Prog., 1872-1873, pp. 32-65, map.

Roché, P.
1939. *Aalénien et Bajocien du Maconnais et de quelques régions voisines.* Lyon Univ., Fac. Sci., Lab. Géol. Trav. f. 35, 355 pp., 12 figs., 13 pls.

Sapunov, I. G.
1971. *Notes on the geographic differentiation of the Lower Jurassic ammonite faunas.* In *Colloque du Jurassique a Luxembourg, 1967.* Fr., Bur. Rech. Geol. Minieres, Mem., 75, pp. 263-270.

Schmidtill, E., and Krumbeck, L.
1938. *Die Coronaten-Schichten von Auerbach (Oberpfalz, Nordbayern).* Dtsch. Geol. Ges., Bd. 90, pp. 297-360, 5 figs., 5 pls.

Sowerby, J.
1812-25. *The mineral conchology of Great Britain.* vols. 1-4, 1812-1825. London.

Sowerby, J. de C.
1825. *The mineral conchology of Great Britain.* vols. 5-7, 1825-1829. London.

Spath, L. F.
1927-33. *Revision of the Jurassic cephalopod faunas of Kachh (Cutch).* India Geol. Sur. Mem., Palaeont. Indica, new ser., vol. 9, mem. 2, pts. 1-6, 945 pp.
1944. *Problems of ammonite nomenclature, IX; The genus* Stephanoceras, *Waagen and some allied genera.* Geol. Mag., vol. 81, pp. 230-234.

Stanton, T. W., and Martin, G. C.
1905. *Mesozoic section on Cook Inlet and Alaska Peninsula.* Geol. Soc. Am., Bull. 16, pp. 391-410, pls. 67-70.

Stevens, G. R.
1971. *Relationship of isotopic temperatures and faunal realms to Jurassic-Cretaceous paleogeography, particularly of the south-west Pacific.* R. Soc. N. Z., J., vol. 1, pp. 145-158, illus.

Stevens, G. R., and Clayton, R. N.
1971. *Oxygen isotope studies on Jurassic and Cretaceous belemnites from New Zealand and their biogeographic significance.* N. Z. J. Geol. Geophys., vol. 14, pp. 829-897, illus.

Sturani, C.
1964. *La successione delle faune ad ammoniti nelle formazioni mediogiurassiche delle Prealpi Venete occidentali (regione tra il Lago di Garda e la valle del Brenta).* Padua Univ., Inst. Geol. Miner., Mem. XXIV, pp. 1-65, 6 pls.
1971. *Ammonites and stratigraphy of the* "Posidonia alpina" *beds of the Venetian Alps (Middle Jurassic, mainly Bajocian).* Inst. Geol. Mineral. Univ. Padova, Mem. XXVIII, pp. 1-90.

Sutherland Brown, A.
1968. *Geology of the Queen Charlotte Islands, British Columbia.* Brit. Col. Dept. Mines Petrol. Res., Bull. 54, pp. 1-226, 18 pls., maps.

Uhlig, V.
1911. *Die marinen Reiche des Jura und der Unterkreide.* Geol. Ges. Wien, Mitt. 4, [Heft 3], pp. 329-448, Tab. 13.

Waagen, W.
1869. *Die Formenreihe des* Ammonites subradiatus. Geogn.-Paläont. Beitr., Band 2, Heft 2, pp. 181-256.

Warren, P. S.
1934. *Present status of the Fernie Shale, Alberta.* Am. J. Sci., vol. 27, pp. 56-70, 1 fig.
1947. *Description of Jurassic ammonites from the Fernie Formation, Alberta.* Alberta Research Council, Rept. 49, pp. 67-76, illus.

Weisert, K.
1932. Stephanoceras *im Schwäbischen Braunen Jura-Delta.* Palaeontographica, vol. 76, pp. 121-191, 44 text-figs., 5 taf.

Westermann, G. E. G.
1954. *Monographie der Otoitidae (Ammonoidea).* Beih. Geol. Jb., Heft 15, pp. 1-364.
1956. *Monographie der Bajocian-Gattungen* Sphaeroceras *und* Chondroceras *(Ammonoidea).* Beih. Geol. Jb., Heft 24, pp. 1-125.
1964a. *Sexual-Dimorphismus bei Ammonoideen und seine Bedeutung für die Taxionomie der Otoitidae.* Palaeontogr. Abt. A., Bd. 124, Lf. 1-3, pp. 33-71, illus.
1964b. *Occurrence and significance of the Arctic* Arkelloceras *in the Middle Bajocian of the Alberta Foothills (Ammonitina, Jurassic).* J. Paleontol., vol. 38, pp. 405-409, illus.

Westermann, G. E. G.

1964c. *The ammonite fauna of the Kialagvik Formation at Wide Bay, Alaska Peninsula. Part I. Lower Bajocian (Aalenian).* Bull. Am. Paleontol. vol. 47, No. 216, pp. 325-503, incl. 32 pls., 37 text-figs.

1969a. *Proposal: Classification and nomenclature of dimorphs at the genus-group level.* In *Sexual dimorphism in fossil Metazoa and taxonomic implications.* G. E. G. Westermann (Ed.). Intl. Union of Geol. Sci. ser. A, vol. 1, pp. 234-238.

1969b. *The ammonite fauna of the Kialagvik Formation at Wide Bay, Alaska Peninsula. Part II.* Sonninia sowerbyi *Zone (Bajocian).* Bull. Am. Paleontol., vol. 57, 226 pp., incl. 47 pls.

Westermann, G. E. G., and Getty, T. A.

1970. *New Middle Jurassic Ammonitina from New Guinea.* Bull. Am. Paleontol., vol. 57, pp. 227-321, pls. 48-62.

Westermann, G. E. G., and Riccardi, A. C.

1972. *Middle Jurassic ammonoid fauna and biochronology of the Argentine-Chilean Andes; Part I, Hildoceratacea.* Palaeontogr., Abt. A, vol. 140, pp. 1-109, illus.

1976. *Middle Jurassic ammonite distribution and the affinities of the Andean faunas.* Prim. Congreso Geol. Chileno, vol. 1, pp. C23-39.

1979. *Middle Jurassic ammonoid fauna and biochronology of the Argentine-Chilean Andes. Part II. Bajocian Stephanocerataceae.* Palaeontogr., Abt. A, vol. 164, pp. 1-188, 28 pls.

White, J. F., and Gould, S. J.

1965. *Interpretation of the coefficient in the allometric equation.* Am. Nat., vol. 99, pp. 5-18.

Whiteaves, J. F.

1876. *Mesozoic fossils. Vol. 1. Pt. I. On some invertebrates from the coal-bearing rocks of the Queen Charlotte Islands, collected by J. Richardson in 1872.* Can., Geol. Surv., 92 pp., 10 pls., map.

1884. *Mesozoic fossils. Vol. 1. Pt. III. On the fossils of the coal-bearing deposits of the Queen Charlotte Islands collected by G. M. Dawson in 1878.* Geol. Sur. Canada, pp. 191-262, 22 pls.

1900. *Mesozoic fossils. Vol. 1. Pt. IV. On some additional or imperfectly understood fossils from Cretaceous rocks of the Queen Charlotte Islands, with a revised list of the species from these rocks.* Geol. Sur. Canada, pp. 263-307, pls. 33-39.

Zittel, K. A. von

1869. *Bemerkungen über* Phylloceras tatricum *Puch sp. und einige andere Cephalopoden.* Jahrb. Kais.-Kögl. Reichsanst., vol. 19, pp. 59-81.

1884. *Handbuch der Palaeontologie.* Abt. 1, Band 2, 893 pp.

Appendix 1. — Data and Regression Equations for the "mass curves" shown on Text-figures for each taxon. Statistics calculated on measurements from phragmocone whorls only. *indicates regression coefficient significantly different from zero at 5% level of significance. **indicates correlation coefficient significant at 5% level of significance.

	Text-figure Number	Number of specimens	Number of individual measurements	Regression Equation	Standard error of slope	Correlation coefficient
Zemistephanus richardsoni (Whiteaves, 1876) ♀ & ♂						
D vs. W ♀ [macroconch]	5	16	67	Y = 0.67X — 0.54	.0104*	0.99**
♂ [microconch]	5	5	22	Y = 0.63X — 0.64	.019*	0.98**
D vs. H ♀ [macroconch]	5	16	73	Y = 0.38X — 0.24	.0056*	0.99**
♂ [microconch]	5	5	22	Y = 0.34X + 0.02	.0109*	0.98**
D vs. U ♀ [macroconch]	5	16	66	Y = 0.36X + 0.35	.0068*	0.98**
♂ [microconch]	5	5	19	Y = 0.42X — 0.37	.017*	0.98**
Zemistephanus crickmayi (McLearn, 1927) ♂						
D vs. W ♂ [microconch]	9	13	62	Y = 0.54X + 0.73	.011*	0.98**
D vs. H ♂ [microconch]	9	13	62	Y = 0.35X + 0.04	.005*	0.99**
D vs. U ♂ [microconch]	9	13	59	Y = 0.41X — 0.22	.0057*	0.99**
Stephanoceras itinsae (McLearn, 1927) ♀ & ♂						
D vs. W ♀ [macroconch]	13	13	55	Y = 0.43X + 4.55	.0117*	0.97**
♂ [microconch]	13	11	40	Y = 0.53X + 0.46	.014*	0.98**
D vs. H ♀ [macroconch]	13	13	55	Y = 0.31X + 1.22	.006*	0.98**
♂ [microconch]	13	11	40	Y = 0.33X + 0.08	.0057*	0.99**
D vs. U ♀ [macroconch]	13	13	56	Y = 0.46X — 1.62	.009*	0.98**
♂ [microconch]	13	11	34	Y = 0.45X — 0.32	.01*	0.99**
Stephanoceras skidegatense (Whiteaves, 1876) ♀ & ♂						
D vs. W ♀ [macroconch]	18	9	25	Y = 0.38X + 3.57	.017*	0.97**
♂ [microconch]	18	2	14	Y = 0.59X — 0.15	.016*	0.99**
D vs. H ♀ [macroconch]	18	9	25	Y = 0.31X + 0.68	.0086*	0.99**
♂ [microconch]	18	2	14	Y = 0.38X — 0.35	.008*	0.99**
D vs. U ♀ [macroconch]	18	9	25	Y = 0.49X — 1.84	.010*	0.99**
♂ [microconch]	18	2	11	Y = 0.42X — 0.12	.0056*	0.99**
Chondroceras defontii (McLearn, 1927) ♀						
D vs. W ♀ [macroconch]	27	4	8	Y = 0.76X — 0.17	.07*	0.96**
D vs. H ♀ [macroconch]	27	4	8	Y = 0.49X + 0.84	.029*	0.98**
D vs. U ♀ [macroconch]	27	3	7	Y = 0.05X + 1.2	.012*	0.79**

Chondroceras oblatum (Whiteaves, 1876) ♀ & ♂						
D vs. W ♀ [macroconch]	24	4	20	Y = 0.76X — 0.14	.02*	0.99**
♂ [microconch]	24	1	6	Y = 0.76X + 0.28		0.99
D vs. H ♀ [macroconch]	24	4	21	Y = 0.48X + 0.24	.025*	0.97**
♂ [microconch]	24	1	6	Y = 0.48X + 0.07		0.98
"Saxitoniceras marshalli" McLearn, 1927 ♀ [= *Chondroceras oblatum*]						
D vs. W ♀ [macroconch]	24	9	48	Y = 0.79X — 0.31	.018*	0.98**
D vs. H ♀ [macroconch]	24	9	48	Y = 0.51X + 0.03	.011*	0.98**
D vs. U ♀ [macroconch]	24	8	24	Y = 0.13X — 0.06	.01*	0.92**
Chondroceras allani (McLearn, 1927) ♀						
D vs. W ♀ [macroconch]	31	8	46	Y = 0.69X — 0.05	.009*	0.99**
D vs. H ♀ [macroconch]	31	8	46	Y = 0.50X + 0.13	.006*	0.99**
D vs. U ♀ [macroconch]	31	6	17	Y = 0.16X — 0.28	.018*	0.88**

Appendix 2.—Measurements and data used in formal statistical comparisons of sexual dimorphs. Only measurements from phragmocone whorls were used; there are marked changes in growth ratios at the end of the nepionic whorl and the beginning of the body chamber. A single measurement from each specimen was used (all measurements are in mm).

	D	W	H	U
Zemistephanus richardsoni (Whiteaves, 1876) ♀ [macroconch]	44.0	30.0	18.5	17.0
	48.0	34.0	17.0	19.8
	29.5	16.4	11.2	10.0
	17.7	10.3	7.0	6.4
	11.3	6.5	3.4	—
	5.4	3.4	1.9	—
	3.9	3.0	1.4	1.3
	12.0	6.9	3.5	5.0
	41.6	27.0	15.2	15.0
	38.3	23.0	13.5	12.9
Zemistephanus richardsoni (Whiteaves, 1876) ♂ [microconch]	8.7	4.6	2.8	3.6
	21.6	11.9	7.4	6.7
	35.5	23.5	12.3	14.0
	28.5	18.0	10.3	—
	37.1	22.5	12.6	14.3

Regression Analysis for *Zemistephanus richardsoni*

D vs. W ♀ $Y = 0.679X - 1.06$
♂ $Y = 0.675X - 1.658$
(a) Variance: $F = 2.21$; $F_{.05}(9, 4) = 6.0$; hypothesis of equal variances accepted
(b) Slopes: $t = 0.06$; $t_{(.05, 11)} = 1.796$; hypothesis of equal slopes accepted
(c) Intercepts: $t = 23.33$; $t_{(.05, 12)} = 1.782$; hypothesis of equal intercepts rejected

D vs. H ♀ $Y = 0.383X - 0.4$
♂ $Y = 0.351X - 0.14$
(a) Variance: $F = 9.1$; $F_{.05}(9, 4) = 6.0$; hypothesis of equal variances rejected

D vs. U ♀ $Y = 0.385X - 0.38$
♂ $Y = 0.395 - 0.53$
(a) Variance: $F = 1.06$; $F_{.05}(7, 3) = 8.89$; hypothesis of equal variances accepted
(b) Slopes: $t = 0.17$; $t_{(.05, 8)} = 1.86$; hypothesis of equal slopes accepted
(c) Intercepts: $t = 20$; $t_{(.05, 9)} = 1.833$; hypothesis of equal intercepts rejected

W vs. H ♀ $Y = 0.556X + 0.334$
♂ $Y = 0.513X + 0.814$
(a) Variance: $F = 5.81$; $F_{.05}(9, 4) = 6.0$; hypothesis of equal variances accepted
(b) Slopes: $t = 0.4$; $t_{(.05, 11)} = 1.796$; hypothesis of equal slopes accepted
(c) Intercepts: $t = 10.5$; $t_{(.05, 12)} = 1.782$; hypothesis of equal intercepts rejected

	D	W	H	U
Stephanoceras itinsae (McLearn, 1927) ♀ [macroconch]	31.0	15.7	9.9	—
	26.4	15.6	9.1	—
	30.0	17.4	10.5	—
	39.0	19.5	15.5	13.0
	27.4	17.5	10.0	14.4
	34.5	19.0	11.9	15.3
	7.0	4.7	2.6	3.0
	33.8	19.6	12.5	14.6
	25.0	12.7	8.2	10.5
	30.0	—	—	12.2
Stephanoceras itinsae (McLearn, 1927) ♂ [microconch]	24.8	13.9	8.6	11.2
	15.0	9.0	5.0	—
	18.7	10.0	5.5	7.3
	21.2	13.4	7.4	9.1
	36.2	20.5	10.7	17.0
	40.6	18.1	13.3	18.2
	30.5	12.4	9.7	14.0
	5.0	2.9	1.5	2.2
	10.0	5.7	3.3	4.2
	29.4	17.2	10.2	13.2

Regression Analysis for *Stephanoceras itinsae*

D vs. W ♀ $Y = 0.496X + 1.75$
♂ $Y = 0.459 + 1.67$
(a) Variance: $F = 1.95$; $F_{.05}(9, 8) = 3.39$; hypothesis of equal variances accepted
(b) Slopes: $t = -0.445$; $t_{(.05, 10)} = 1.753$; hypothesis of equal slopes accepted
(c) Intercepts: $t = 4.76$; $t_{(.05, 10)} = 1.746$; hypothesis of equal intercepts rejected

D vs. H ♀ $Y = 0.378X - 0.64$
♂ $Y = 0.319X + 0.13$
(a) Variance: $F = 2.17$; $F_{.05}(8, 9) = 3.23$; hypothesis of equal variances accepted
(b) Slopes: $t = 1.79$; $t_{(.05, 15)} = 1.753$; hypothesis of equal slopes rejected
(c) Intercepts: $t = 9.02$; $t_{(.05, 15)} = 1.753$; hypothesis of equal intercepts rejected

D vs. U ♀ $Y = 0.369X + 1.49$
♂ $Y = 0.469X - 0.58$
(a) Variance: $F = 16.6$; $F_{.05}(6, 8) = 3.58$; hypothesis of equal variances rejected

W vs. H ♀ $Y = 0.706X - 1.09$
♂ $Y = 0.621X - 0.13$
(a) Variance: $F = 1.1$; $F_{.05}(9, 8) = 3.39$; hypothesis of equal variances accepted
(b) Slopes: $t = 0.016$; $t_{(.05, 15)} = 1.753$; hypothesis of equal slopes accepted
(c) Intercepts: $t = 1.5$; $t_{(.05, 15)} = 1.753$; hypothesis of equal intercepts accepted

	D	W	H	U
Stephanoceras skidegatense (Whiteaves, 1876) ♀ [macroconch]	43.5	18.5	16.3	18.2
	20.6	10.5	7.3	7.5
	33.7	19.2	13.2	13.7
	32.7	16.0	11.7	16.4
	15.0	9.0	4.4	7.5
	35.3	19.8	12.2	13.7
	30.5	14.3	9.3	11.4
Stephanoceras skidegatense (Whiteaves, 1876) ♂ [microconch]	36.5	23.0	14.0	15.0
	14.2	7.6	4.2	—
	43.4	—	—	18.6

Regression Analysis for *Stephanoceras skidegatense*

D vs. W ♀ $Y = 0.41X + 3.0$
♂ $Y = 0.69X - 2.2$
(a) Variance: $F = 3.8$; $F_{.05}(6, 1) = 234.0$; accept hypothesis of equal variances
(b) Slopes: $t = 1.89$; $t_{(.05, 5)} = 2.015$; accept hypothesis of equal slopes
(c) Intercepts: $t = 1.65$; $t_{(.05, 6)} = 1.943$; accept hypothesis of equal intercepts

D vs. H ♀ $Y = 0.41X - 1.64$
♂ $Y = 0.44X - 2.0$
(a) Variance: $F = 0.8$; $F_{.05}(6, 1) = 234.0$; accept hypothesis of equal variances
(b) Slopes: $t = 0.44$; $t_{(.05, 5)} = 2.015$; accept hypothesis of equal slopes
(c) Intercepts: $t = 0.642$; $t_{(.05, 6)} = 1.943$; accept hypothesis of equal intercepts

D vs. U ♀ $Y = 0.40X + 0.46$
♂ $Y = 0.52X - 4.0$
(a) Variance: $F = 4.6$; $F_{.05}(6, 1) = 234.0$; accept hypothesis of equal variance
(b) Slopes: $t = 0.356$; $t_{(.05, 5)} = 2.015$; accept hypothesis of equal slopes
(c) Intercepts: $t = 8.3$; $t_{(.05, 6)} = 1.943$; reject hypothesis of equal intercepts

W vs. H ♀ $Y = 0.84X - 2.28$
♂ $Y = 0.64X - 0.6$
(a) Variance: $F = 3.2$; $F_{.05}(6, 1) = 234.0$; accept hypothesis of equal variances
(b) Slopes: $t = 0.87$; $t_{(.05, 5)} = 2.015$; accept hypothesis of equal slopes
(c) Intercepts: $t = 0.608$; $t_{(.05, 6)} = 1.943$; accept hypothesis of equal intercepts

	D	W	H
Chondroceras oblatum (Whiteaves, 1876) ♀ [macroconch]	37.0	30.0	20.9
	10.0	8.3	5.7
	6.2	4.3	3.3
	23.4	19.2	11.7
"*Saxitoniceras marshalli*" McLearn, 1927 ♀ [= *C. oblatum*] [macroconch]	34.3	34.0	17.5
	40.8	28.7	20.8
	35.6	29.2	19.8
	24.8	19.3	12.0
	2.6	1.7	1.0
	31.8	28.8	14.0
	15.4	12.2	7.5
	10.9	8.8	5.7
	7.0	5.0	3.5
	5.4	4.2	2.8

Regression Analysis for *Chondroceras oblatum* and "*Saxitoniceras marshalli*"

D vs. W *(C. oblatum)* $Y = 0.825X - 0.35$
(S. marshalli) $Y = 0.839X - 0.32$
(a) Variance: $F = 26.9$; $F_{.05}(9, 3) = 8.84$; reject hypothesis of equal variances

D vs. H *(C. oblatum)* $Y = 0.557X - 0.27$
(S. marshalli) $Y = 0.51X - 0.22$
(a) Variance: $F = 0.944$; $F_{.05}(9, 3) = 8.84$; accept hypothesis of equal variances
(b) Slopes: $t = 1.02$; $t_{(.05, 10)} = 1.812$; accept hypothesis of equal slopes
(c) Intercepts: $t = 1.33$; $t_{(.05, 11)} = 1.796$; accept hypothesis of equal intercepts

W vs. H *(C. oblatum)* $Y = 0.674X - 0.013$
(S. marshalli) $Y = 0.577X + 0.54$
(a) Variance: $F = 4.18$; $F_{.05}(9, 3) = 8.84$; accept hypothesis of equal variances
(b) Slopes: $t = 0.86$; $t_{(.05, 10)} = 1.812$; accept hypothesis of equal slopes
(c) Intercepts: $t = 0.817$; $t_{(.05, 11)} = 1.796$; accept hypothesis of equal intercepts

	D	W	H
"*Saxitoniceras marshalli*" McLearn, 1927 ♀ [= *C. oblatum*] [macroconch]	34.3	34.0	17.5
	40.8	28.7	20.8
	35.6	29.3	19.8
	24.8	19.3	12.0
	2.6	1.7	1.0
	31.8	18.8	14.0
	15.4	12.2	7.5
	10.9	8.8	5.7
	7.0	5.0	3.5
	5.4	4.2	2.8
Chondroceras allani (McLearn, 1927) ♀ [macroconch]	26.0	17.6	13.8
	37.0	24.9	18.2
	12.8	8.9	7.0
	18.7	15.0	9.4
	10.0	7.5	6.4
	4.5	2.9	2.3
	7.5	5.1	3.8
	25.0	17.2	12.4

Regression Analysis for "*Saxitoniceras marshalli*" and *Chondroceras allani*

D vs. W *(S. marshalli)* $Y = 0.839X - 0.32$
(C. allani) $Y = 0.673X + 0.487$
(a) Variance: $F = 4.96$; $F_{.05}(9, 7) = 3.68$; reject hypothesis of equal variances

D vs. H *(S. marshalli)* $Y = 0.51X - 0.22$
(C. allani) $Y = 0.483X + 0.62$
(a) Variance: $F = 0.71$; $F_{.05}(9, 7) = 3.68$; accept hypothesis of equal variances
(b) Slopes: $t = 0.16$; $t_{(.05, 14)} = 1.761$; accept hypothesis of equal slopes
(c) Intercepts: $t = 0.578$; $t_{(.05, 15)} = 1.753$; accept hypothesis of equal intercepts

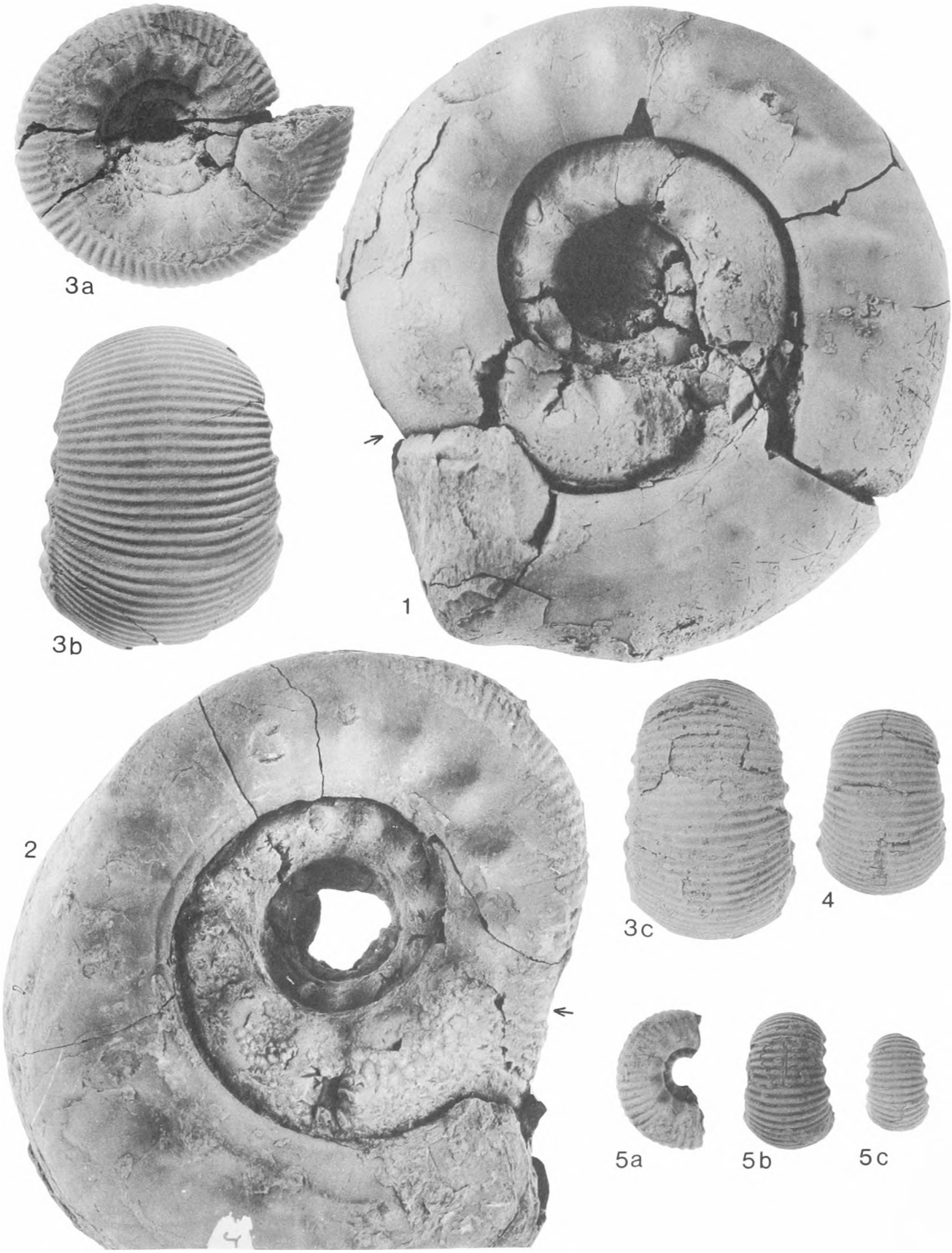
W vs. H *(S. marshalli)* $Y = 0.577X + 0.54$
(C. allani) $Y = 0.71X + 0.37$
(a) Variance: $F = 2.93$; $F_{.50}(9, 7) = 2.93$; accept hypothesis of equal variances
(b) Slopes: $t = 1.37$; $t_{(.05, 14)} = 1.761$; accept hypothesis of equal slopes
(c) Intercepts: $t = 1.84$; $t_{(.05, 15)} = 1.753$; reject hypothesis of equal intercepts

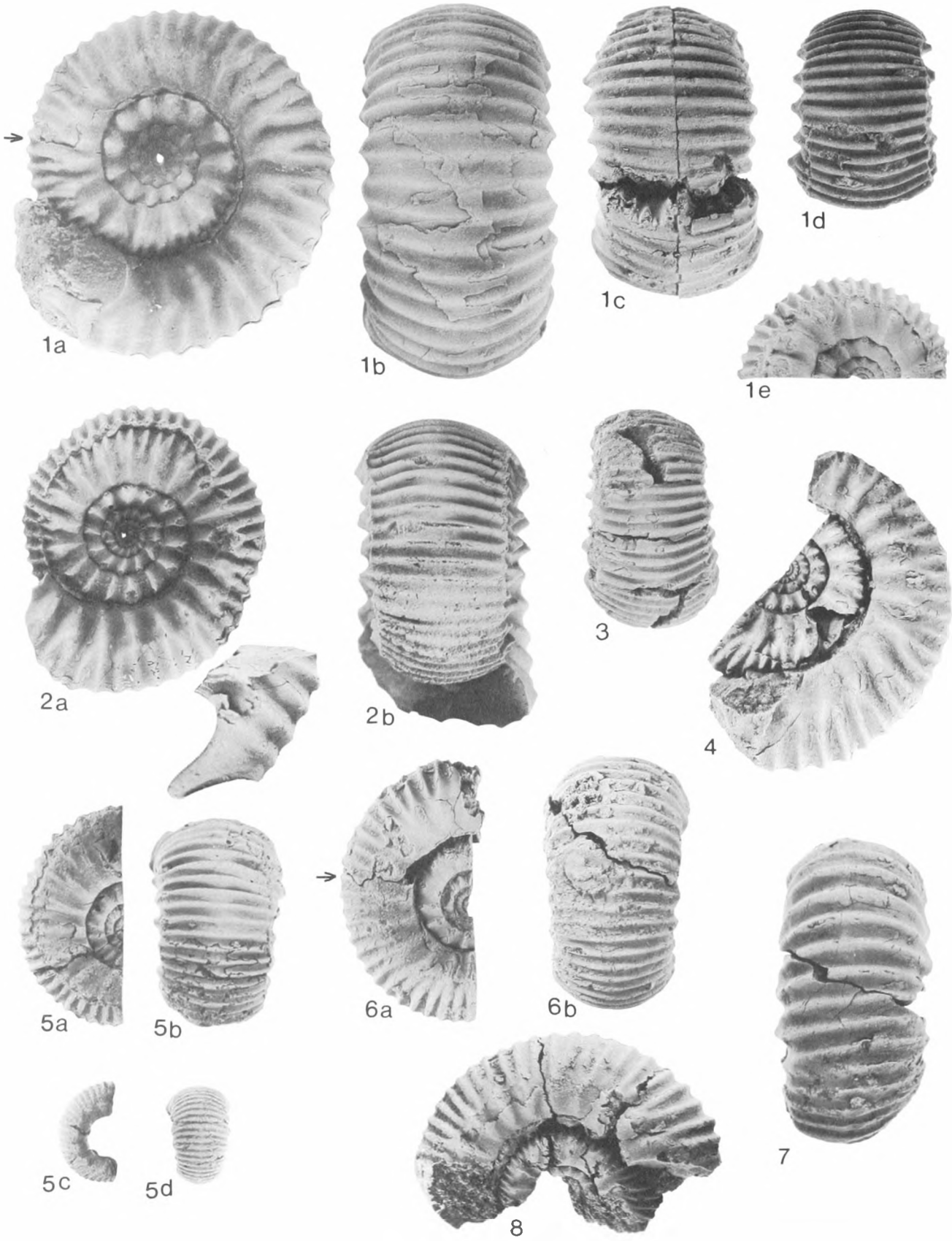
PLATES

EXPLANATION OF PLATE 1

(All figures natural size)

Figure		Page
1-5.	**Zemistephanus richardsoni** (Whiteaves) ♀	25
1.	Complete specimen with aperture, McM J1797a, from bed *a* of Yakoun Formation at Mackenzie Bay, Queen Charlotte Islands. Note egression of body chamber, rounded nodes and deep umbilicus.	
2.	Almost complete specimen, McM J1797d, from same bed as specimen shown in figure 1. Note egression of body chamber, loss of ribbing on body chamber, deep umbilicus and nodes of phragmocone.	
3a-c.	Phragmocone, McM J1797b, from same bed as specimens shown in figures 1 and 2.	
4-5.	Incomplete phragmocones and nuclei, from Fitz Creek Siltstone, southern Alaska (USGS Mesozoic. loc. 2999.3).	


3a
3b
1
2
3c
4
5a
5b
5c


1a
1b
1c
1d
1e
2a
2b
3
4
5a
5b
5c
5d
6a
6b
7
8

EXPLANATION OF PLATE 2

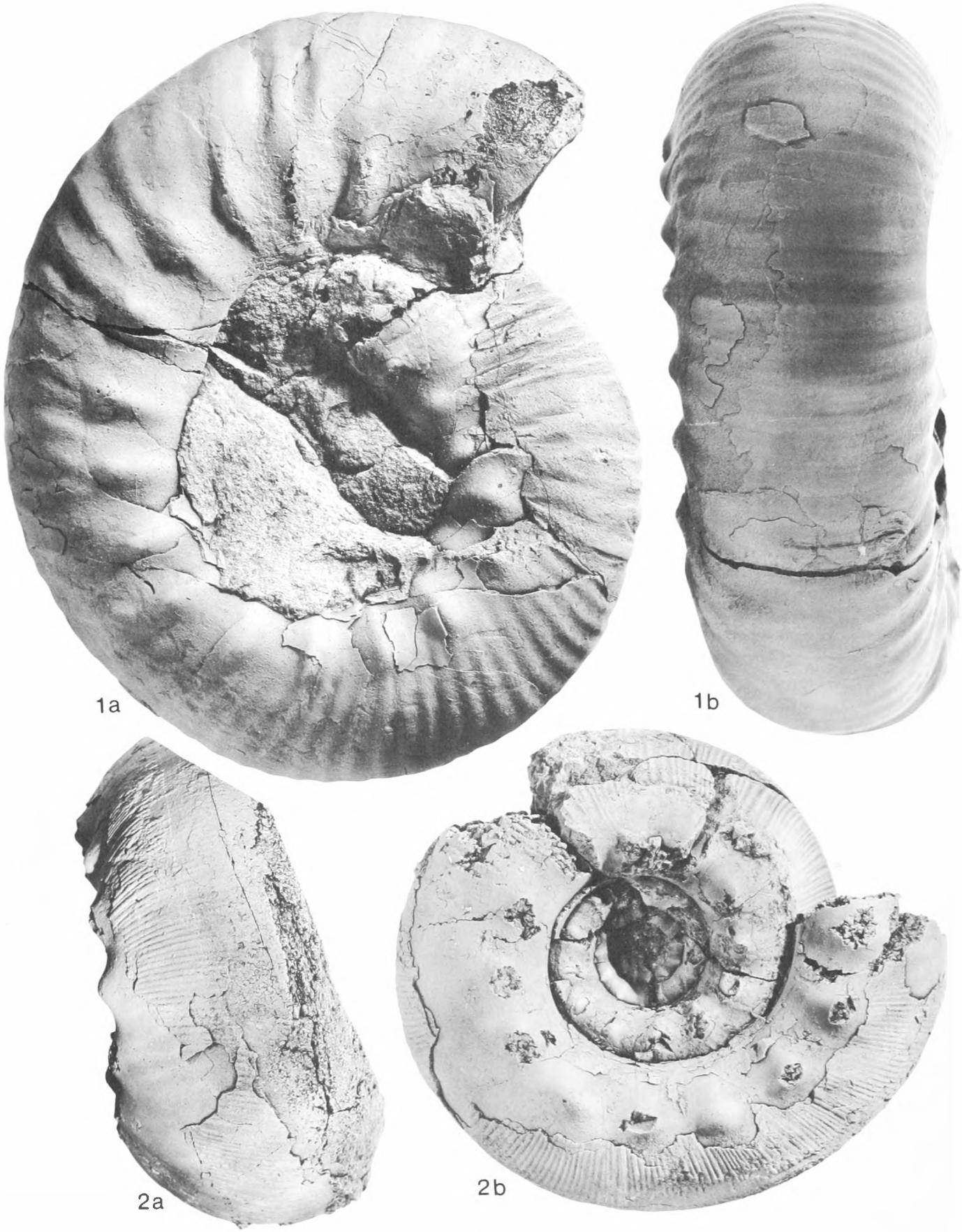
(All figures natural size)

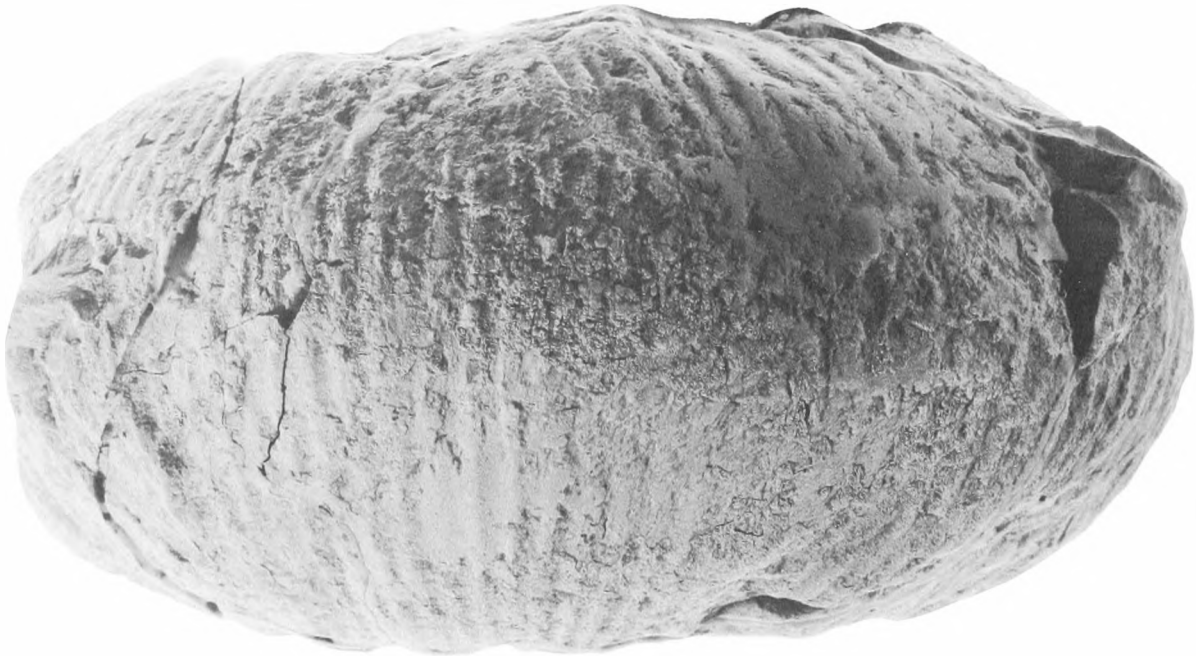
Figure		Page
1-4.	**Zemistephanus richardsoni** (Whiteaves) ♂	25
	1a-e. Allotype, complete specimen with lappet, McM J1796a, from bed *a* of Yakoun Formation at Mackenzie Bay, Queen Charlotte Islands; c-e, phragmocone whorls. Note strong egression of body chamber, straight primaries and large conical nodes on phragmocone, and wider spacing of secondary ribbing on body chamber.	
	2a-b. Phragmocone with body chamber fragment bearing lappets. GSC 56686, from the Yakoun Formation at Mackenzie Bay, Queen Charlotte Islands, a, aperture located in approximate position. Note large conical nodes and prominent ribbing.	
	3. Phragmocone, McM J1796b, from same bed as specimens shown in figure 1.	
	4. Incomplete phragmocone with body chamber bearing lappets, McM J1796c, from same bed as specimens shown in figures 1 and 3.	
5-8.	**Zemistephanus crickmayi** (McLearn) ♂	30
	5a-d. Incomplete phragmocone with its nucleus, from Fitz Creek Siltstone, southern Alaska (USGS Mesozoic loc. 2999.4).	
	6a-b. Incomplete phragmocone with beginning of body chamber, McM J1798a, from bed *a* of Yakoun Formation at Mackenzie Bay, Queen Charlotte Islands.	
	7. Body chamber with prominent ribbing, McM J1798e, from same bed as specimen shown in figure. 1.	
	8. Incomplete phragmocone and body chamber, McM J1798f, from same bed as specimen shown in figure 1.	

EXPLANATION OF PLATE 3

(All figures natural size)

Figure		Page
1a-b.	**Zemistephanus carlottensis** (Whiteaves) ♀ ..... Holotype, from Skidegate Inlet, Queen Charlotte Islands, GSC 5010.	25
2a-b.	**Zemistephanus richardsoni** (Whiteaves) ♀ ..... Holotype, from Skidegate Inlet, Queen Charlotte Islands, GSC 5013.	25


1a
1b
2a
2b



EXPLANATION OF PLATE 4

(All figures natural size)

Page

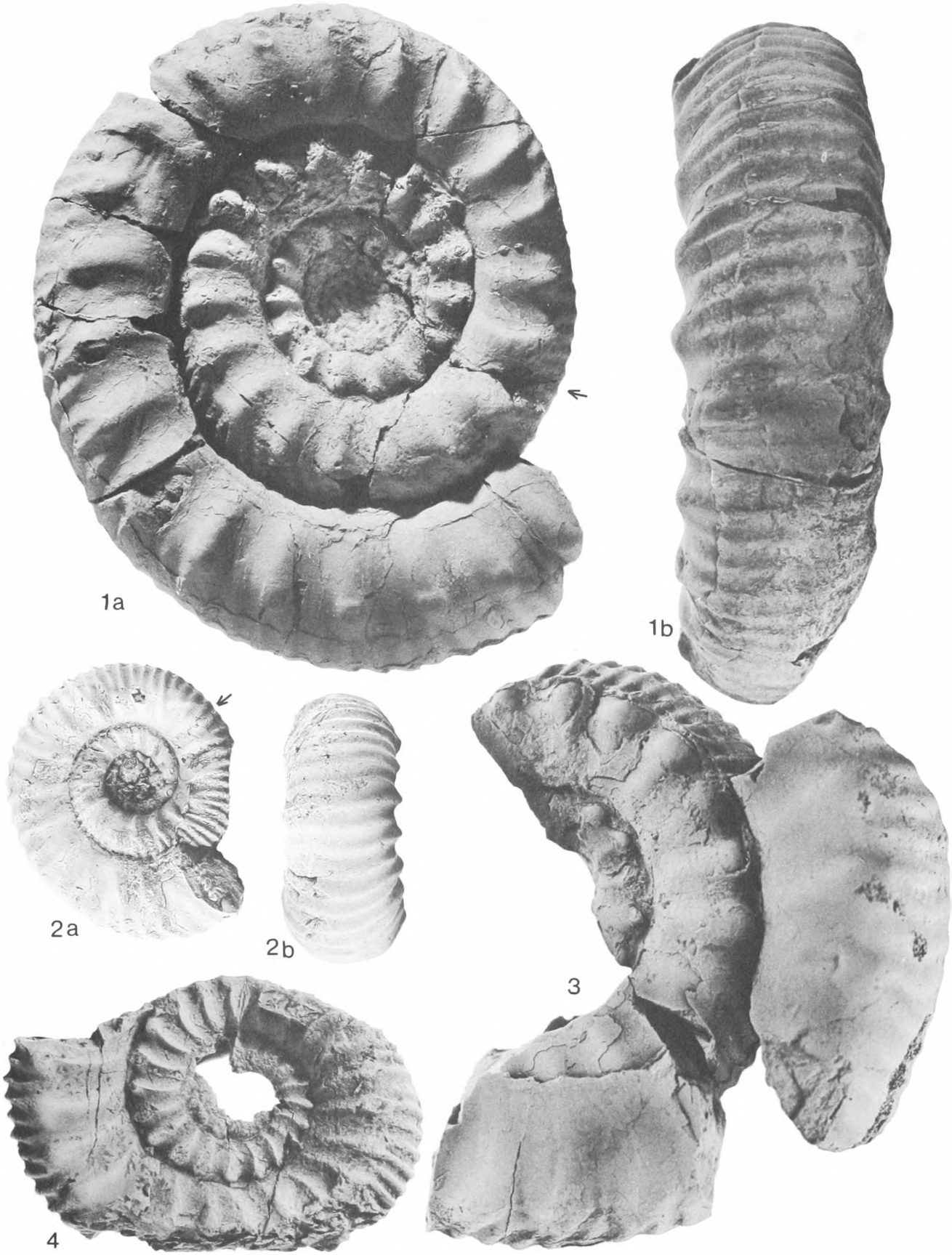
Zemistephanus warreni (McLearn) ♀ 34

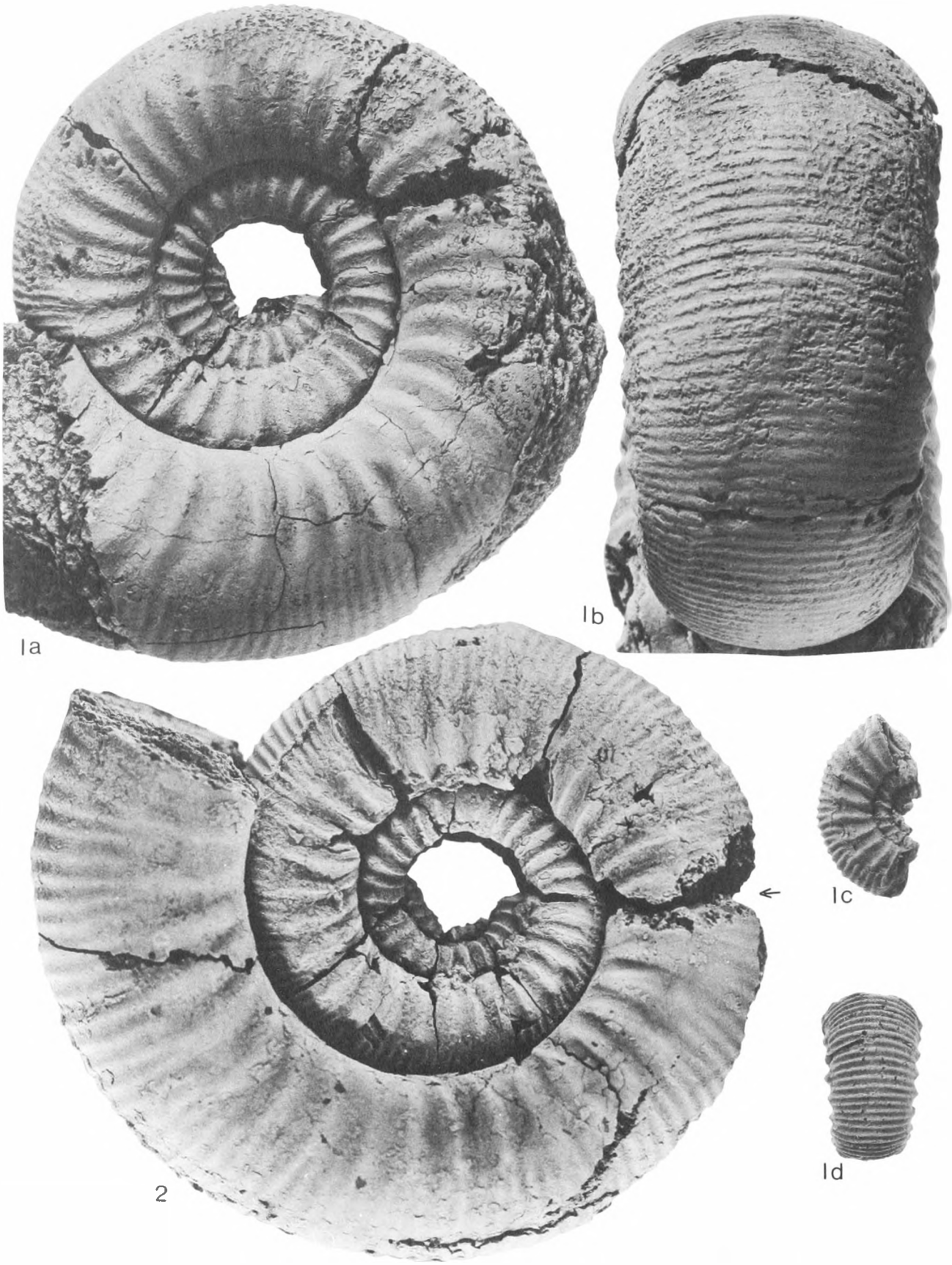
Holotype, almost completely septate, from Porcupine Creek, southern Alberta, UA JR114.

EXPLANATION OF PLATE 5

(Figures natural size unless otherwise indicated)

Figure		Page
1a, b.	**Zemistephanus alaskensis** n. sp. ♀ Almost complete specimen with one whorl body chamber, McM J1858a, from bed *c* of the Yakoun Formation at Mackenzie Bay, Queen Charlotte Islands. Note large conical nodes on phragmocone and broad blunt primaries. × 0.75.	32
2.	**Zemistephanus crickmayi** (McLearn) ♂ Holotype, almost complete specimen, GSC 9016, from Yakoun Formation of Maude Island, Queen Charlotte Islands.	30
3-4.	**Stephanoceras (Stemmatoceras)** ex gr. **S. acuticostatum** (Weisert) 3. Lectotype of "*Teloceras itinsae*" McLearn (*nom. dub.*), incomplete phragmocone and body chamber, GSC 56687, from same bed as specimen shown in figure 2. Note broad, flat venter. 4. Damaged phragmocone, McM J1861, from Yakoun Formation of Mackenzie Bay, Queen Charlotte Islands. Note broad, flat venter, lateral shoulder, and sharp primaries on inner whorls.	49


1a
1b
2a
2b
3
4


1a
1b
1c
1d
2

EXPLANATION OF PLATE 6

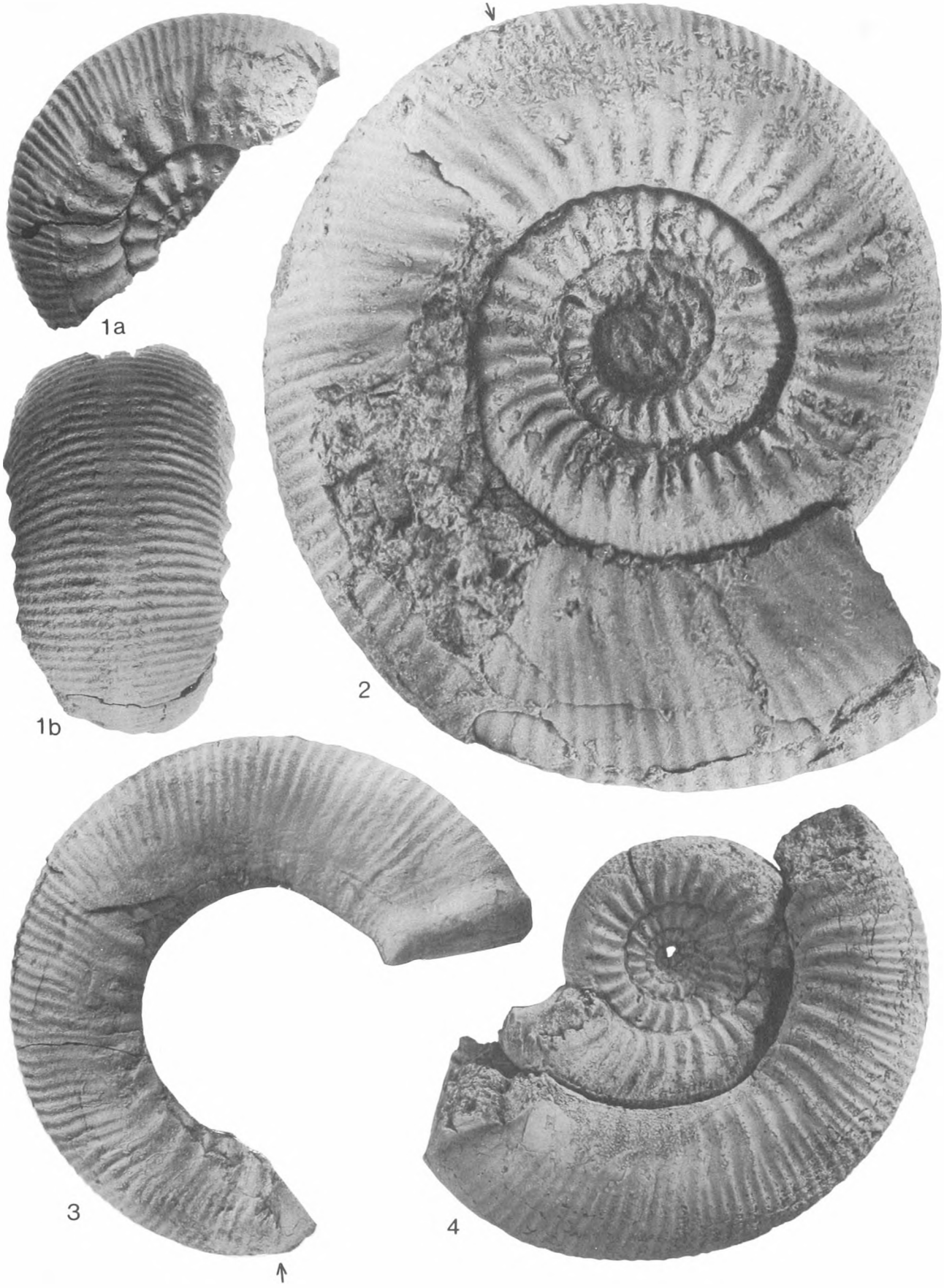
(All figures natural size)

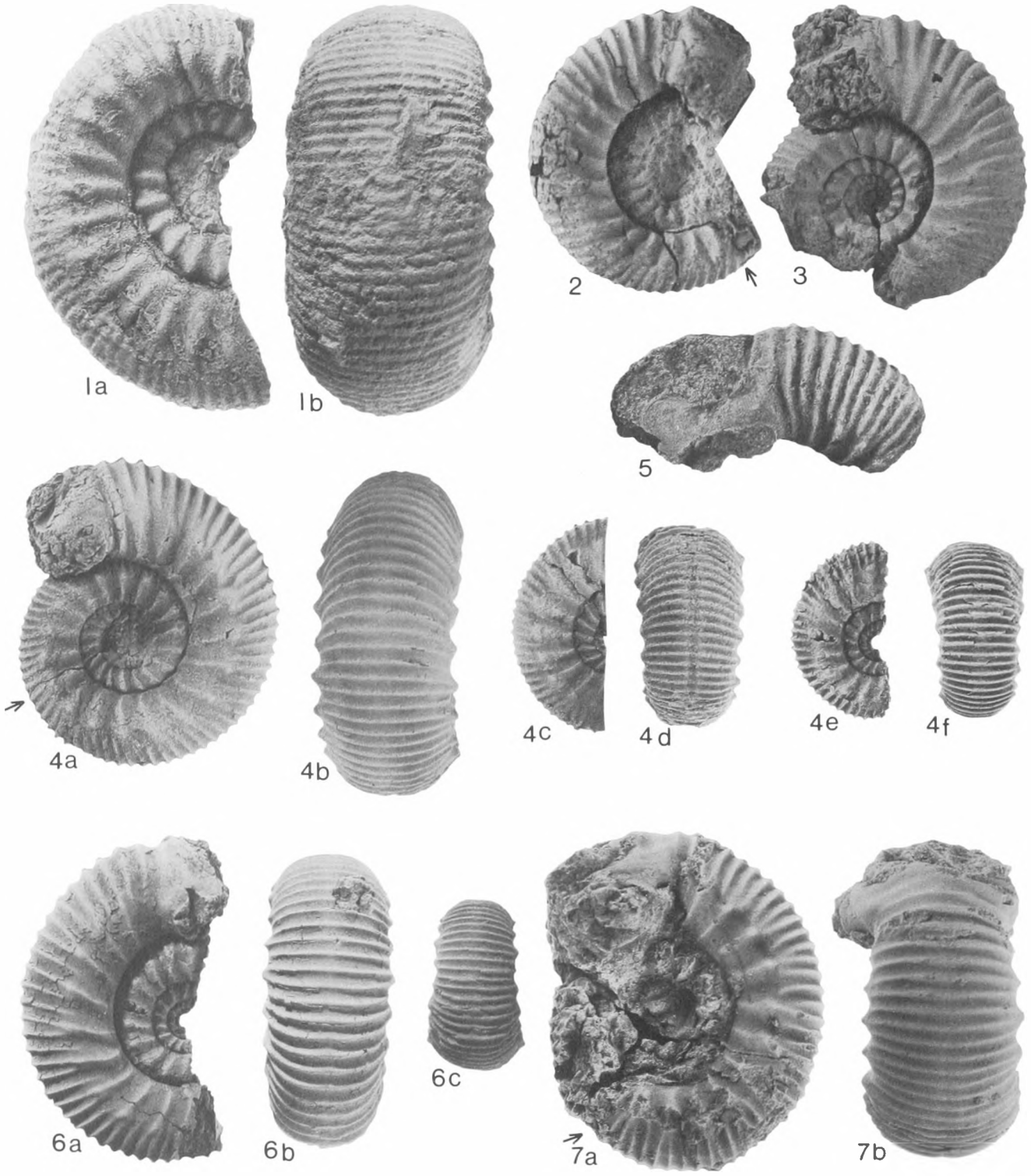
Figure		Page
1-2.	**Stephanoceras (Stephanoceras) itinsae** (McLearn) ♀	37
	1a-d. Phragmocone and its nucleus, UBC collections, unnumbered, from the Yakoun Formation of "Harty's Island", Queen Charlotte Islands.	
	2. Phragmocone and half whorl of body chamber, GSC 56688, from the Yakoun Formation of Reef Island, Queen Charlotte Islands. Note egression of body chamber.	

EXPLANATION OF PLATE 7

(Figures natural size unless otherwise indicated)

Figure		Page
1-4.	**Stephanoceras (Stephanoceras) itinsae** (McLearn) ♀	37
1a, b.	Incomplete phragmocone, McM J1808g, from the Yakoun Formation of South Balch Island, eastern shore, Queen Charlotte Islands.	
2.	Phragmocone with half whorl of body chamber, GSC 56689, from the Yakoun Formation of Reef Island, Queen Charlotte Islands. Note egression of body chamber.	
3.	Complete body chamber with aperture, GSC 56690, from Yakoun Formation of South Balch Island, Queen Charlotte Islands, × 0.5.	
4.	Almost complete specimen with aperture, lateral view, GSC 56691, from same locality as specimen shown in figure 3. × 0.5.	


1a
1b
2
3
4


1a
1b
2
3
5
4a
4b
4c
4d
4e
4f
6a
6b
6c
7a
7b

EXPLANATION OF PLATE 8

(All figures natural size)

Figure		Page
1-7.	**Stephanoceras (Stephanoceras) itinsae** (McLearn)	37

1a, b. ♀ [macroconch]
Incomplete phragmocone, McM J1880, from Rock Creek Member of Fernie Group at Ribbon Creek, southern Alberta.

2-7. ♂ [microconch: **'Itinsaites'**]

2-3. Damaged specimens with peristome, bearing lappets, GSC 56692-3, from Yakoun Formation of South Balch Island, Queen Charlotte Islands.

4a-f. Complete specimen with lappet, entire specimen and parts of phragmocone, McM J1799a, from same locality as specimen shown in figures 2-3.

5. Fragment with well-preserved lappet, McM J1801b, from interval *d* at same locality as that from which the specimens shown in figures 2-4 were collected.

6a-c. Damaged specimen, McM J1800, with aperture, body chamber and phragmocone, from bed *c* at same locality as that from which the specimens shown in figures 2-5 were collected.

7a, b. Damaged specimen with body chamber, McM J1838, from Rock Creek Member of Fernie Group at Ribbon Creek, southern Alberta.

EXPLANATION OF PLATE 9

(Figures natural size unless otherwise indicated)

Figure — Page

1-5. **Stephanoceras (Stephanoceras) skidegatense** (Whiteaves) **43**

1a-b. ♀ [macroconch]
Incomplete specimen with part of body chamber, lateral and ventral views, McM J1878, × 0.7, from bed *d* of Yakoun Formation at Richardson Bay, Queen Charlotte Islands. Note prominent sharp secondaries, with only two per primary on body chamber.

2-3. ♂ [microconch] **45**

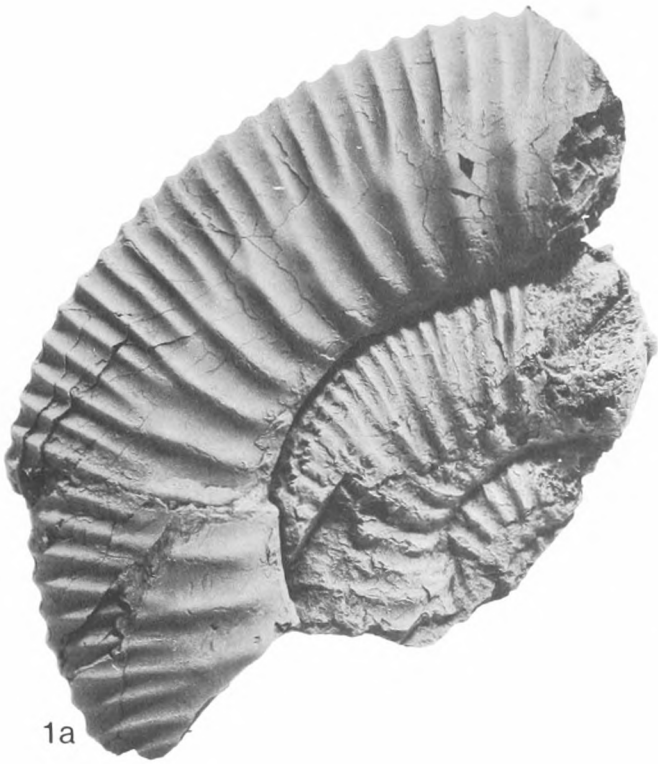
2a, b. Almost complete paratype of Whiteaves, GSC 56694, purchased by Dawson on Queen Charlotte Islands.

3a-c. Allotype, complete specimen with base of lappet: a,b, entire specimen and c, phragmocone, McM J1802b, collected from same bed as specimen shown in figure 1.

4-5. Inner whorls of juveniles (?) × 2.

4a, b. McM J1802f, from same bed as that from which specimens shown in figures 1 and 3 were collected.

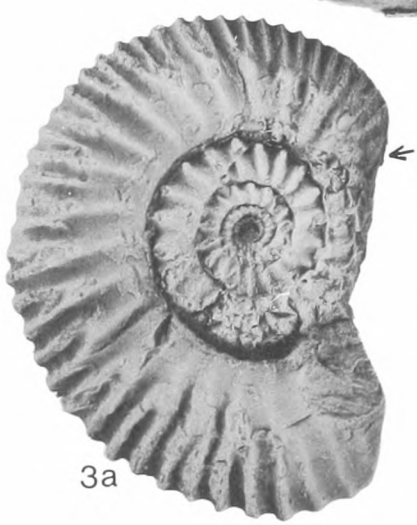
5a, b. GSC 56695, from Yakoun Formation, "between MacKenzie Bay and Clapp Bay", Queen Charlotte Islands.


1a


1b


2a


2b


3a


3b

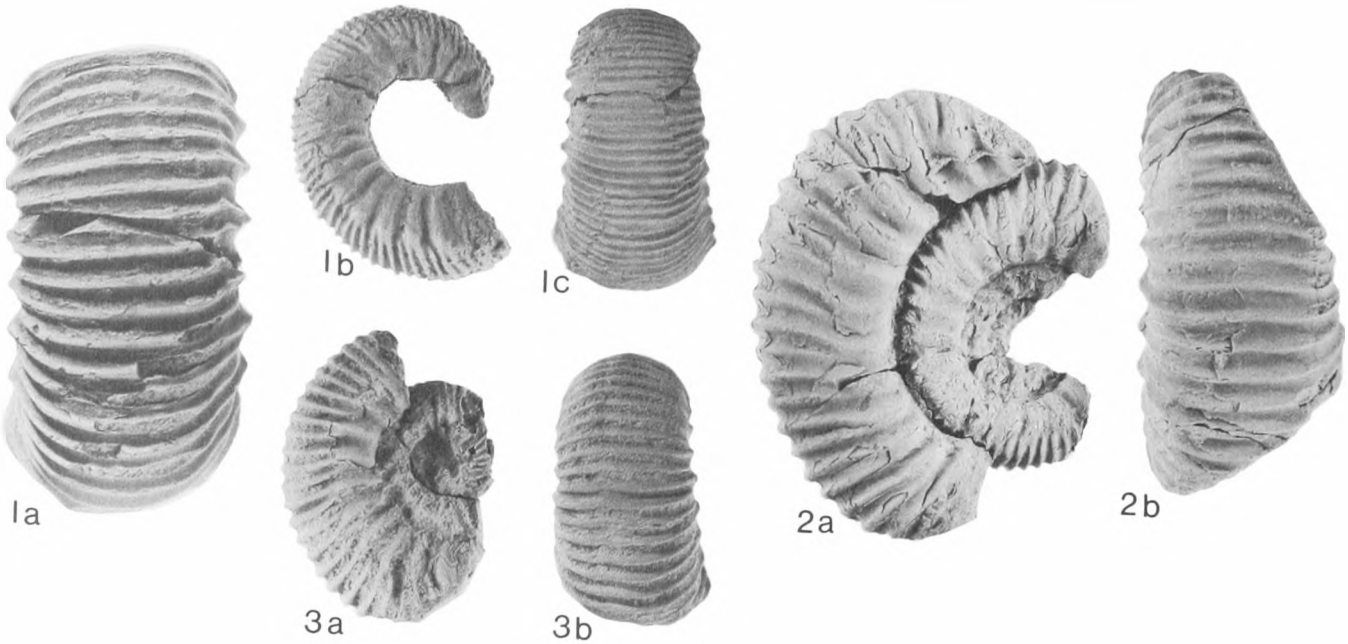

4a


4b


5a


5b


3c


1a
1b
1c
2a
2b
3a
3b


4a


4b

EXPLANATION OF PLATE 10

(All figures natural size)

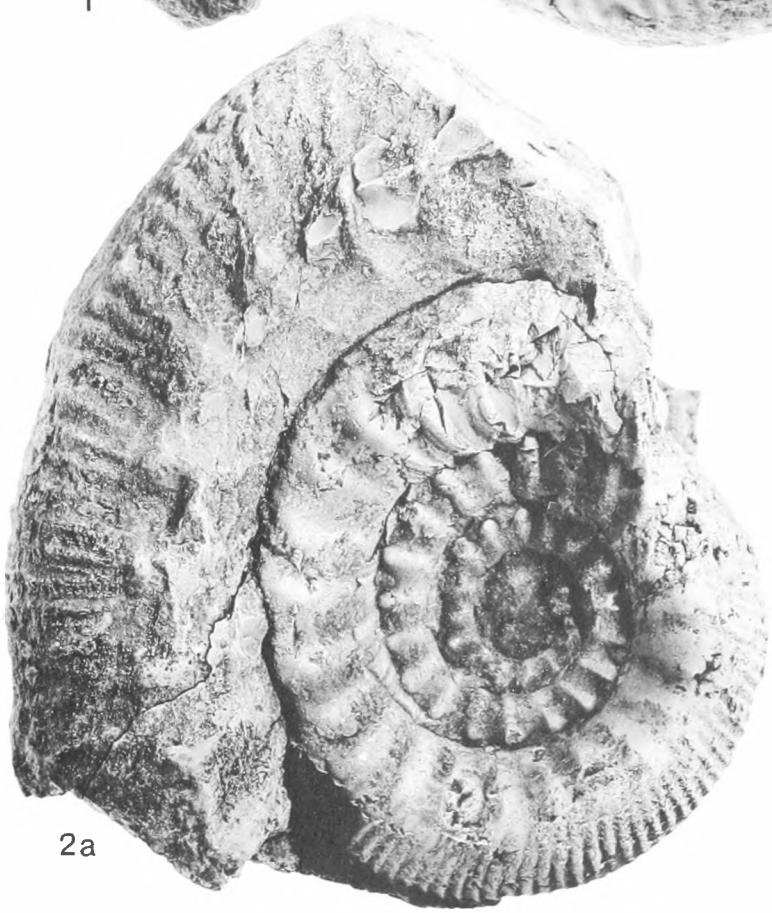
Figure		Page
1a-c.	**Stephanoceras skidegatense** (Whiteaves) ♂ ..... a, body chamber, b and c, phragmocone, McM J1802a, from bed *d* of the Yakoun Formation at Richardson Bay, Queen Charlotte Islands.	43
2-3.	**Stephanoceras** sp. ♂ aff. **S. skidegatense** (Whiteaves) ..... From Yakoun Formation at Richardson Bay, Queen Charlotte Islands. 2a, b. Phragmocone with part of body chamber, McM J1804, from bed *c*. 3a, b. Phragmocone, McM J1806, from bed *b*.	47
4a-b.	**Stephanoceras (Stephanoceras) pyritosum caamanoi** McLearn ♀ ..... Holotype, almost completely septate, GSC 9056, from Yakoun Formation of South Balch Island, Queen Charlotte Islands.	47

EXPLANATION OF PLATE 11

(All figures natural size)

Figure		Page
1.	**"Stemmatoceras mclearni"** Warren [= **Stephanoceras (S.) itinsae** McLearn ♀] ..... Holotype, septate almost to the end, UA JR192, float in Miners Creek near Cadomin, Alberta.	37
2a, b.	**Stephanoceras (Stemmatoceras) dowlingi** (McLearn) ♀ ..... Holotype, with part of body chamber, GSC 9050, from (?) Rock Creek Member, Fernie Group, at Ribbon Creek, southern Alberta.	48

1



2a

2b

1

2

EXPLANATION OF PLATE 12

(Figures natural size unless otherwise indicated)

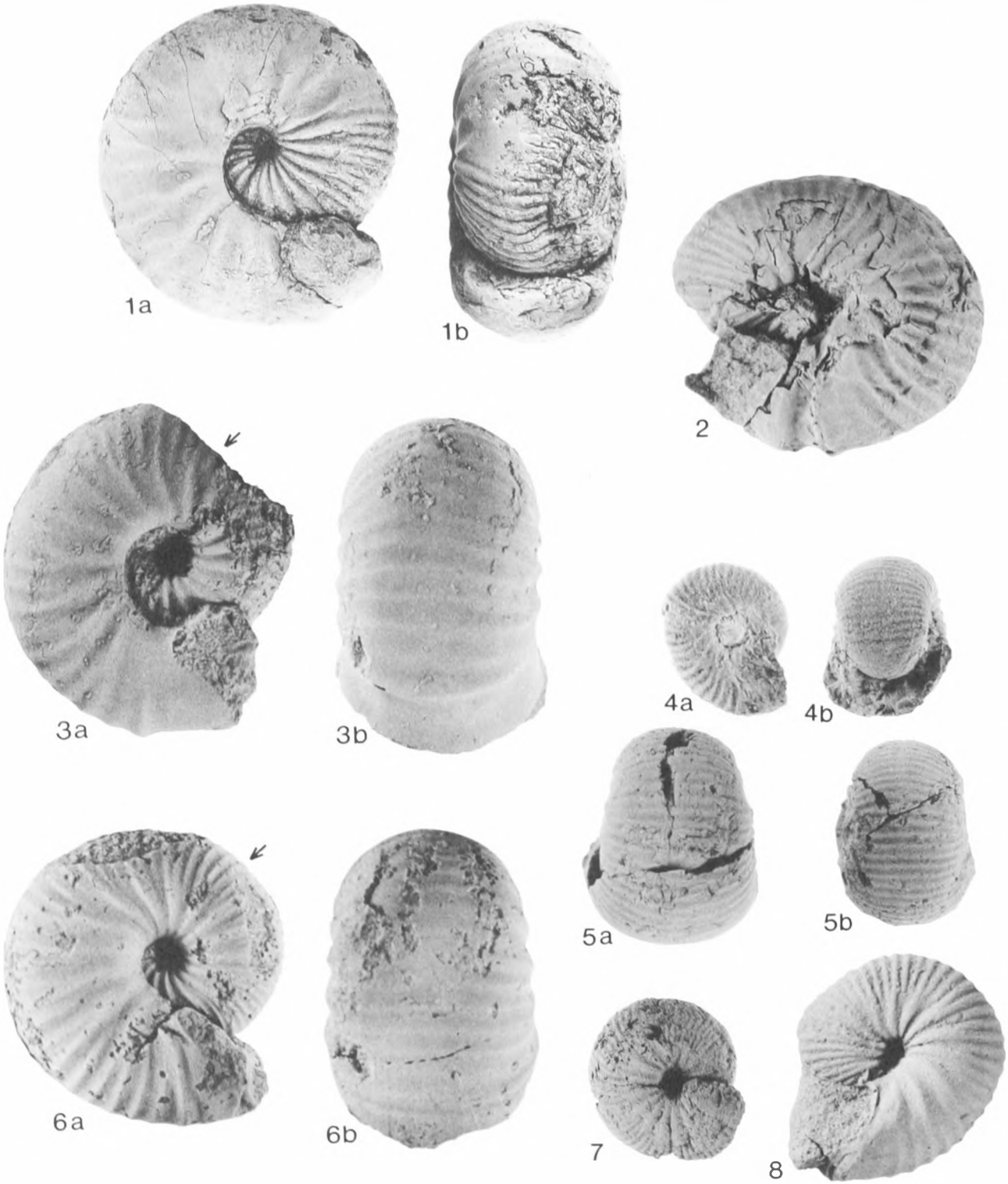
Figure		Page
1.	**Stephanoceras (Stemmatoceras?) palliseri** (Warren) ♀ Holotype, UA JR115, from the Fernie Group, upper Whitehorse River, Mountain Park area, Alberta.	49
2.	**Stephanoceras (Stemmatoceras?) allani** (Warren) ♀ Holotype, almost complete with 3/4 whorl body chamber, UA JR479, from the Rock Creek Member, Fernie Group, southern Alberta, × 0.6.	48

EXPLANATION OF PLATE 13

(Figures natural size unless otherwise indicated)

Figure		Page
1.	**Stephanoceras (Stemmatoceras?) albertense** (McLearn) ♀ ..... Holotype, UA JR113, from the Fernie Group, headwaters of Sheep Creek, Alberta.	49
2.	**"Stemmatoceras carri"** Warren [= **Stephanoceras (S.) itinsae** ♀] ..... Lectotype (here designated), UA JR484, from the Rock Creek Member of the Fernie Group, six miles above the mouth of Whitehorse River near Cadomin, Alberta, × 0.8.	40


2
1


1a
1b
2
3a
3b
4a
4b
5a
5b
6a
6b
7
8

EXPLANATION OF PLATE 14

(Figures natural size unless otherwise indicated)

Figure		Page
1-7.	**Chondroceras oblatum** (Whiteaves) ♀ [macroconch]	52
1a, b.	Holotype, complete, GSC 4964, from Skidegate Channel, Queen Charlotte Islands.	
2.	Distorted complete specimen, McM J1834a, from beds *a* - *b*, of the Yakoun Formation on South Balch Island, Queen Charlotte Islands.	
3a, b.	Complete specimen, McM J1836b, from the Rock Creek Member of the Fernie Group at Ribbon Creek, southern Alberta.	
4a, b.	Phragmocone, McM J1877, from undifferentiated Yakoun Formation of eastern shore of South Balch Island. × 2.	
5a, b.	Phragmocone, McM J1795, collected from same locality as specimen shown in figure 4, × 1.3.	
6a, b.	Complete specimen, McM J1836c, from the Rock Creek Member of the Fernie Group at Ribbon Creek, southern Alberta.	
7.	Phragmocone, McM J1832b, from base of interval *d* of the Yakoun Formation at South Balch Island.	
8.	**Chondroceras oblatum** (Whiteaves) ♂ [microconch]	52
	Allotype, complete specimen, McM J1794a, from undifferentiated Yakoun Formation of South Balch Island, Queen Charlotte Islands, × 2.	

EXPLANATION OF PLATE 15

(Figures natural size unless otherwise indicated)

Figure — Page

1-6. **Chondroceras defontii** (McLearn) 54

1-5. ♀ [macroconch]

1a, b. Holotype, complete, GSC 9009, from Richardson Bay, Queen Charlotte Islands.

2. Complete specimen, McM J1792b, from bed *c* of the Yakoun Formation at Richardson Bay, Queen Charlotte Islands. Note sudden egression of body chamber and apertural constriction.

3. Complete specimen, McM J1829, from bed *a* of same section as that from which the specimen shown in figure 2 was collected.

4a, b. Complete specimen, GSC 56696, from the Yakoun Formation at Richardson Bay.

5. Phragmocone, McM J1792, collected from same locality as the specimen shown in figure 4. Note dense, curved ribbing.

6. ♂ [microconch]

Allotype, complete body chamber, McM J1793a, from bed *d* of the Yakoun Formation at Richardson Bay, × 1.5.

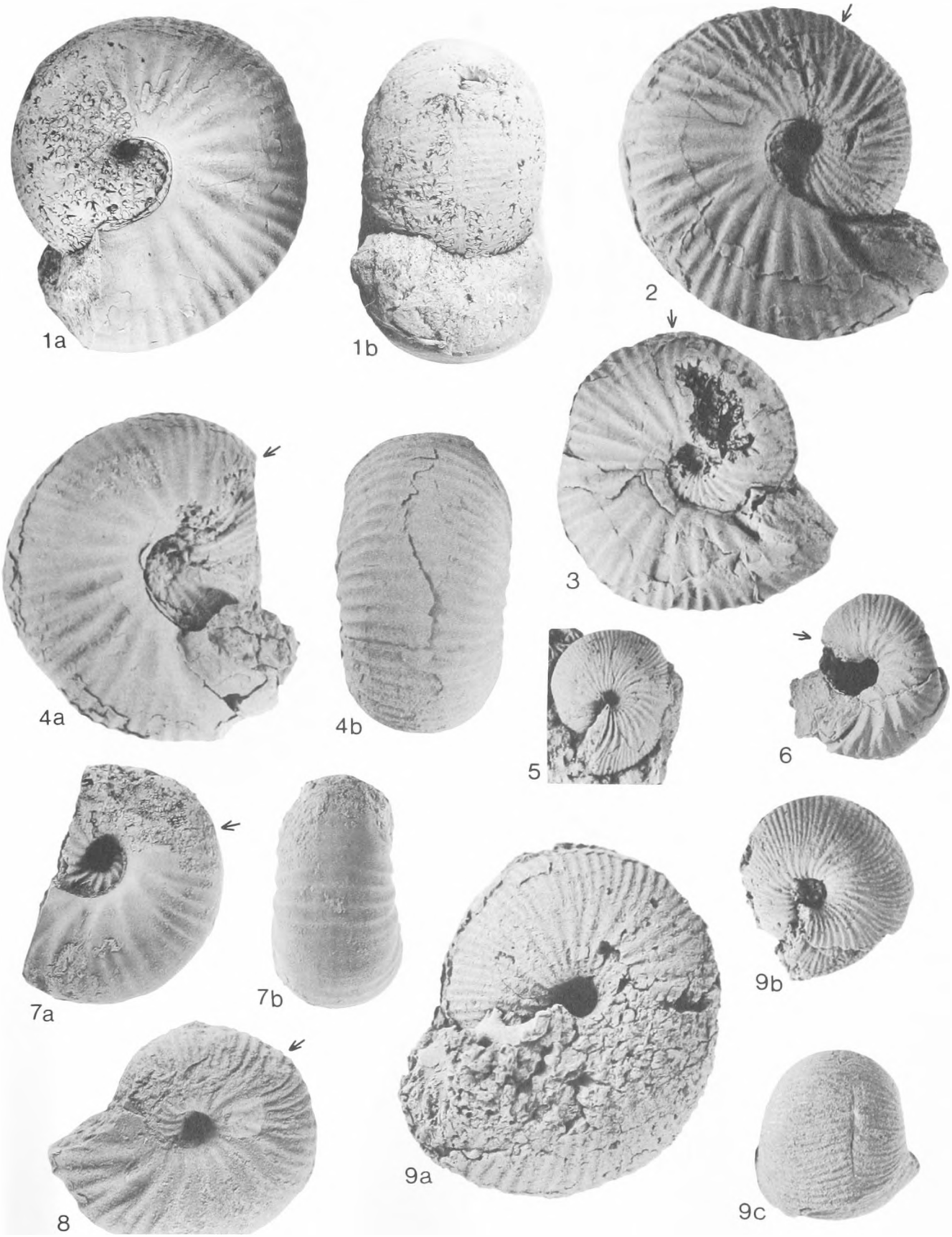
7-8. **Chondroceras allani** (McLearn) ♀ 58

7a, b. Almost complete specimen, McM J1830h, from the Rock Creek Member of the Fernie Group, Ribbon Creek, southern Alberta. Note curved primaries and egression of body chamber.

8. Almost complete specimen, McM J1831d, collected from the same locality as specimen shown in figure 7. Note curved primaries and egression of body chamber.

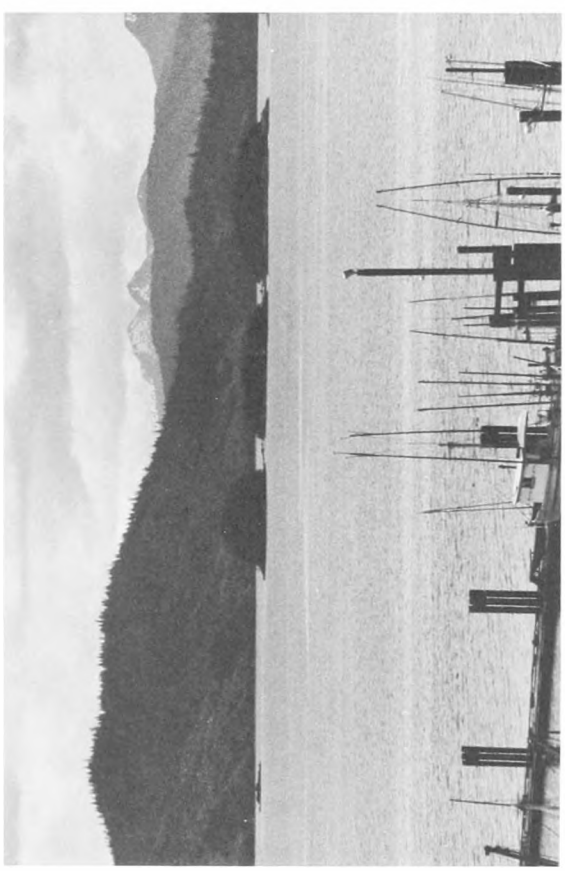
9a-c. **Chondroceras** sp. indet. ♀ 60

Phragmocone, McM J1857a, from bed *a* of the Yakoun Formation at Mackenzie Bay, Queen Charlotte Islands; b-c, inner whorls with strongly curved ribbing, × 2.

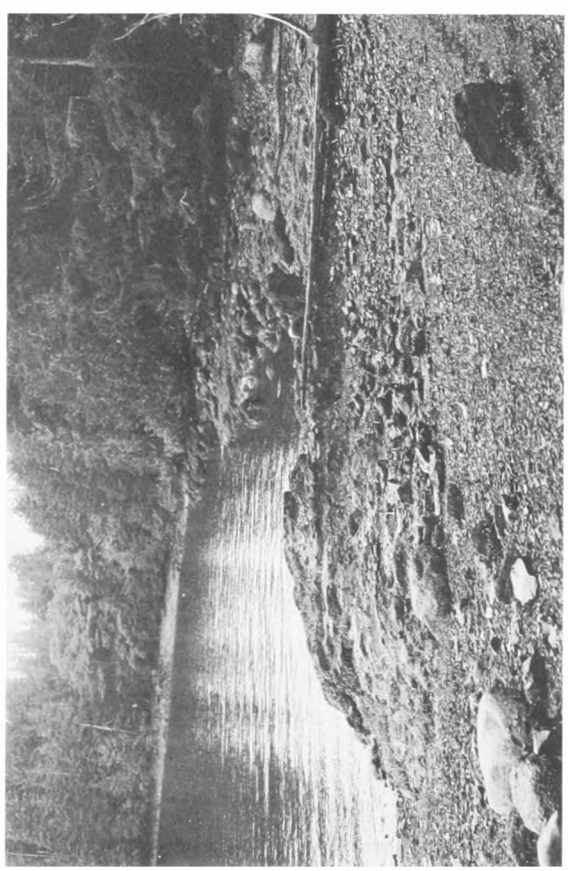

1a
1b
2
3
4a
4b
5
6
7a
7b
8
9a
9b
9c

2

4



1



3

EXPLANATION OF PLATE 16

Figure		Page
1.	General view southward across Skidegate Inlet from Queen Charlotte City showing Channel Islands (South Balch Island on the left) with part of Maude Island behind	7
2.	Mackenzie Bay, Maude Island; interbedded shales and volcanogenic sandstones of the lower Yakoun Formation; horizon *a*, in which *Zemistephanus richardsoni* (Whiteaves) is abundant	9
3.	Outcrop of lower Yakoun Formation on shore of Richardson Bay, Maude Island at high tide. Interbedded shales and sandstones with abundant *Chondroceras defontii* (McLearn) and *Stephanoceras skidegatense* (Whiteaves)	11
4.	South Balch Island, shales and interbedded sandstones of the lower Yakoun Formation; horizon *c*, yielding *Stephanoceras itinsae* (McLearn) and *Chondroceras oblatum* (Whiteaves)	11

INDEX

Number 52

Note: Page numbers in light face; plate numbers in **bold face** type.

A

B

C

D

E

F

G

H

I

J

K

L

M

N

O

P

Q

R

S

T

U

V

W

Y

Z